



## *Atmospheric Pollution*

### *Lecture 4*

*Sahraei  
Physics Department  
Razi university*

<http://www.razi.ac.ir/sahraei>

## Equation of State

Boyle's Law

$$p \propto 1/V$$

Lifting of a parcel of air

Charles' Law

$$V \propto T$$

Avogadro's Law

$$V \propto n$$

Low T, Low P, High V

Ideal gas law  
(simplified equation of state)

$$p = nR^*T/V = nA/V(R^*/A)T$$

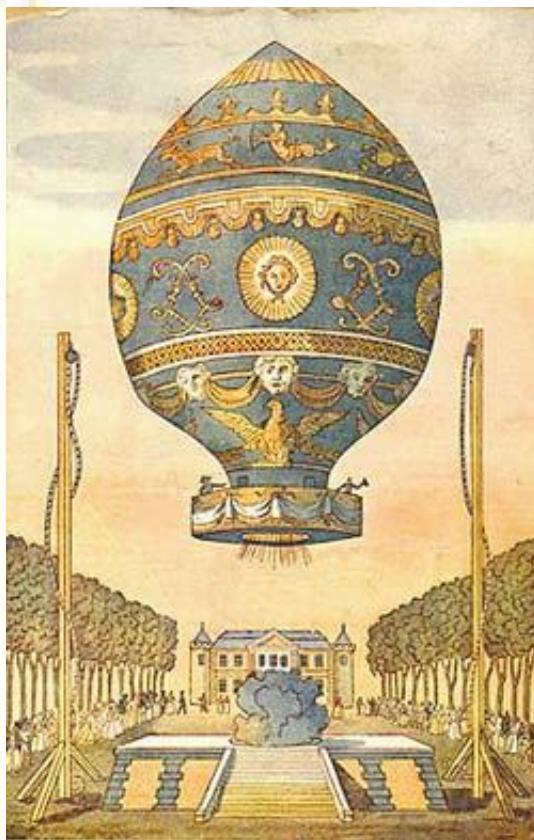
High T, High P, Low V

Number concentration     $N = nA/V$

Boltzmann's constant     $k_B = R^*/A = 1.3807 \times 10^{-19} \text{ cm}^3 \text{ hPa K}^{-1}$

$$P = Nk_B T$$

# Charles Law and Hot Air Balloon



Montgolfier  
Hot Air Balloon  
June 4, 1783  
[www.aps.org](http://www.aps.org)



Charles's  
Hydrogen Balloon  
Aug. 27, 1783  
[www.aps.org](http://www.aps.org)

The country people who saw it fall were frightened and attacked it with stones and knives so that it was much mangled

## Example

$$p = Nk_B T \quad k_B = 1.3807 \times 10^{-19} \text{ cm}^3 \text{ hPa K}^{-1}$$

Earth's surface       $p = 1013 \text{ hPa}$        $T = 288 \text{ K}$

$$N = 2.55 \times 10^{19} \text{ molecules cm}^{-3}$$

48 km altitude       $p = 1 \text{ hPa}$        $T = 270 \text{ K}$

$$N = 2.68 \times 10^{16} \text{ molecules cm}^{-3}$$

## Dalton's Law of Partial Pressure

Total air pressure equals sum of partial pressures of individual gases in the air.

Total atmospheric pressures  $p_a = p_d + p_v$

Equation of state for dry air  $p_d = N_d k_B T$

The number concentration of a gas (molecules per unit volume of air) is an absolute quantity.

The abundance of a gas may also be expressed in terms of a relative quantity,



## volume mixing ratio

the number of gas molecules per molecule of dry air, and expressed for gas  $q$  as

$$\chi_q = \frac{N_q}{N_d} = \frac{p_q}{p_d} \quad (\text{molecules of gas/molecule of dry air})$$

where  $N_q$  and  $p_q$  are the number concentration and partial pressure, respectively, of gas  $q$ .

volume mixing ratios may be multiplied by 100 and expressed as a percentage of dry air volume,

multiplied by  $10^6$  → parts per million volume (ppmv)

multiplied by  $10^9$  → parts per billion volume (ppbv)

multiplied by  $10^{12}$  → parts per trillion volume (pptv)

$$1\% = 0.01 = 10^4 \text{ ppmv}$$

$$1 \text{ ppmv} = 0.000001 = 0.0001\% = 1000 \text{ ppbv} = 10^6 \text{ pptv}$$

### Example

Find the number concentration and partial pressure of ozone if its volume mixing ratio is 0.10 ppmv. Assume  $T = 288 \text{ K}$  and  $p_d = 1013 \text{ mb}$

$$p_d = N_d k_B T \rightarrow N_d = 2.55 \times 10^{19} \text{ molecules cm}^{-3}$$

$$\chi_q = \frac{N_q}{N_d} = \frac{p_q}{p_d} \rightarrow N_q = 0.10 \text{ ppmv} \times 10^{-6} \times 2.55 \times 10^{19} \text{ molecules cm}^{-3}$$
$$N_q = 2.55 \times 10^{12} \text{ molecules cm}^{-3}$$

$$P_q = N_q k_B T = 0.000101 \text{ mb}$$

# Composition of the Air Well-Mixed Gases

Gas	Volume mixing ratio (percent)	(ppmv)
Nitrogen (N <sub>2</sub> )	78.08	780,800
Oxygen (O <sub>2</sub> )	20.95	209,500
Argon (Ar)	0.93	9,300
Neon (Ne)	0.0015	15
Helium (He)	0.0005	5
Krypton (Kr)	0.0001	1
Xenon (Xe)	0.000005	0.05

# Variable Gases

Gas Name	Chemical Formula	Volume Mixing Ratio (ppbv)		
		Clean Troposphere	Polluted Troposphere	Stratosphere
<b>Inorganic</b>				
Water vapor	H <sub>2</sub> O(g)	3000-40(+7)	5.0(+6)-40(+7)	3000-6000
Carbon dioxide	CO <sub>2</sub> (g)	365,000	365,000	365,000
Carbon monoxide	CO(g)	40-200	2000-40,000	10-60
Ozone	O <sub>3</sub> (g)	10-400	10-350	1000-12,000
Sulfur dioxide	SO <sub>2</sub> (g)	0.02-1	1-30	0.01-1
Nitric oxide	NO(g)	0.005-0.1	0.05-300	0.005-10
Nitrogen dioxide	NO <sub>2</sub> (g)	0.01-0.3	0.2-200	0.005-10
CFC-12	CF <sub>2</sub> Cl <sub>2</sub> (g)	0.55	0.55	0.22
<b>Organic</b>				
Methane	CH <sub>4</sub> (g)	1800	1800-2500	150-1700
Ethane	C <sub>2</sub> H <sub>6</sub> (g)	0.2-5	1-50	---
Ethene	C <sub>2</sub> H <sub>4</sub> (g)	0.1-1	1-30	---
Formaldehyde	HCHO(g)	0.1-1	1-200	---
Toluene	C <sub>6</sub> H <sub>5</sub> CH <sub>3</sub> (g)	---	1-30	---
Xylene	C <sub>6</sub> H <sub>4</sub> (CH <sub>3</sub> ) <sub>2</sub> (g)	---	1-30	---
Methyl chloride	CH <sub>3</sub> Cl(g)	0.61	0.61	0.36

## **CHARACTERISTICS OF SELECTED GASES AND AEROSOL PARTICLE COMPONENTS**

gases and aerosol particle components relevant to each of five air pollution problems

### **Pollutants for Different Problems**

each air pollution problem involves a different set of pollutants, although some pollutants are common to two or more problems.

a few gases and aerosol particle components are discussed in terms of their relevance, abundance, sources, sinks, and health effects.

## Indoor air pollution

Gases: NO<sub>2</sub>, CO, HCHO, SO<sub>2</sub>, organic gases, radon

Particles: Black carbon, organic matter, sulfate, nitrate, ammonium, allergens, asbestos, fungal spores, pollen, tobacco smoke

## Outdoor urban air pollution

Gases: O<sub>3</sub>, NO, NO<sub>2</sub>, CO, ethene, toluene, xylene, PAN

Particles: Black carbon, organic matter, sulfate, nitrate, ammonium, soil dust, sea spray, tire particles, lead

## **Acid deposition**

Gases:  $\text{SO}_2$ ,  $\text{H}_2\text{SO}_4$ ,  $\text{NO}_2$ ,  $\text{HNO}_3$ ,  $\text{HCl}$ ,  $\text{CO}_2$

Particles: Sulfate, nitrate, chloride

## **Stratospheric ozone reduction**

Gases:  $\text{O}_3$ ,  $\text{NO}$ ,  $\text{HNO}_3$ ,  $\text{HCl}$ ,  $\text{ClONO}_2$ , chlorofluorocarbons

Particles: chloride, sulfate, nitrate

## Global climate change

Gases:  $H_2O$ ,  $CO_2$ ,  $CH_4$ ,  $N_2O$ ,  $O_3$ , chlorofluorocarbons

Particles: black carbon, organic matter, sulfate, nitrate, ammonium, soil dust, sea spray

# **Carbon Dioxide [CO<sub>2</sub>(g)]**

Colorless, odorless, greenhouse gas

## **Sources**

Bacterial fermentation, respiration

Plant, animal, fungus, protozoa respiration

Evaporation from the oceans, chemical reaction

Volcanos; biomass, biofuel, fossil-fuel burning

# Sinks

Photosynthesis

Autotrophic bacterial respiration

Dissolution into oceans, lakes; transfer to ice caps, soil

Chemical weathering, photolysis in upper atmosphere

## Health effects

>15,000 ppmv affect respiration;

> 30,000 ppmv --> headaches, dizziness, nausea

# Carbon Storage Reservoirs

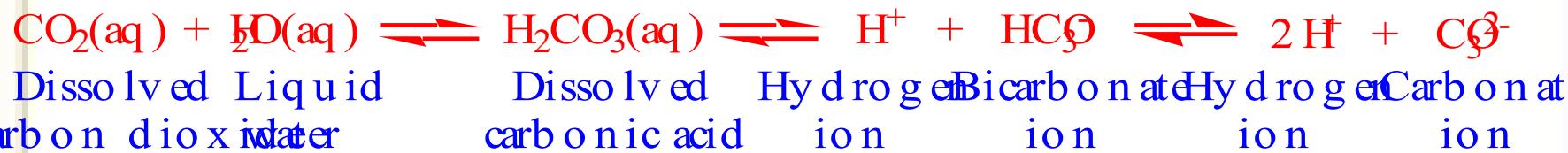
Location	GT-C
<b>Atmosphere</b>	
Gas and particle	859
<b>Surface oceans</b>	
Live organic carbon	5
Dead organic carbon	30
Bicarbonate ion	500
<b>Deep oceans</b>	
Dead organic carbon	3000
Bicarbonate ion	40,000
<b>Ocean sediments</b>	
Dead organic carbon	10,000,000
<b>Land/ocean sediments</b>	
Carbonate rock	60,000,000
<b>Land</b>	
Live organic carbon	800
Dead organic carbon	2000

# Carbon Dioxide Aqueous Chemistry

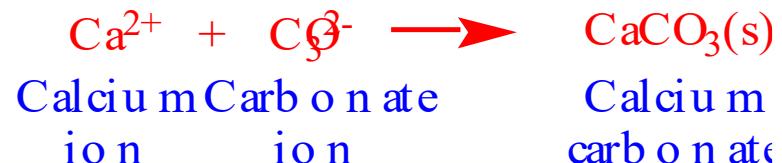
## Dissolution/Dissociation



Gaseous      Dissolve  
carbon        carbon  
dioxide        dioxide

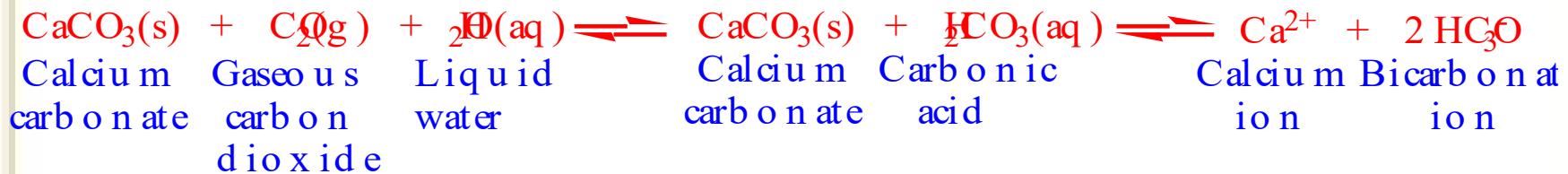
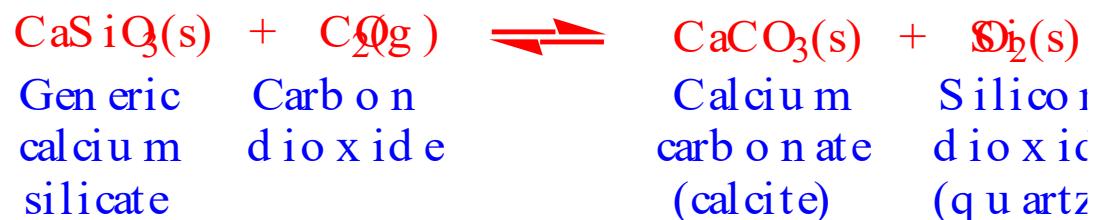


## Formation of calcium carbonate

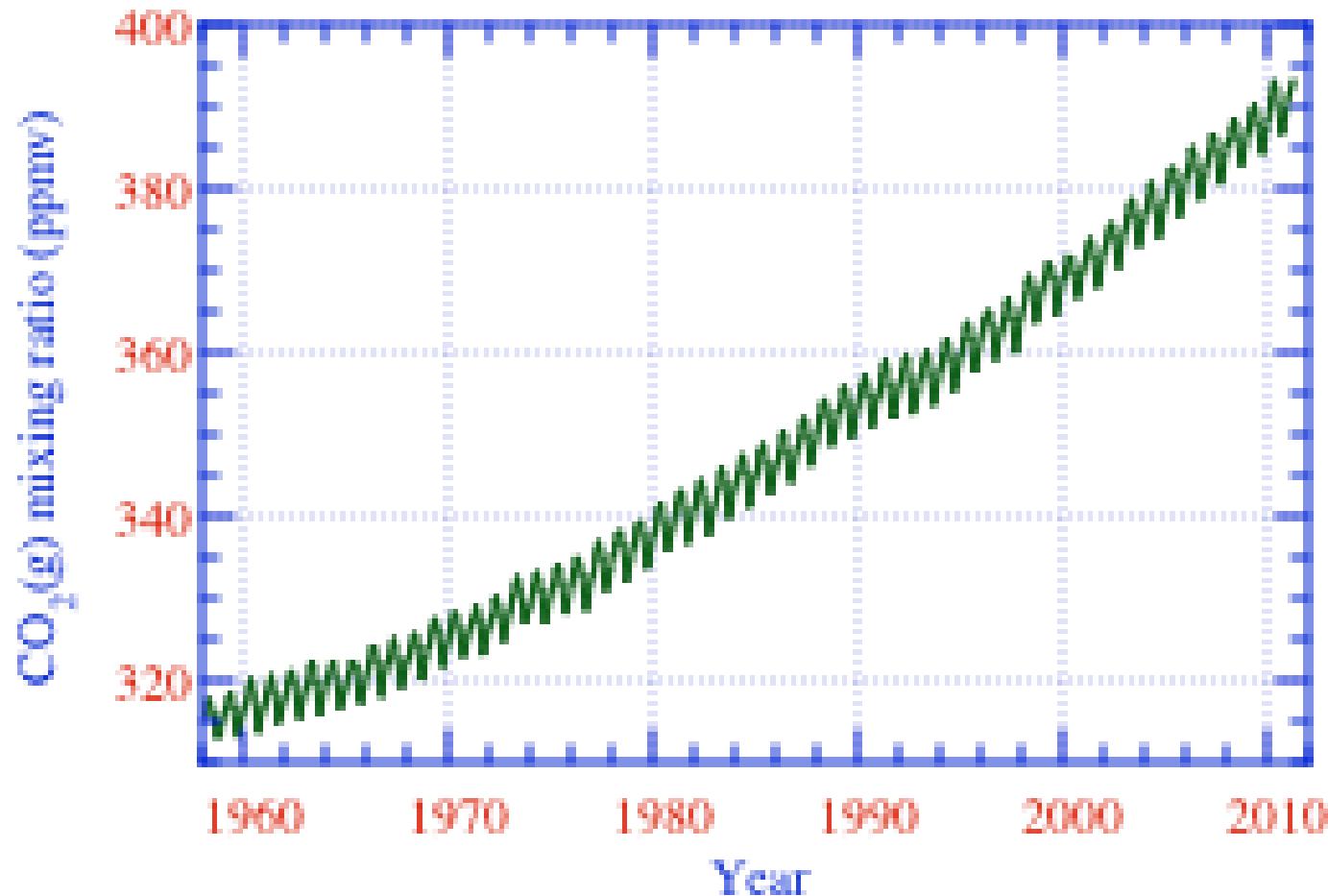


# Chemical Weathering

Breakdown and reformation of rocks and minerals at the atomic and molecular level by chemical reaction



## Carbon Dioxide Mixing Ratio



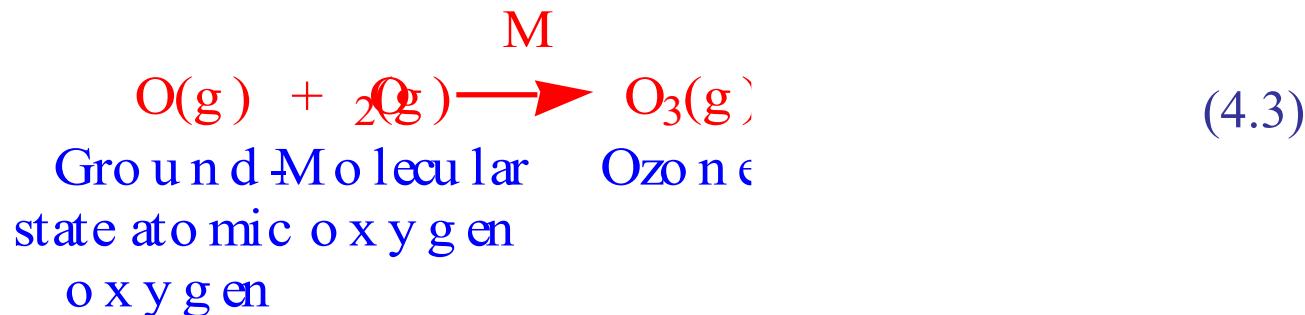
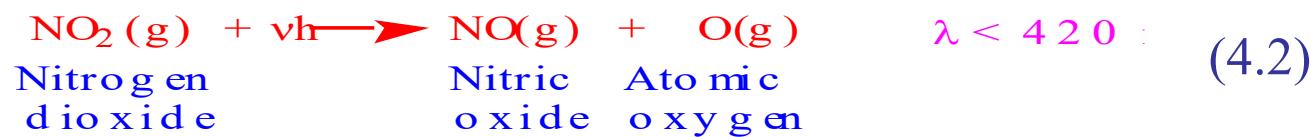
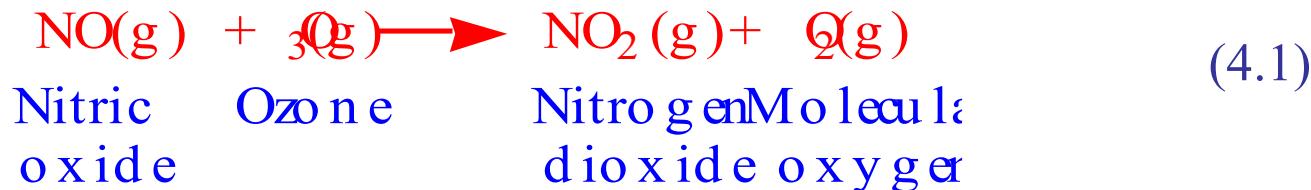


Los Angeles Smog

# Los Angeles (July 23, 2000)



## Photostationary State Ozone



# Photostationary State Ozone

$$\chi_{O_3} = (J/N_d k_1) (\chi_{NO_2(g)} / \chi_{NO(g)})$$

(4.4)

## Example 4.1.

Estimate ozone mixing ratio when

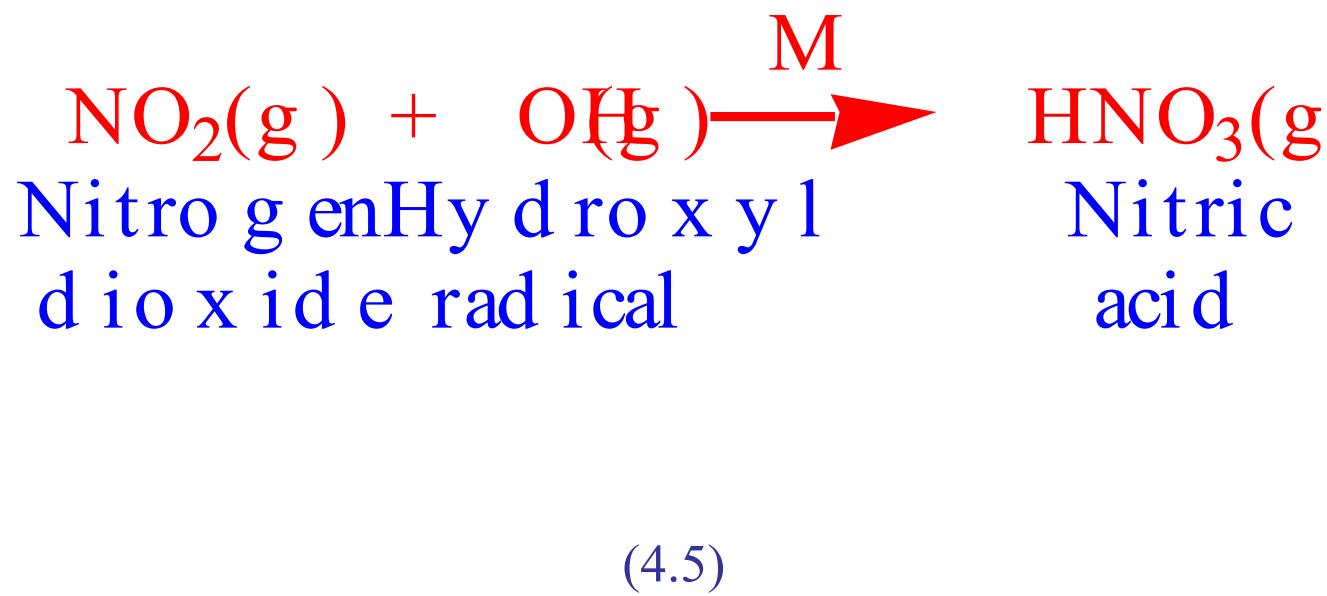
$p_d$	= 1013 mb	$T$	= 298 K
$\chi_{NO(g)}$	= 5 pptv	$\chi_{NO_2(g)}$	= 10 pptv
$k_1$	= $1.8 \times 10^{-14}$ cm <sup>3</sup> molec. <sup>-1</sup> s <sup>-1</sup>	$J$	= 0.01 s <sup>-1</sup>

$$\Rightarrow N_{O_3(g)} = 1.1 \times 10^{12} \text{ molec. cm}^{-3}$$

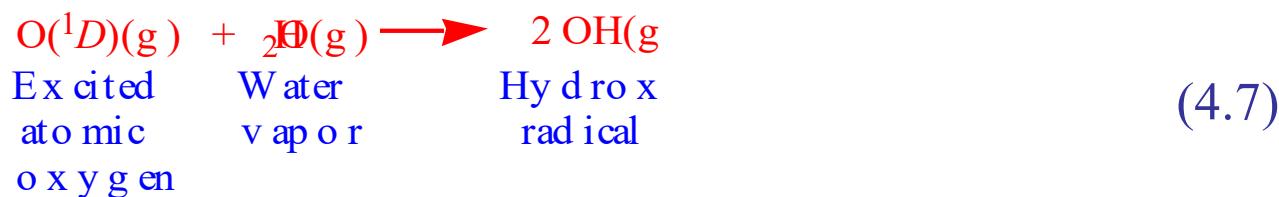
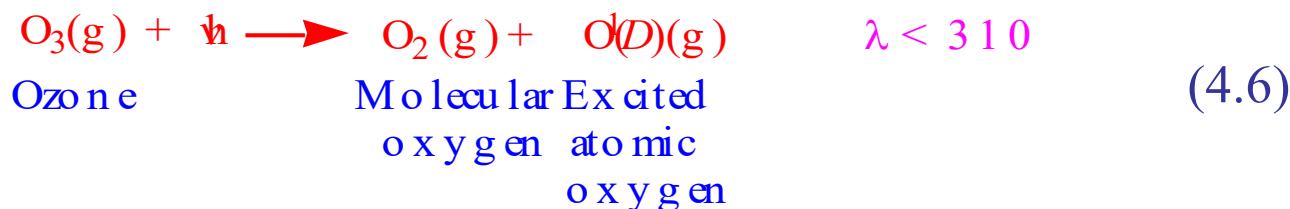
$$\Rightarrow N_d = 2.46 \times 10^{19} \text{ molec. cm}^{-3}$$

$$\Rightarrow \chi_{O_3(g)} = 44.7 \text{ ppbv}$$

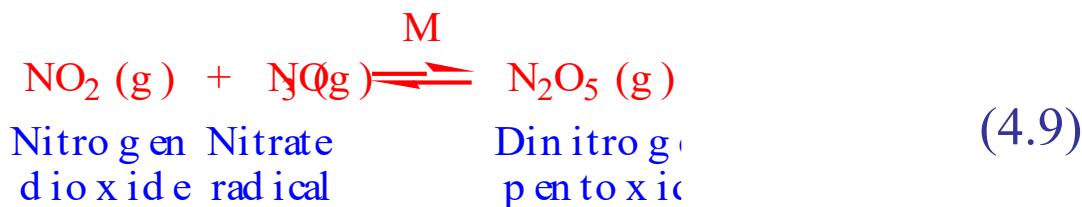
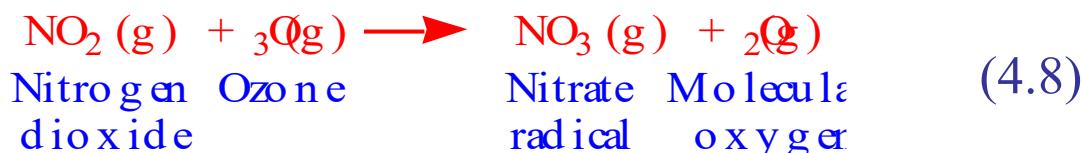
## Daytime Nitrogen Oxide Removal



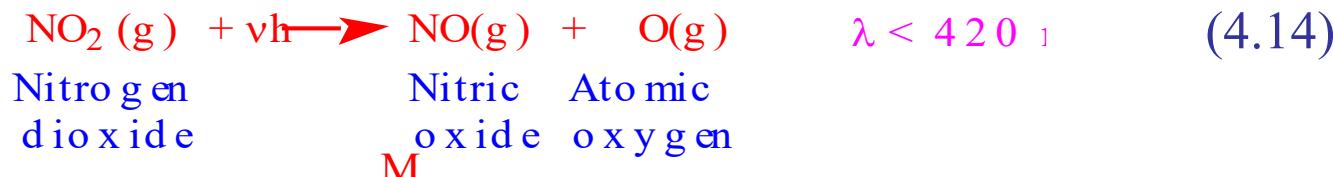
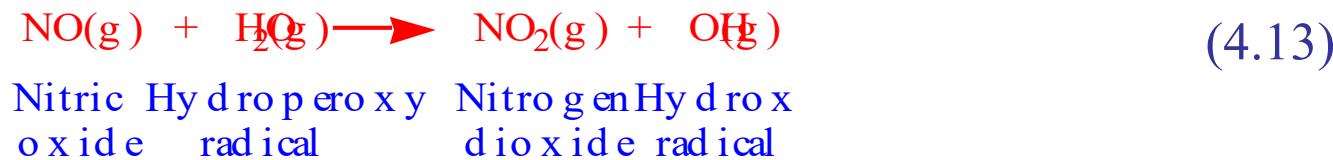
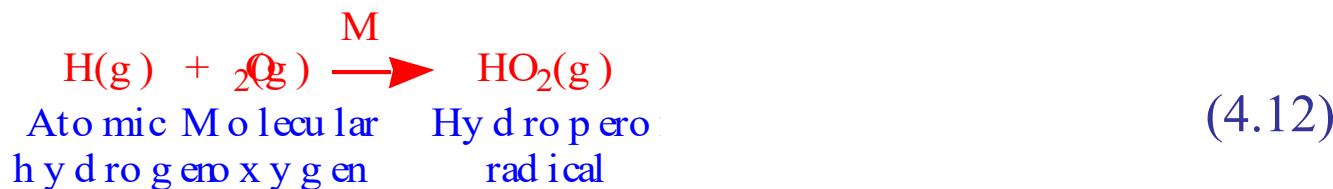
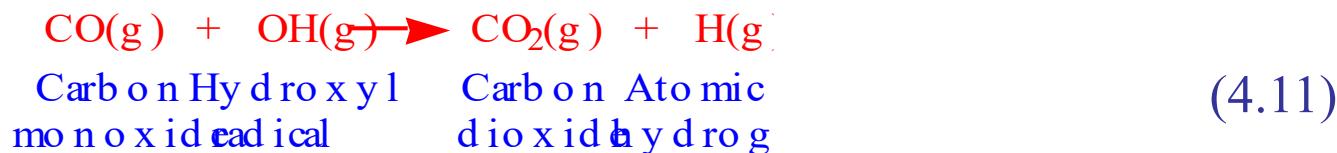
# Hydroxyl Radical Production



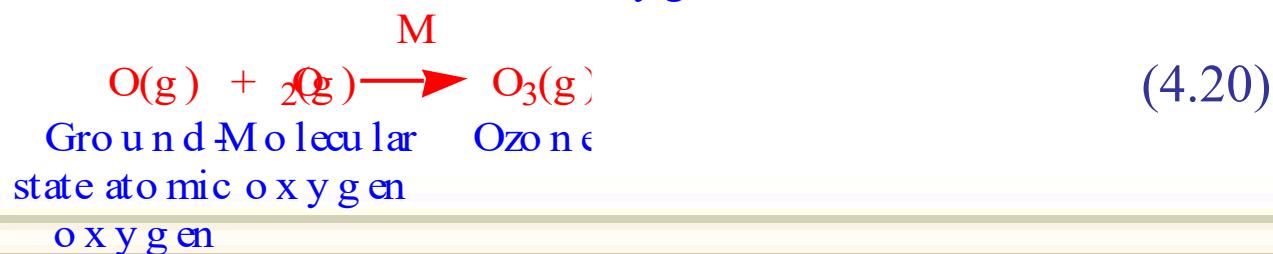
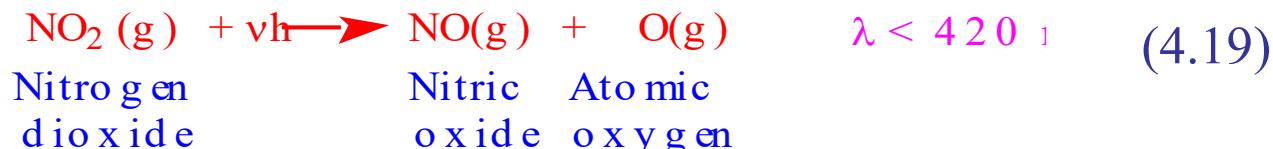
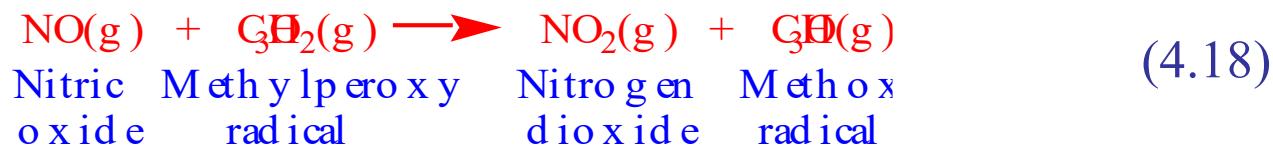
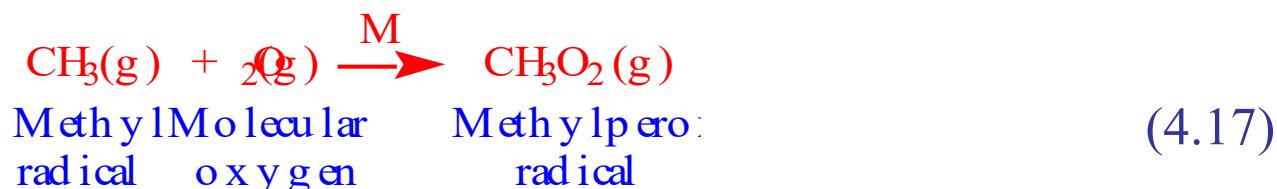
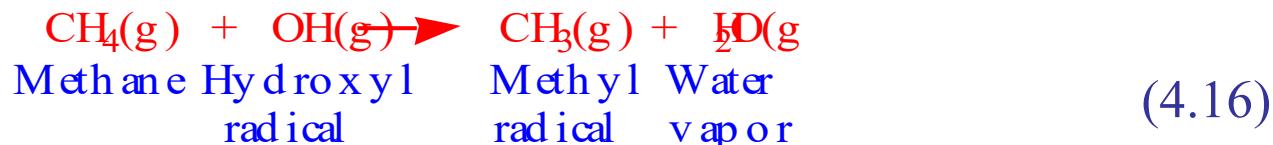
# Nighttime Nitrogen Chemistry



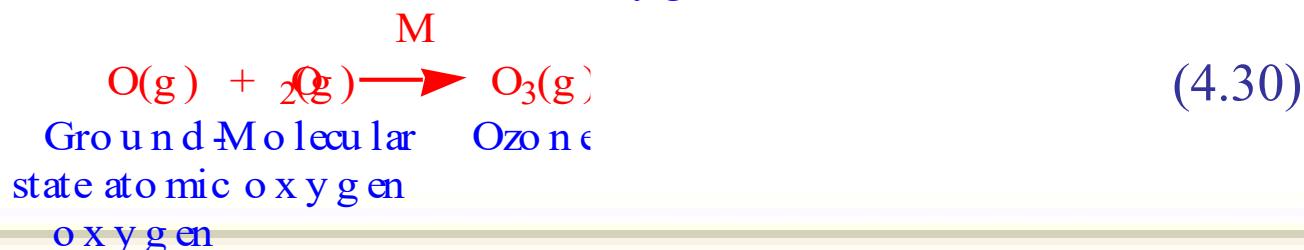
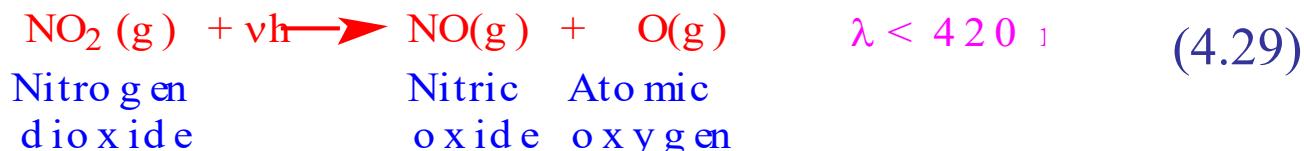
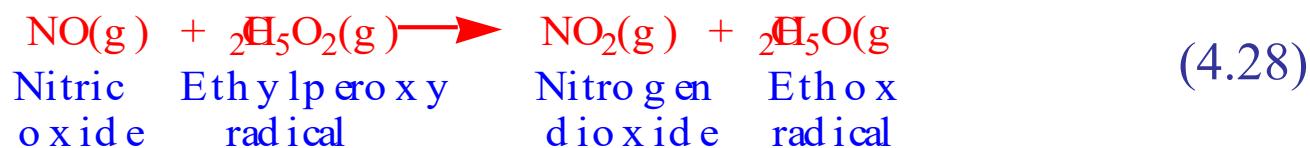
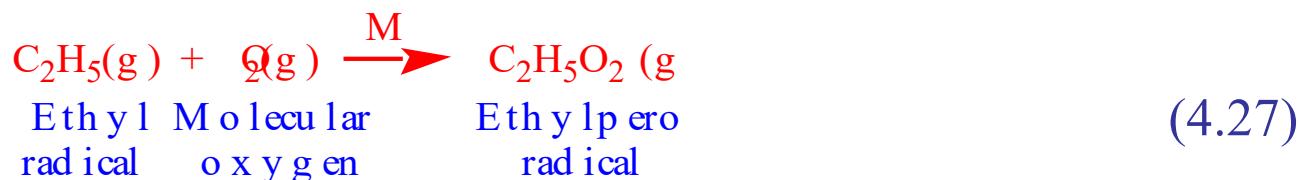
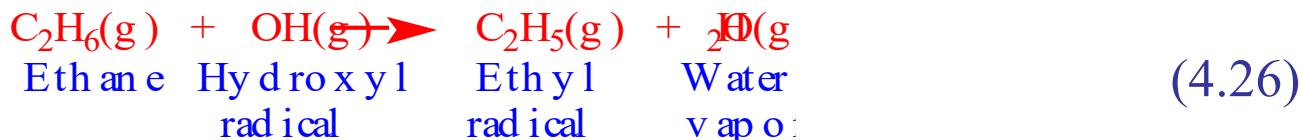
# Ozone Production From Carbon Monoxide



# Ozone Production From Methane



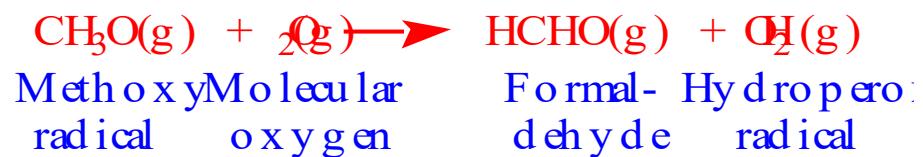
# Ozone Production From Ethane



# Production of Formaldehyde and Acetaldehyde

Formaldehyde

(4.21)

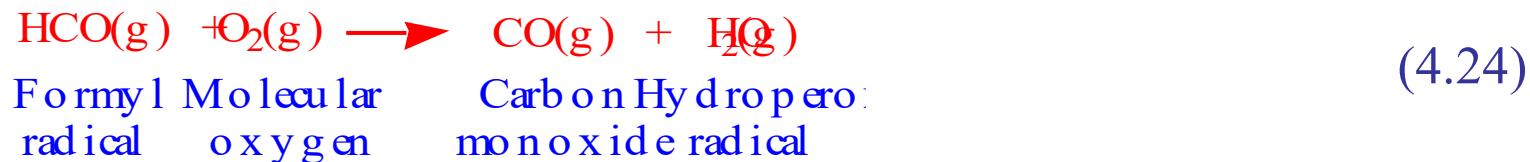
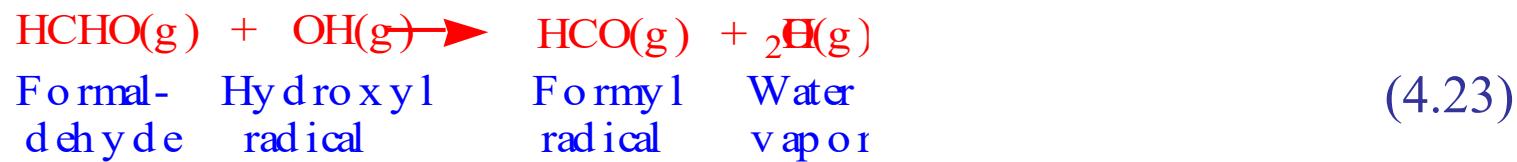
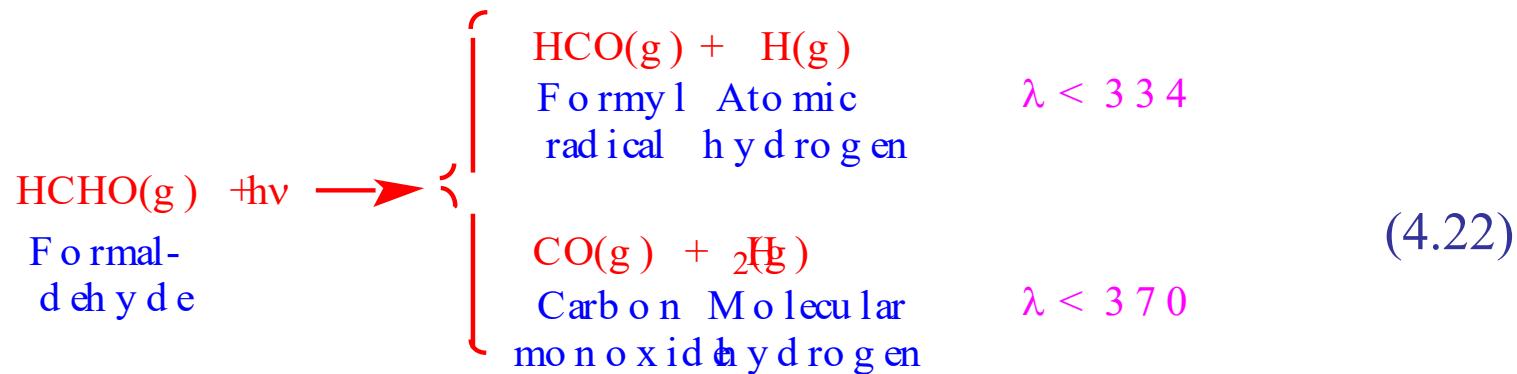


Acetaldehyde

(4.22)

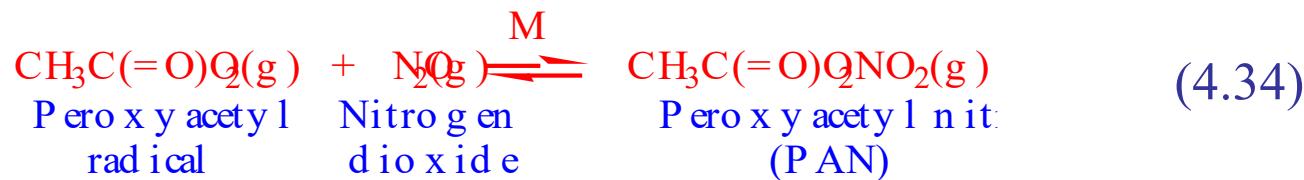
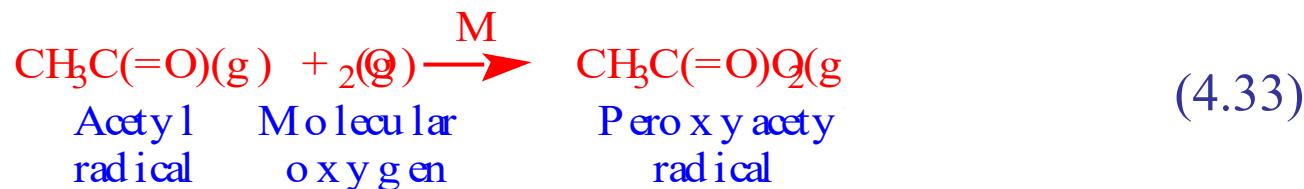
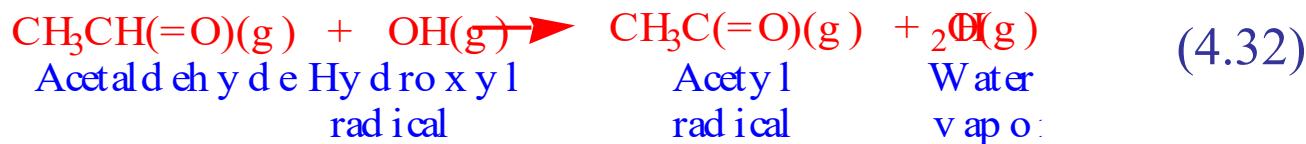


# Ozone From Formaldehyde

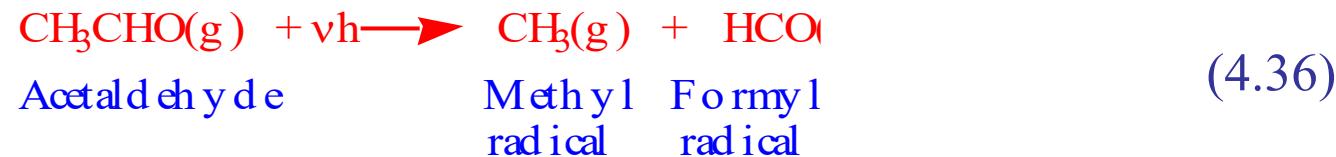


--> Form O<sub>3</sub> from both CO and HO<sub>2</sub>

# PAN Production From Acetaldehyde

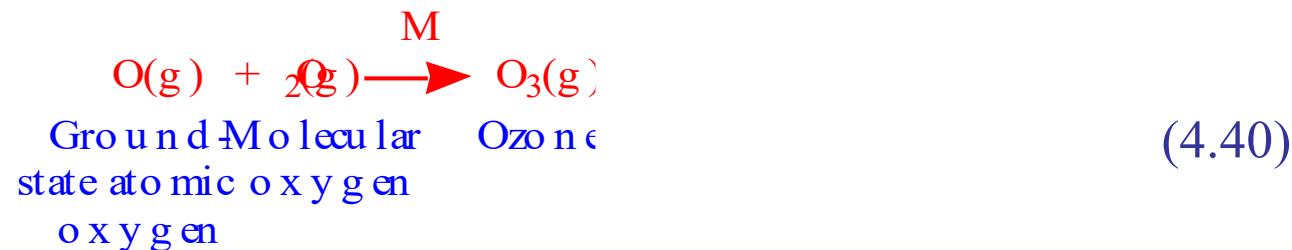
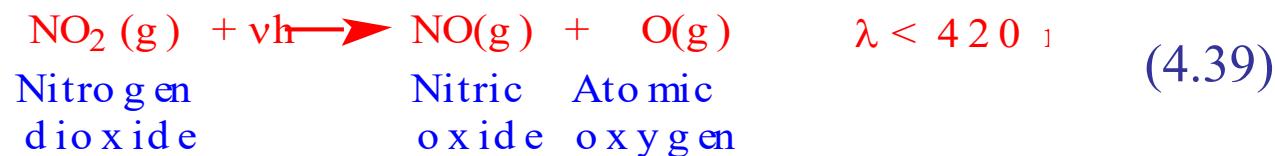
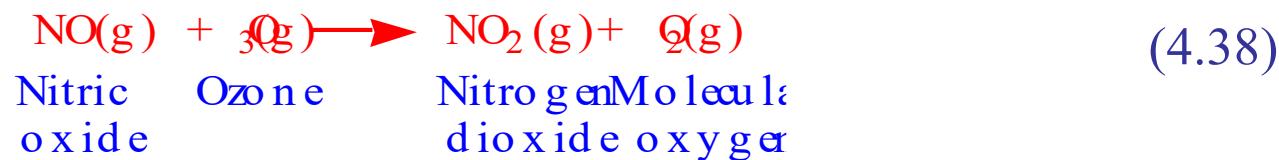
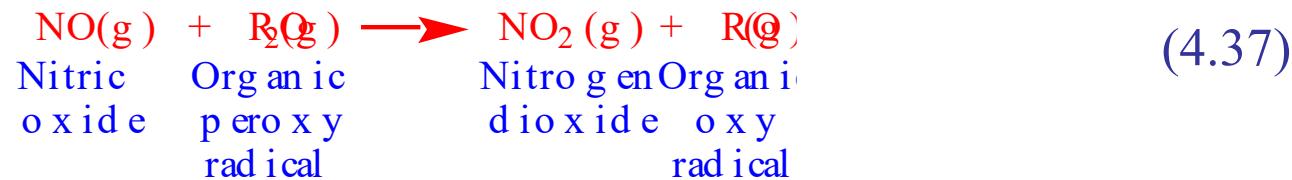


# Ozone Production From Acetaldehyde

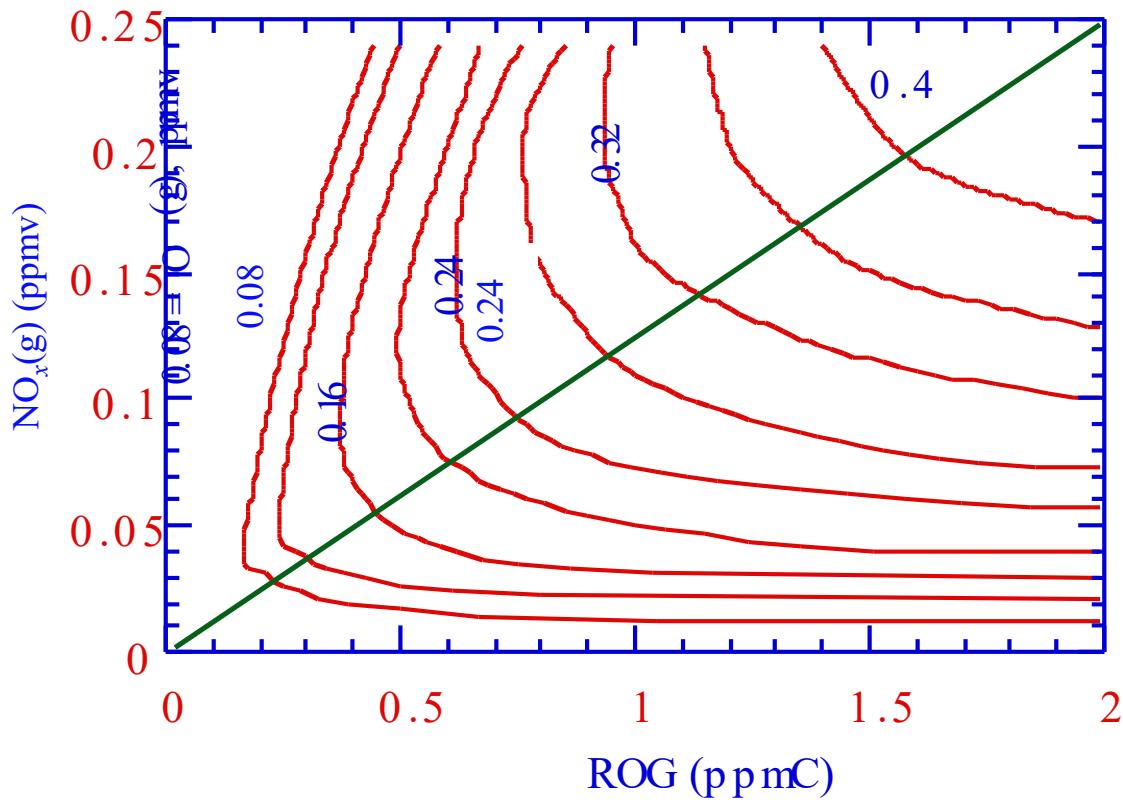


--> Form O<sub>3</sub> from NO<sub>2</sub>, CH<sub>3</sub>, and HCO

# Photochemical Smog Formation



# Ozone Isopleth



Contours are ozone (ppmv)

Figure 4.9

# Source/Receptor Regions in Los Angeles

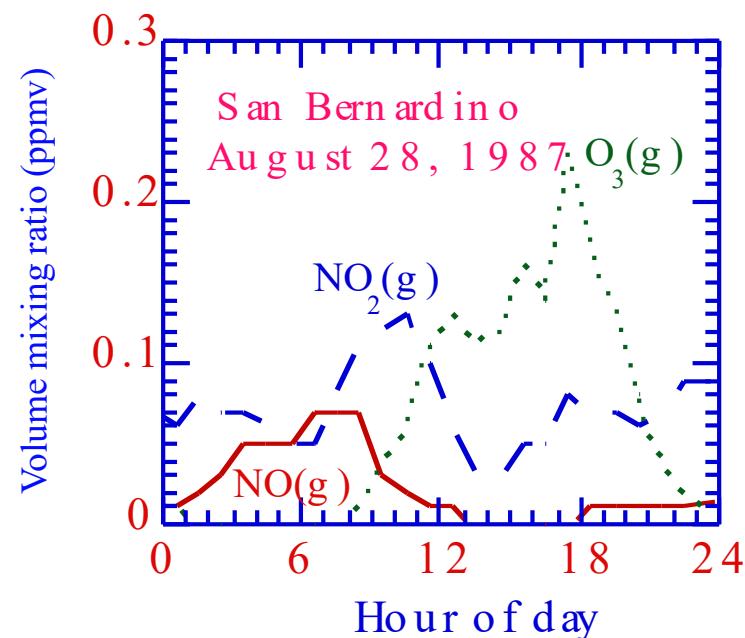
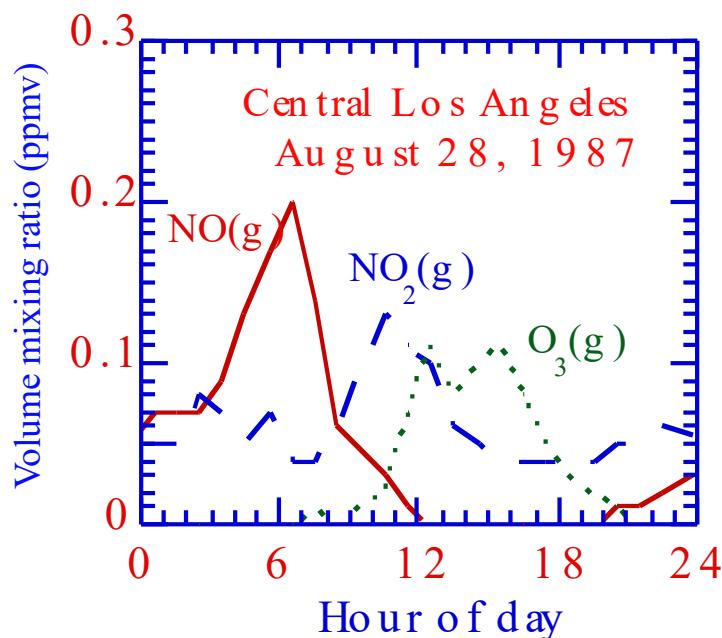


Figure 4.10

# Daily Emissions in Los Angeles (1987)

Substance	Emissions (tons day <sup>-1</sup> )	Percent of total
Carbon monoxide [CO(g)]	9796	69.3
Nitric oxide [NO(g)]	754	
Nitrogen dioxide [NO <sub>2</sub> (g)]	129	
Nitrous acid [HONO(g)]	6.5	
<b>Totals NO<sub>x</sub>(g) + HONO(g)</b>	<b>889.5</b>	<b>6.3</b>
Sulfur dioxide [SO <sub>2</sub> (g)]	109	
Sulfur trioxide [SO <sub>3</sub> (g)]	4.5	
<b>Totals SO<sub>x</sub>(g)</b>	<b>113.5</b>	<b>0.8</b>
Alkanes	1399	
Alkenes	313	
Aldehydes	108	
Ketones	29	
Alcohols	33	
Aromatics	500	
Hemiterpenes	47	
<b>Totals IROGs</b>	<b>2429</b>	<b>17.2</b>
Methane [CH <sub>4</sub> (g)]	904	6.4
<b>Totals Emissions</b>	<b>14132</b>	<b>100</b>

Table 4.1, Allen and Wagner, 1992

# Percent Emission by Source

Source Category	CO(g)	NO <sub>x</sub> (g)	SO <sub>x</sub> (g)	ROG
Stationary	2	24	38	50
Mobile	98	76	62	50
Total	100	100	100	100

Table 4.2, Chang et al., 1991

# Lifetimes of Organic Gases in Urban Air

ROG Species	Photolysis	[OH(g)]				
		5 ×	×	×	×	×
<i>trans</i> -2-Butene	---	52 m	4 y	6.3 d	4 m	17 m
Acetylene	---	3.0 d	---	2.5 y	---	200 d
Formaldehyde	7 h	6.0 h	1.8 h	2.5 y	2.0 d	3200 y
Acetone	23 d	9.6 d	---	---	---	---
Ethanol	---	19 h	---	---	---	---
Toluene	---	9.0 h	---	6 y	33 d	200 d
Isoprene	---	34 m	---	4 d	5 m	4.6 h

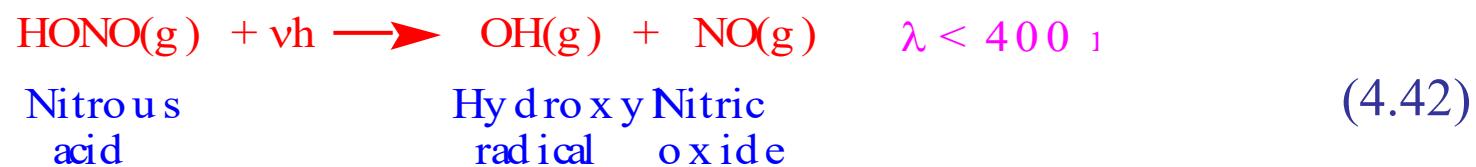
Table 4.3

# Most Abundant Species in Terms of Abundance and Reactivity

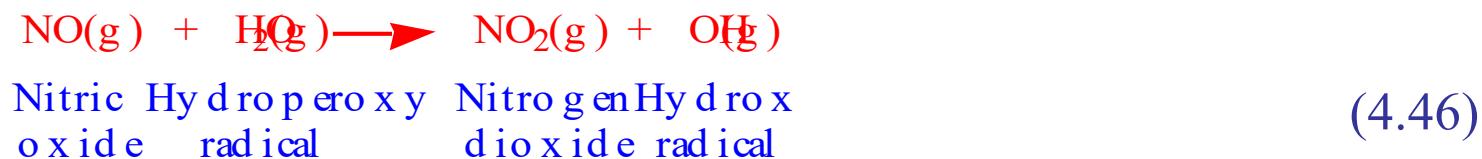
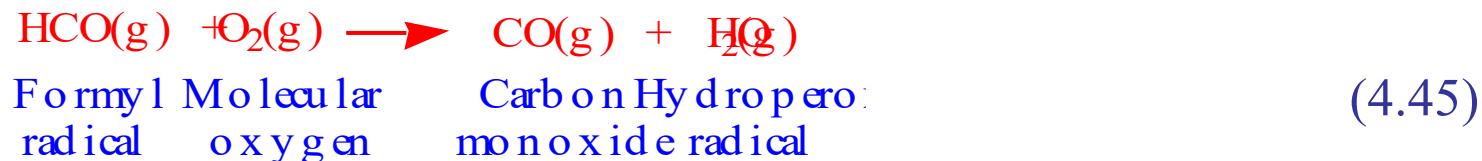
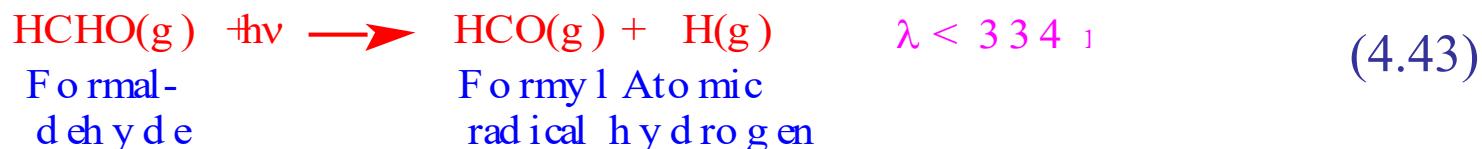
1. <i>m</i> - and <i>p</i> -Xylene	6. <i>i</i> -Pentane
2. Ethene	7. Propene
3. Acetaldehyde	8. <i>o</i> -Xylene
4. Toluene	9. Butane
5. Formaldehyde	10. Methylcyclopentane

Table 4.4, Lurmann et al., 1992

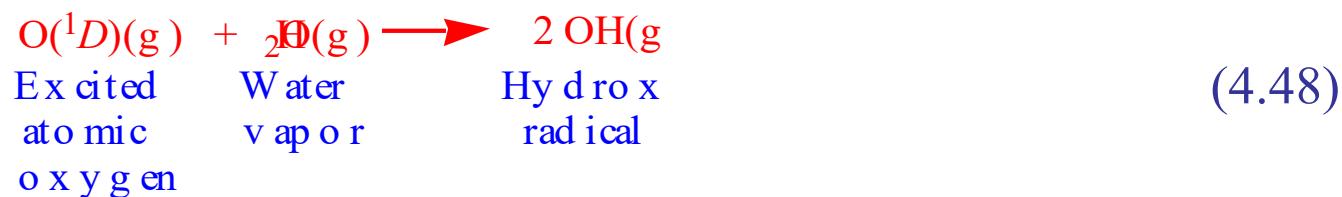
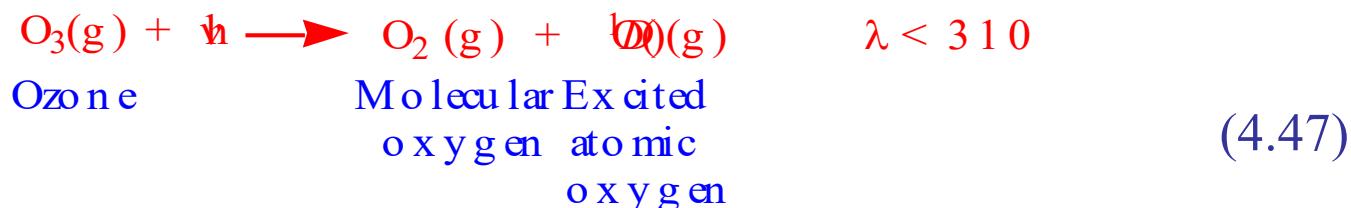
# Early Morning Source of OH in Polluted Air



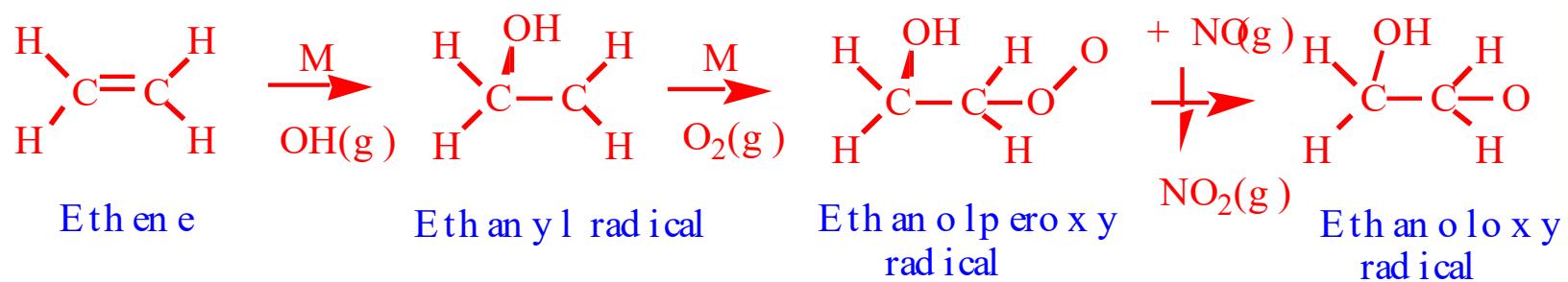
# Mid-Morning Source of OH in Polluted Air



# Afternoon Source of OH in Polluted Air

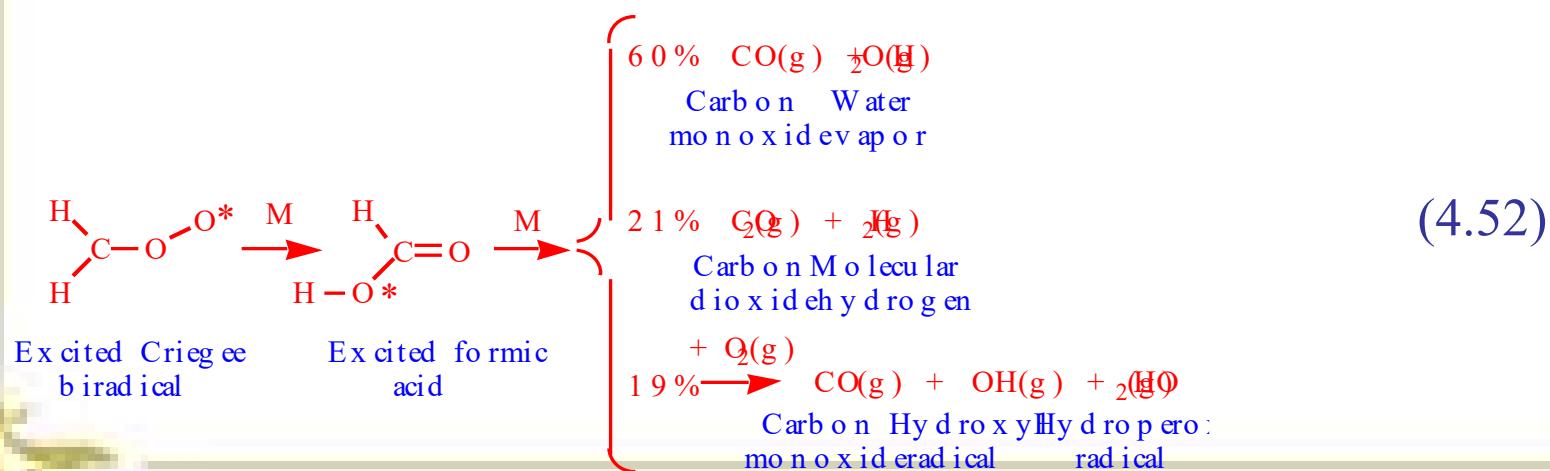
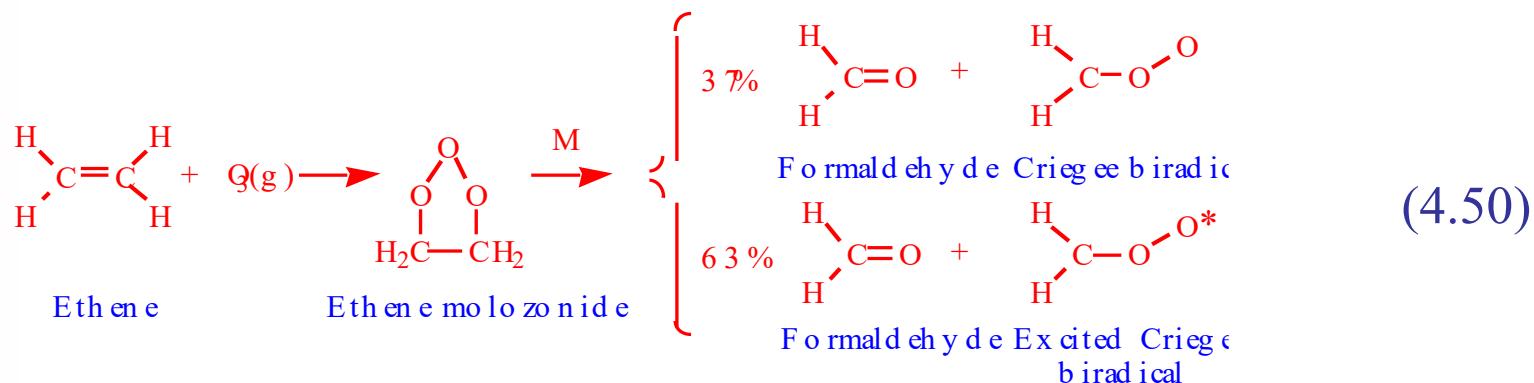


# Alkene Reaction with OH

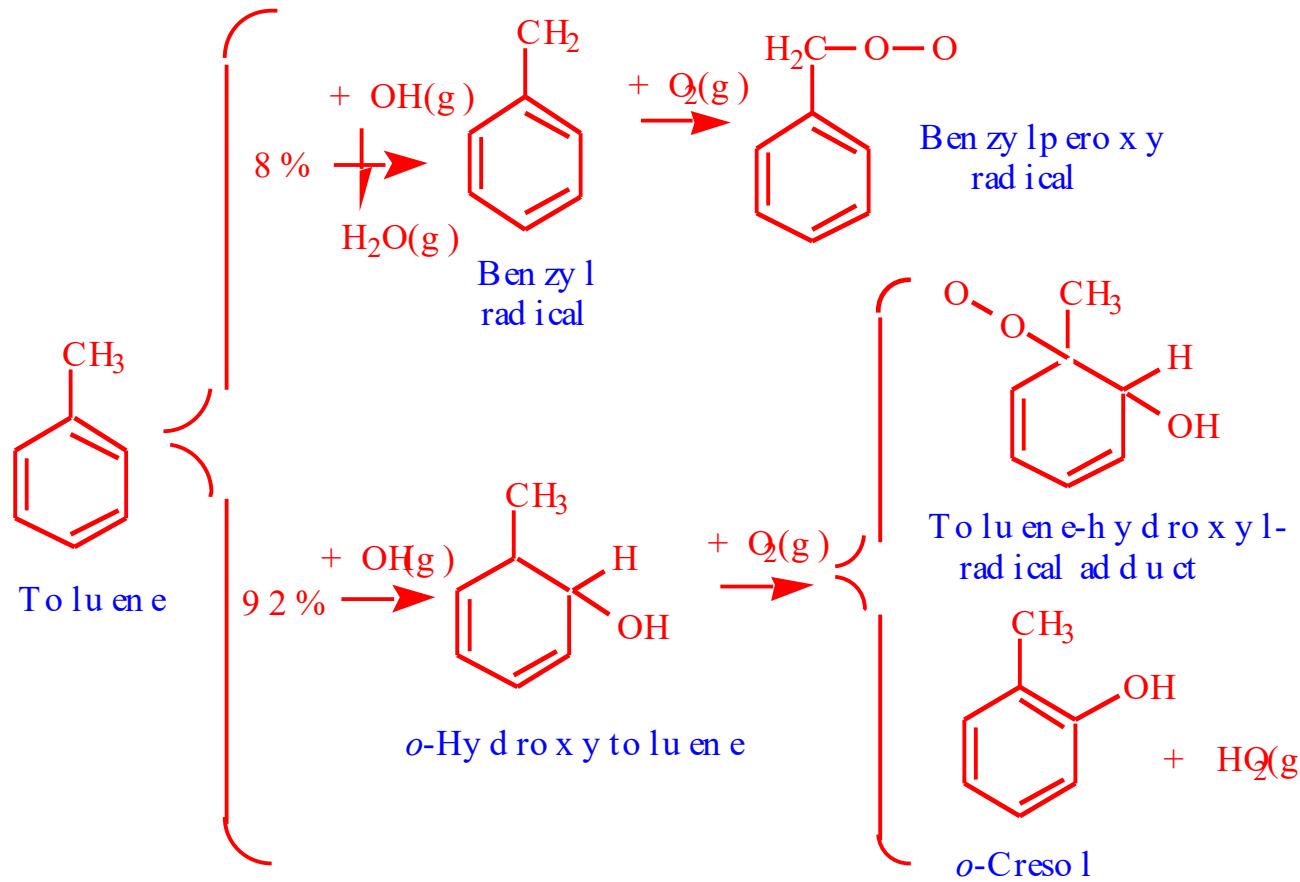


(4.49)

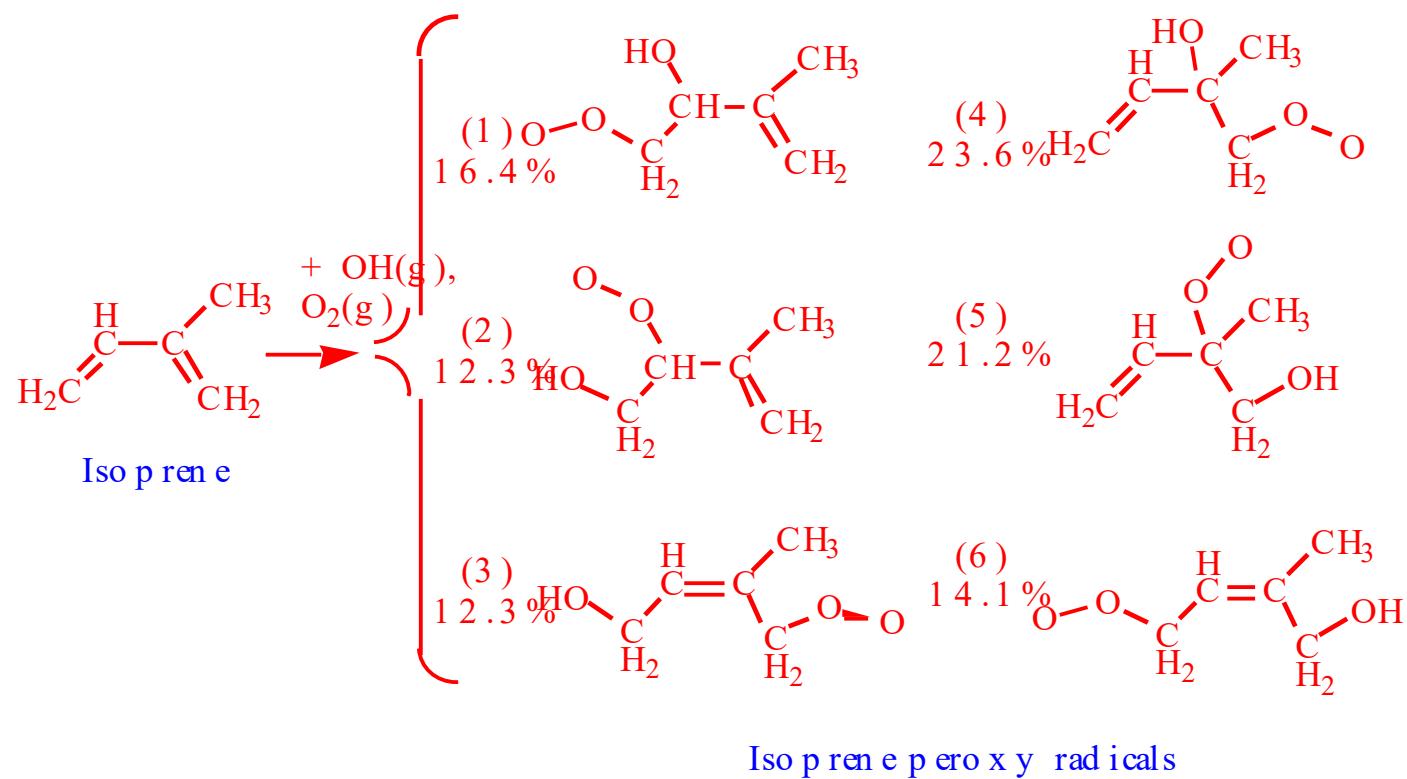
# Alkene Reaction with Ozone



# Toluene Reaction with OH

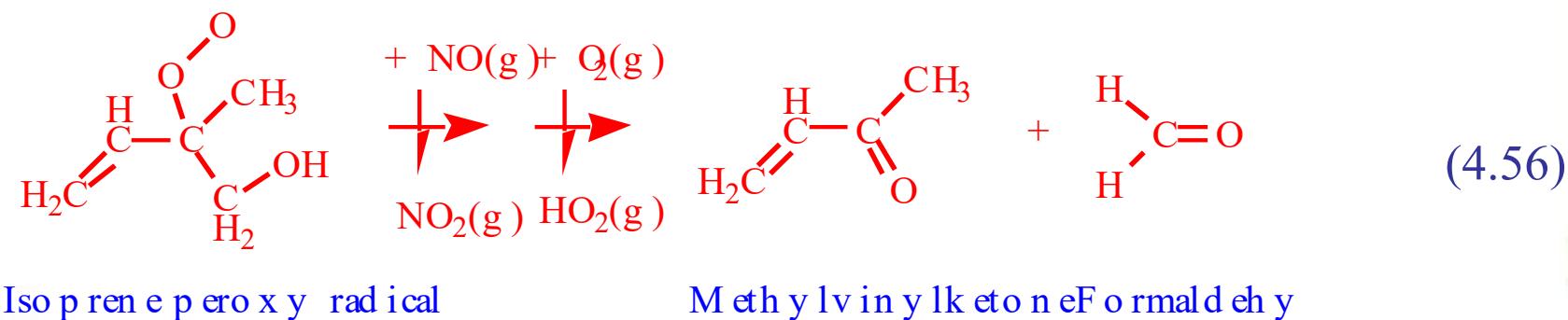
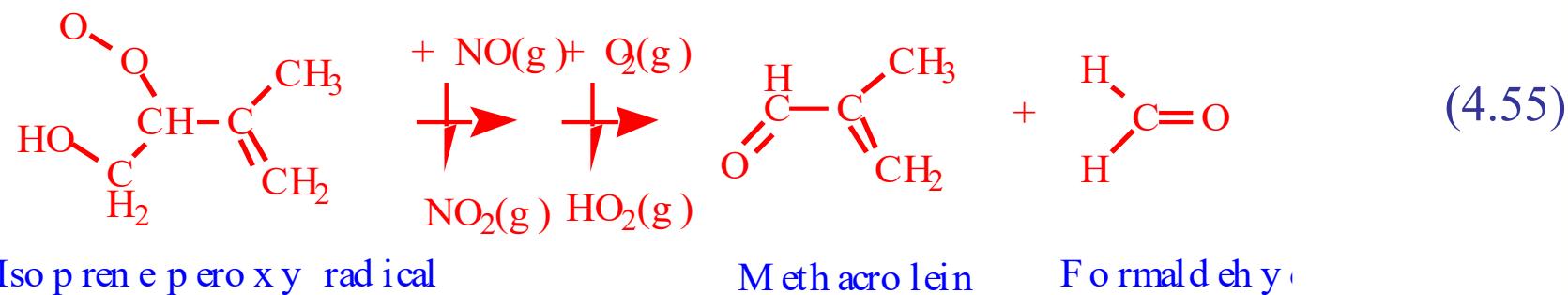


# Isoprene Reaction with OH

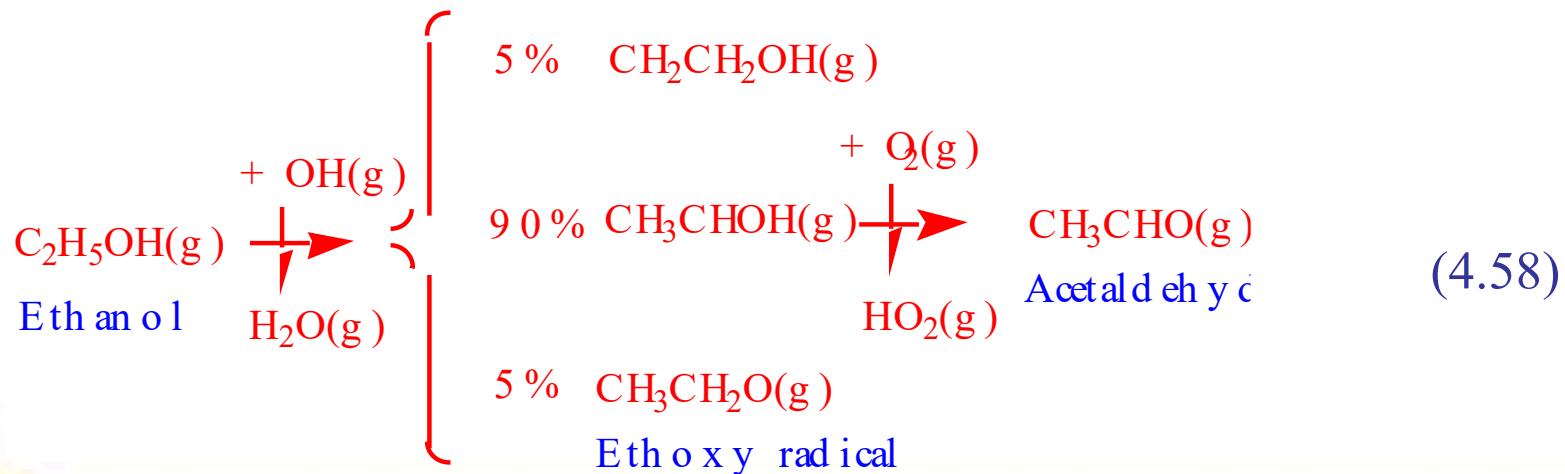
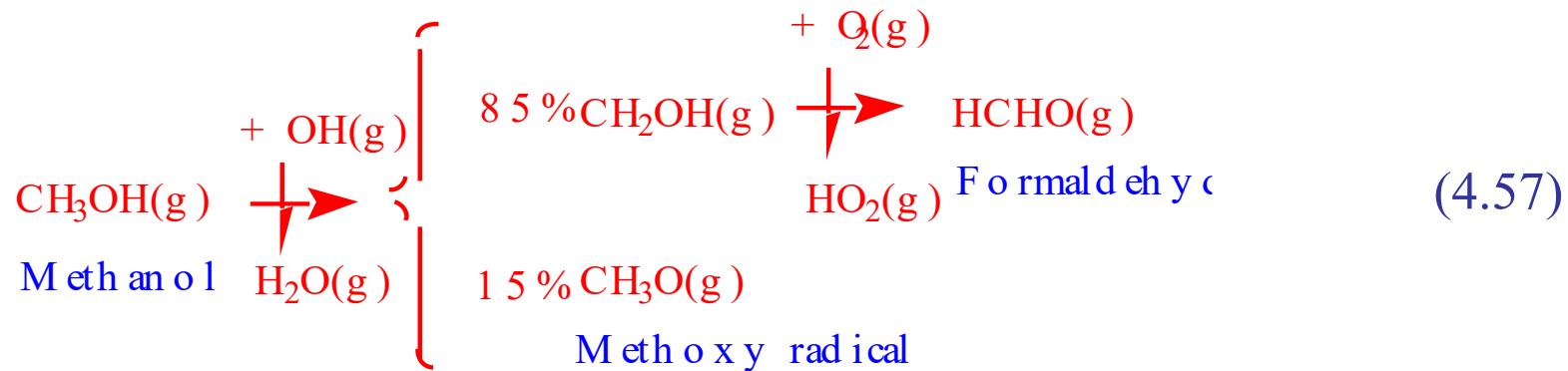


(4.54)

# Methacrolein and Methyvinylketone Production



# Alcohol Reactions with OH



# **Chapter 2:**

## **The Sun, the Earth, and the Evolution of the Earth's Atmosphere**

# Origin of the Sun

15 billion years ago (bya). Big bang. All mass in universe compressed to single point  $10^9 \text{ kg m}^{-3}$  density,  $T=10^{12} \text{ K}$

Aggregates of ejected material collapsed gravitationally to form earliest stars.

Temperatures in cores increased due to compressional heating

When temperatures reached 10 million K, nuclear fusion of H into He and other elements began, releasing energy to power the stars.

As early stars aged, they ultimately exploded, ejecting elements to the universe around

# Cosmic Abundance of Hydrogen Relative to Other Elements

Element	Atomic Mass	Abundance of H Relative to Element
Hydrogen	1.01	1:1
Helium	4.00	14:1
Oxygen	16.0	1400:1
Carbon	12.0	2300:1
Nitrogen	14.0	11,000:1
Magnesium	24.3	24,000:1
Silicon	28.1	26,000:1
Iron	55.8	29,000:1
Aluminum	27.0	306,000:1
Sodium	23.0	433,000:1

# Origin of the Sun

4.6 bya interstellar material aggregated to form cloudy mass, the solar nebula

Sun formed from gravitational collapse of solar nebula

Today's Sun (90% H, 9.9% He)

Core (8-15 million K)

Intermediate interior (5-8 million K)

Hydrogen convection zone (HCZ) (6400 K - 5 million K)

10 million years for photon to travel from core to top of HCZ

Photosphere (4000-6400 K, effective 5785 K)

Chromosphere (4000 K - 1 million K. H energized and decays)

Corona (1 - 2 million K. Consists of ionized gases)

Solar wind --> Aurora Borealis and Aurora Australis on Earth

300-1000 km/s, 200,000 K at Earth

# Structure of the Sun

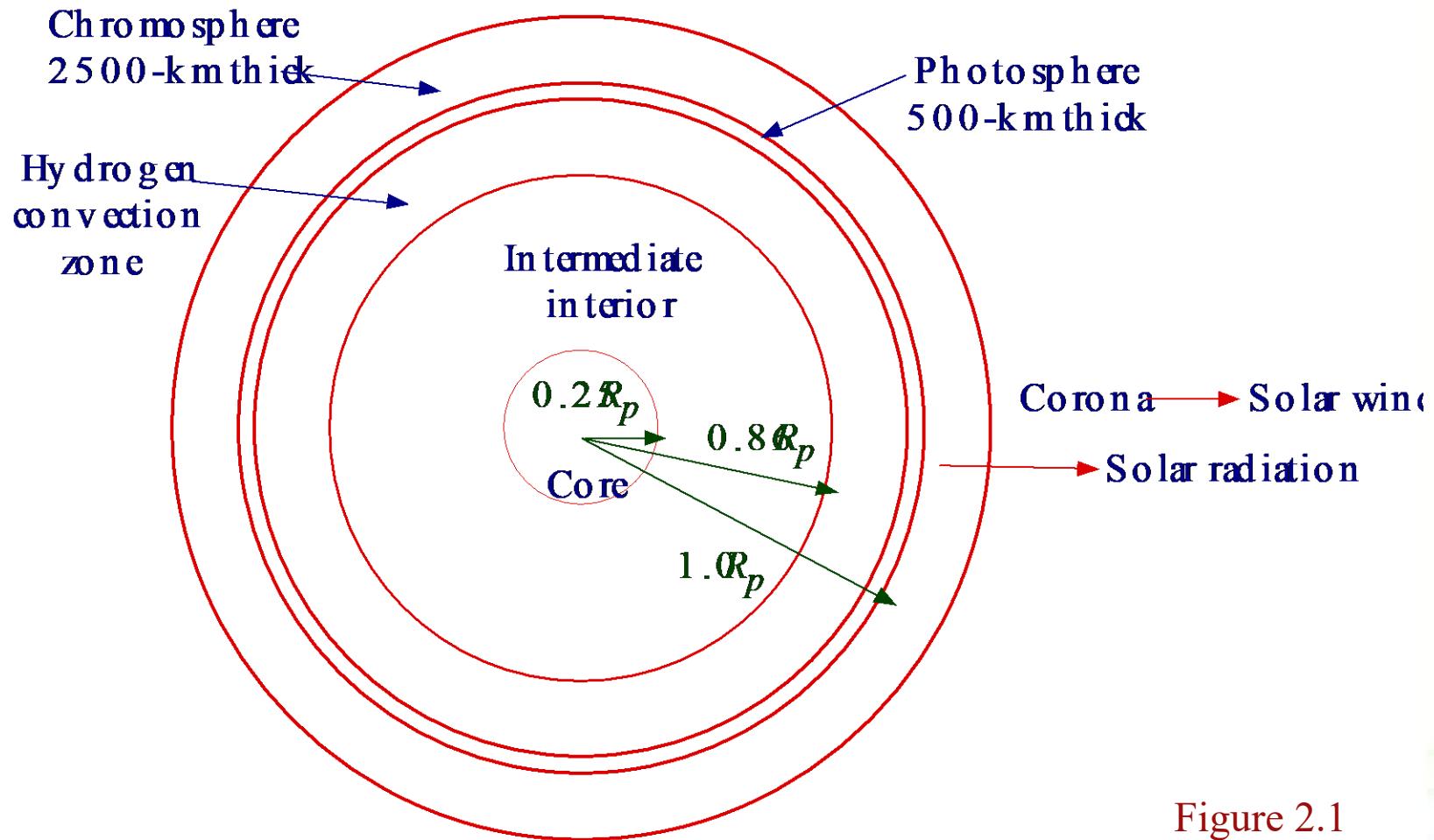


Figure 2.1

# Aurora Australis



David Miller, National Geophysical Data Center, available  
from NOAA Central Library

# Radiation Spectra

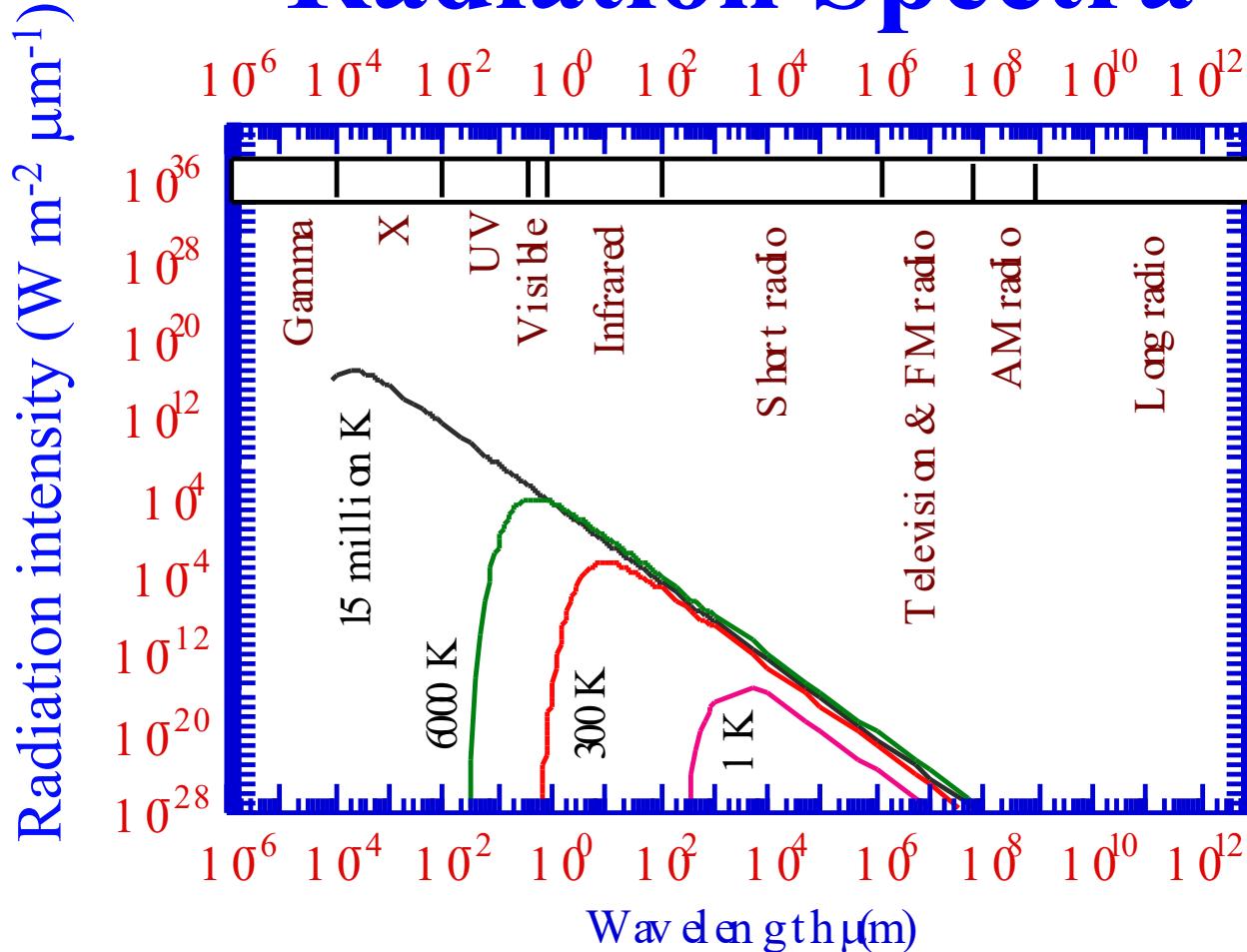


Figure 2.3

Wien's Law (2.1)

$$\lambda_p(\mu\text{m}) = 2897/T(\text{K})$$

Stefan-Boltzmann Law (2.2)

$$F_b (\text{W/m}^2) = \epsilon \sigma T(\text{K})^4$$

# Emission Spectra of the Sun and Earth

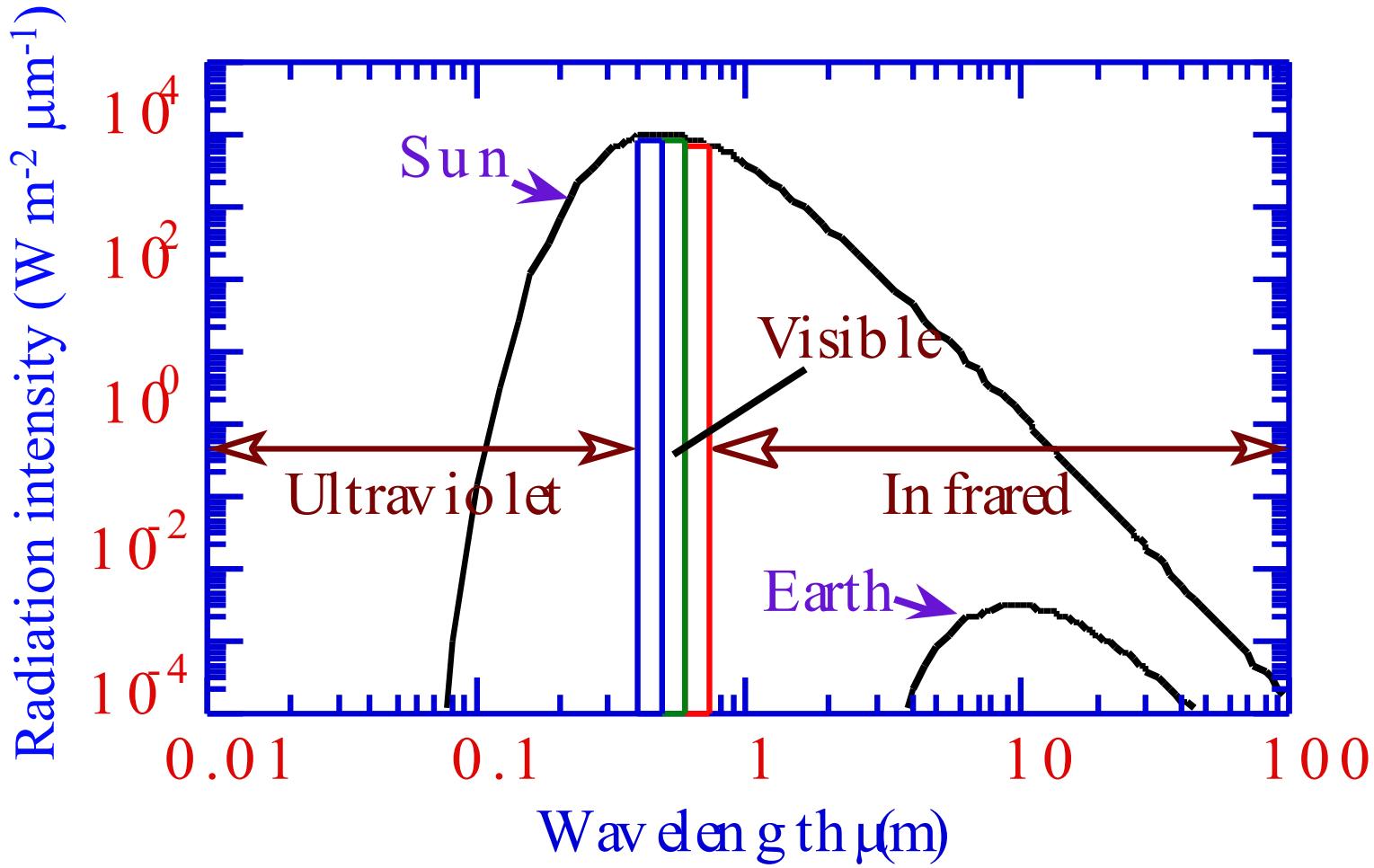


Figure 2.4 57

# Ultraviolet and Visible Spectra of the Sun

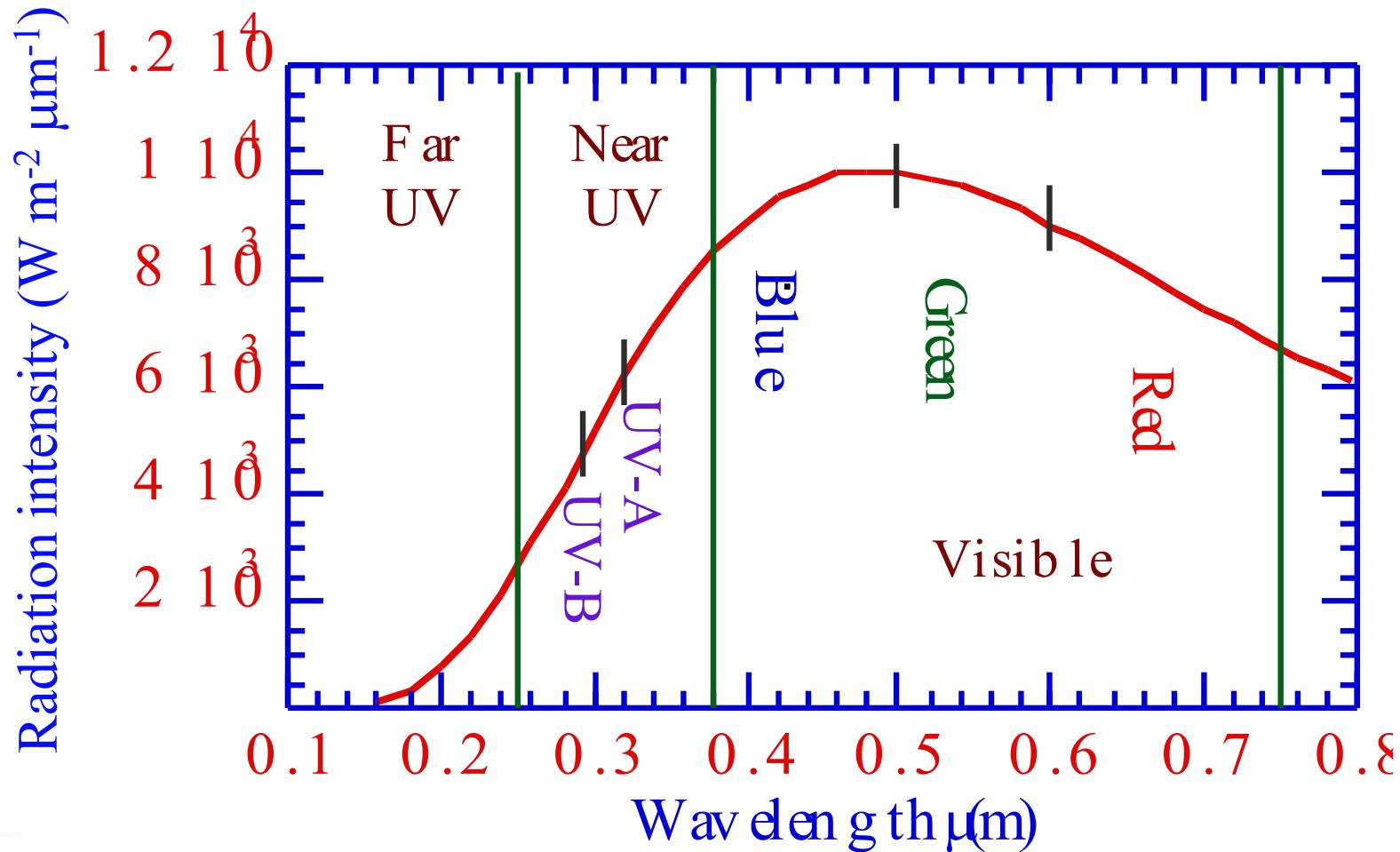


Figure 2.5  
58

# Origin of the Earth

4.6 bya, rock-forming elements, which were gases at high temperature in solar nebula, condensed into small solid grains as nebula cooled.

Grains accreted to planetesimals, such as asteroids and comets.

Asteroid: rocky body 1-1000 km in size that orbits sun.

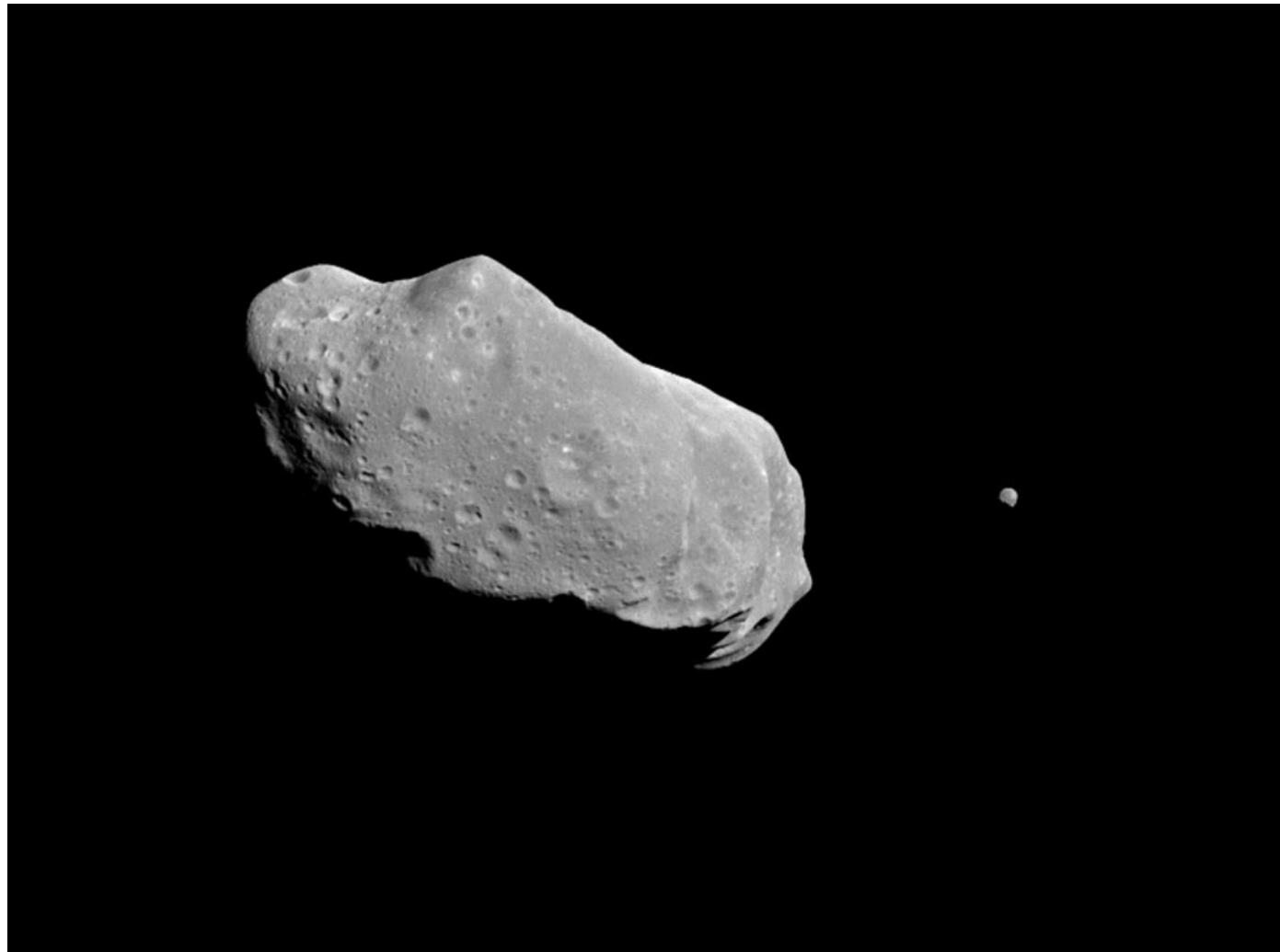
Comet: a small frozen mass that orbits sun

Planetesimals accreted to form the Earth

Meteorite bombardment over 500 million years aided Earth's growth.

Meteorite: Solid mineral or rock that reaches a planet's surface without vaporizing.

# Asteroid Ida and its Moon, Dactyl



# Meteorite Impacts

Meteorites contained rock-forming elements (Mg, Si, Fe, Al, Ca, Na, Ni) that condensed in solar nebula and noncondensable elements (H, He, O, C, Ne, N, S, Ar, P).

How did noncondensable elements enter meteorites?

They chemically reacted as gases to form high molecular-weight compounds that condensed.

Upon impact with Earth, some noncondensable elements (volatiles) evaporated (volatilized) on impact. Others have volatilized over time and have been outgassed through volcanos, fumaroles, steam wells, geysers.

# Composition of Stony Meteorites, Total Earth, and Earth's Crust

Element	Mass percent of element in			
	Stony Meteorites	Total Earth	Soil Crust	Ocean Crust
Oxygen	33.24	29.50	46.6	45.4
Iron	27.24	34.60	5.0	6.4
Silicon	17.10	15.20	27.2	22.8
Magnesium	14.29	12.70	2.1	4.1
Sulfur	1.93	1.93	0.026	0.026
Nickel	1.64	2.39	0.075	0.075
Calcium	1.27	1.13	3.6	8.8
Aluminum	1.22	1.09	8.1	8.7
Sodium	0.64	0.57	2.8	1.9

Table 2.2

# Formation of the Earth's Crust

4.5-4 bya, Earth's core hotter than today. Only mechanism of energy escape was conduction (transfer of energy molecule to molecule).

Because conduction is slow, internal energy could not dissipate, so entire Earth became molten and surface was magma ocean.

At that point, energy could be transferred to the surface by convection, the mass movement of molecules.

Convection allowed energy release and cooling at the surface, forming the Earth's crust 4.3-3.8 bya.

Dense elements (Fe, Ni) settled to core. Light ones (Si, Al, Na, Ca) rose to surface. Certain Mg, Fe silicates settled to mantle.

Today, land crust granite (quartz, potassium feldspar). Ocean crust basalt (plagioclase feldspar, pyroxene). Outer core liquid Fe, Ni; inner core solid Fe, Ni

# The Earth's Interior

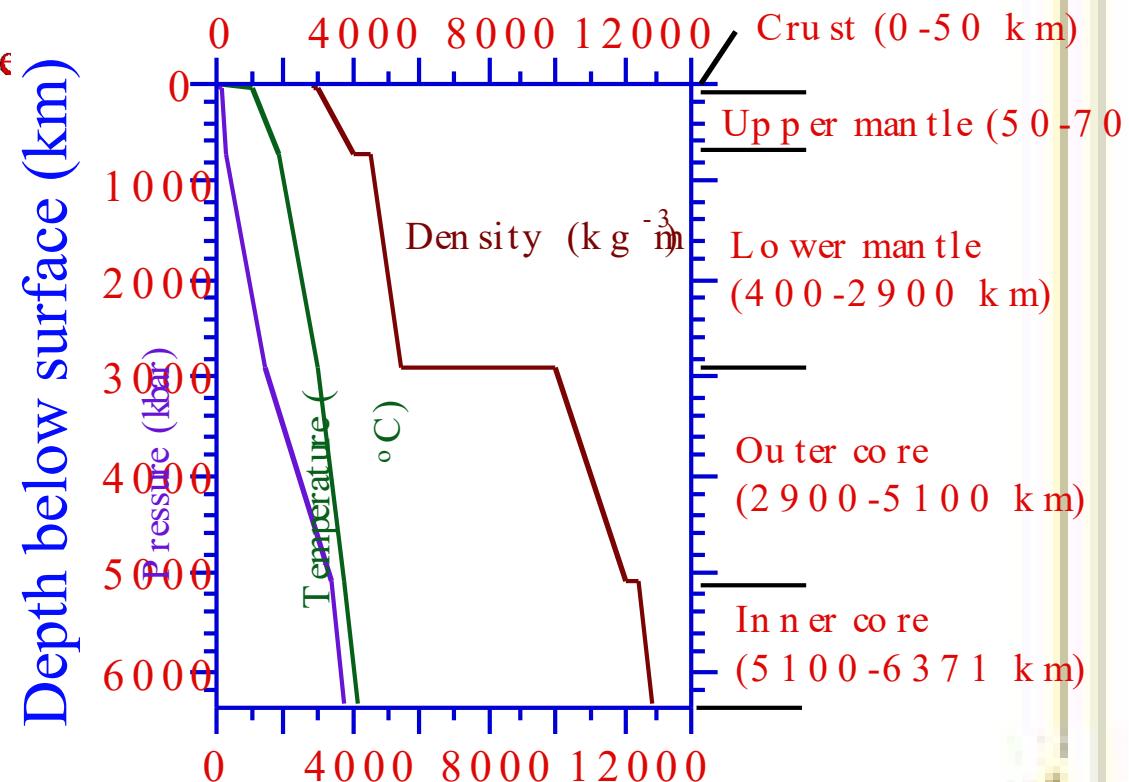
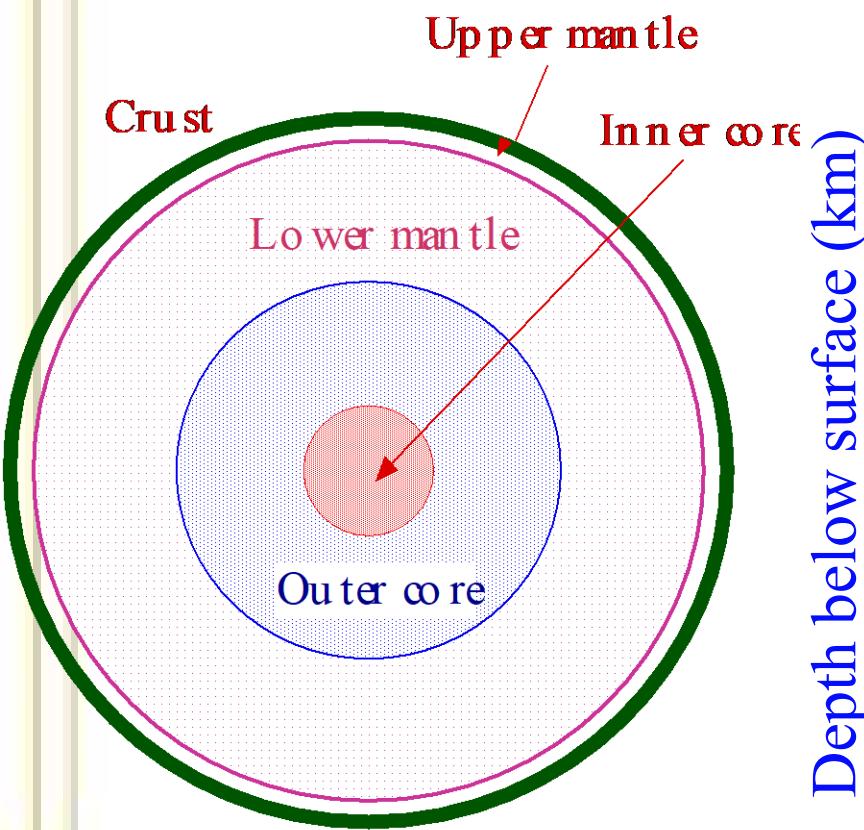


Figure 2.7

# Earth's First Atmosphere

Consisted mostly of H, He

During birth of the Sun, nuclear reactions are enhanced, increasing solar wind speeds and densities (T-Tauri stage of solar evolution).

Enhanced solar wind stripped off most H, He from the Earth.

Additional H, He lost by escape from Earth's gravitational field.

# Earth's Second Atmosphere

Initially due to outgassing by volcanos, fumaroles, steam wells, geysers.

Hydroxyl molecules (OH) bound in crustal minerals, became detached and converted reduced gases to oxidized gases:



Second atmosphere dominated initially by  $\text{CO}_2(\text{g})$ ,  $\text{H}_2(\text{g})$

Outgassed water vapor condensed to form the oceans.

# Timeline of Earth's Evolution

4.6 bya

3.5 bya

Formation of the Earth

Abiotic synthesis,

1953 Miller and Urey



--> first prokaryotes

single strand of DNA but no nucleus

conventional heterotrophs

# Classification of Organisms

## Energy Source

Sunlight	Phototroph
Oxidation of inorganic material	Lithotroph
Oxidation of organic material	Conventional heterotroph

## Carbon source

Carbon dioxide	Autotroph
Organic material	Heterotroph

\*Conventional heterotrophs obtain  
Energy and carbon from organic material

Table 2.3

# Classification of Organisms

## Photoautotrophs

Green plants, most algae, cyanobacteria, some purple and green bacteria

## Photoheterotrophs

Some algae, most purple and green bacteria, some cyanobacteria

## Lithotrophic autotrophs

Hydrogen bacteria, colorless sulfur bacteria, methanogenic bacteria, nitrifying bacteria, iron bacteria

## Lithotrophic heterotrophs

Some colorless sulfur bacteria

## Conventional heterotrophs

Animals, fungi, protozoa, most bacteria

Table 2.4

# Hot Sulfur Springs in Lassen National Park



Lithotrophic autotrophs oxidize  $\text{H}_2\text{S}(\text{aq})$  to  $\text{H}_2\text{SO}_4(\text{aq})$ , which dissolves minerals into a “mud pot.” Alfred Spormann, Stanford University<sup>70</sup>

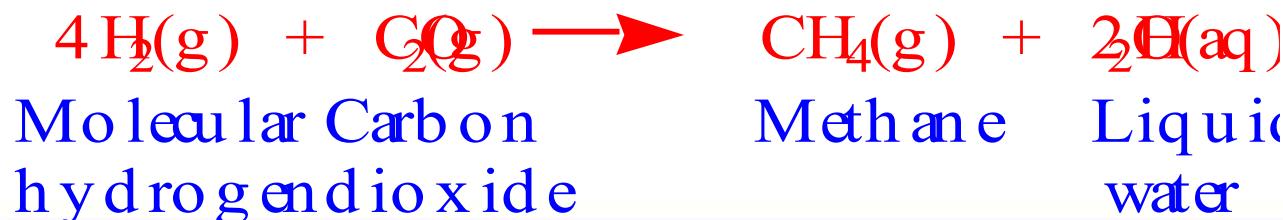
# $\text{CO}_2(\text{g})$ and $\text{CH}_4(\text{g})$ From Bacteria

Anaerobic respiration: production of energy from food where electron acceptor is not oxygen.

$\text{CO}_2(\text{g})$  by fermentation (2.3)



$\text{CH}_4(\text{g})$  by methanogenesis (lithotrophic autotrophs) (2.4)

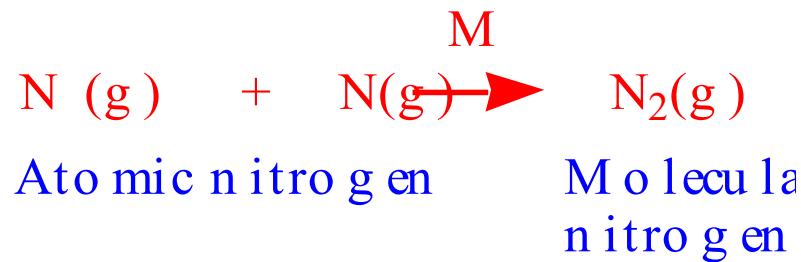
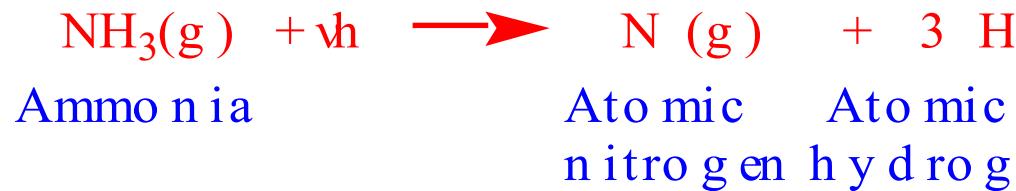


# Evolution of Molecular Nitrogen

4.6 bya

$\text{N}_2(\text{g})$  by ammonia photolysis

(2.5-2.6)



# Evolution of Molecular Nitrogen

3.2 bya

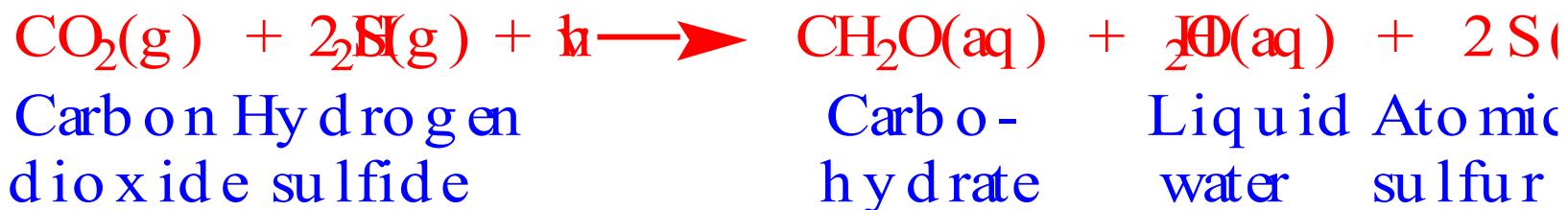
Denitrification: 2-step process

$\text{N}_2(\text{g})$  by anaerobic respiration (2.7-2.8)  
(conventional heterotrophs)

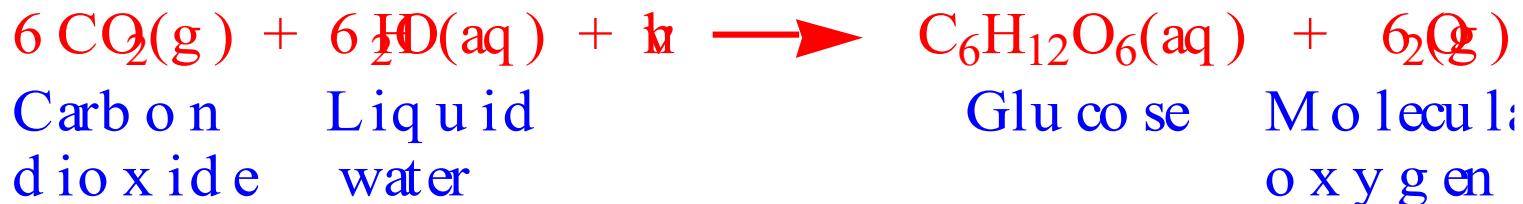


# Photosynthesis

Anoxygenic photosynthesis (photoautotrophs) (2.9)



- 2.3 bya                      Oxygenic photosynthesis (cyanobacteria) (2.10)  
1.4 bya                      Oxygen levels still 1% of today  
0.395-0.43 bya              Green plant photosynthesis



Photosynthesis in chlorophylls *a,b* : pigments that absorb visible  
Chlorophyll *a*. Absorbs red more efficiently  
Chlorophyll *b*. Absorbs blue more efficiently

# Hot Spring in Yellowstone National Park



Different colored photosynthetic cyanobacteria grow at in hot spring due to different temperatures. Alfred Spormann, Stanford University <sup>75</sup>

# Aerobic Respiration

O<sub>2</sub>(g) reacts with organic cell material to produce energy during cellular respiration, which is oxidation of organics in living cells.



Aerobic respiration developed first in prokaryotes (bacteria, blue-green algae), but spread with the advent of eukaryotes.

Eukaryote. Cell containing DNA surrounded by a true membrane-enclosed nucleus. Cells of higher organisms all eukaryotic.

Eukaryotic cells usually switch from fermentation to aerobic respiration when oxygen reaches 1% of present levels -->  
Eukaryotes developed about 1.4 bya, after oxygen rose to 1%

# Timeline of Earth's Evolution

4.6 bya	Formation of the Earth
3.5 bya	Abiotic synthesis, Denitrification
3.2 bya	Oxygen-producing photosynthesis by cyanobacteria
2.3 bya	Start of ozone formation
1.8 bya	Nitrification (aerobic)
1.5 bya	Nitrogen fixation (aerobic)
1.4 bya	Earliest eukaryotes
0.57 bya	First shelled invertebrates
0.43-0.5 bya	Primitive fish
0.395-0.43 bya	First land plants -- oxygen and ozone increase

# The Nitrogen Cycle

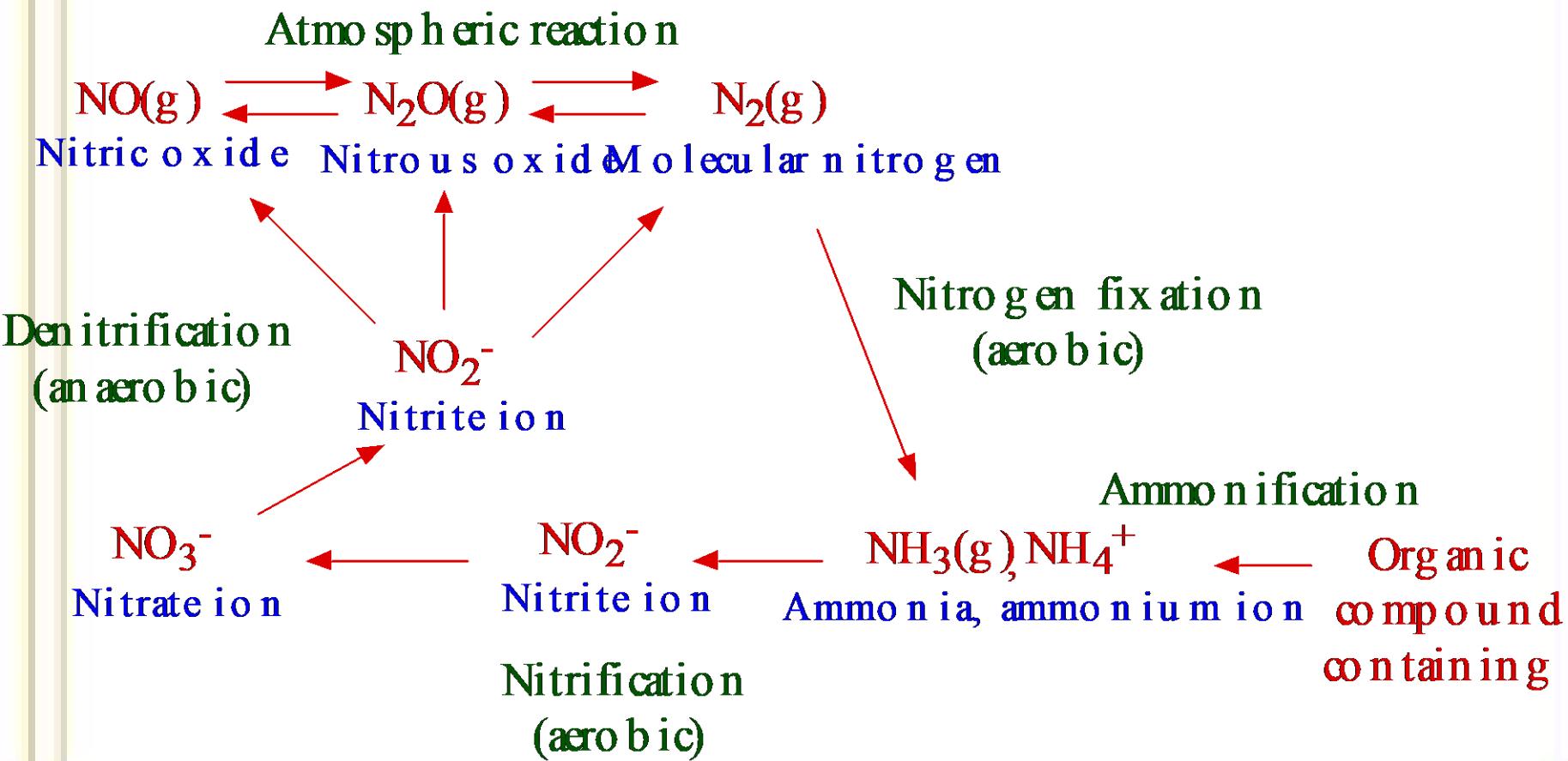


Figure 2.11

# Evolution of the Earth's Second Atmosphere

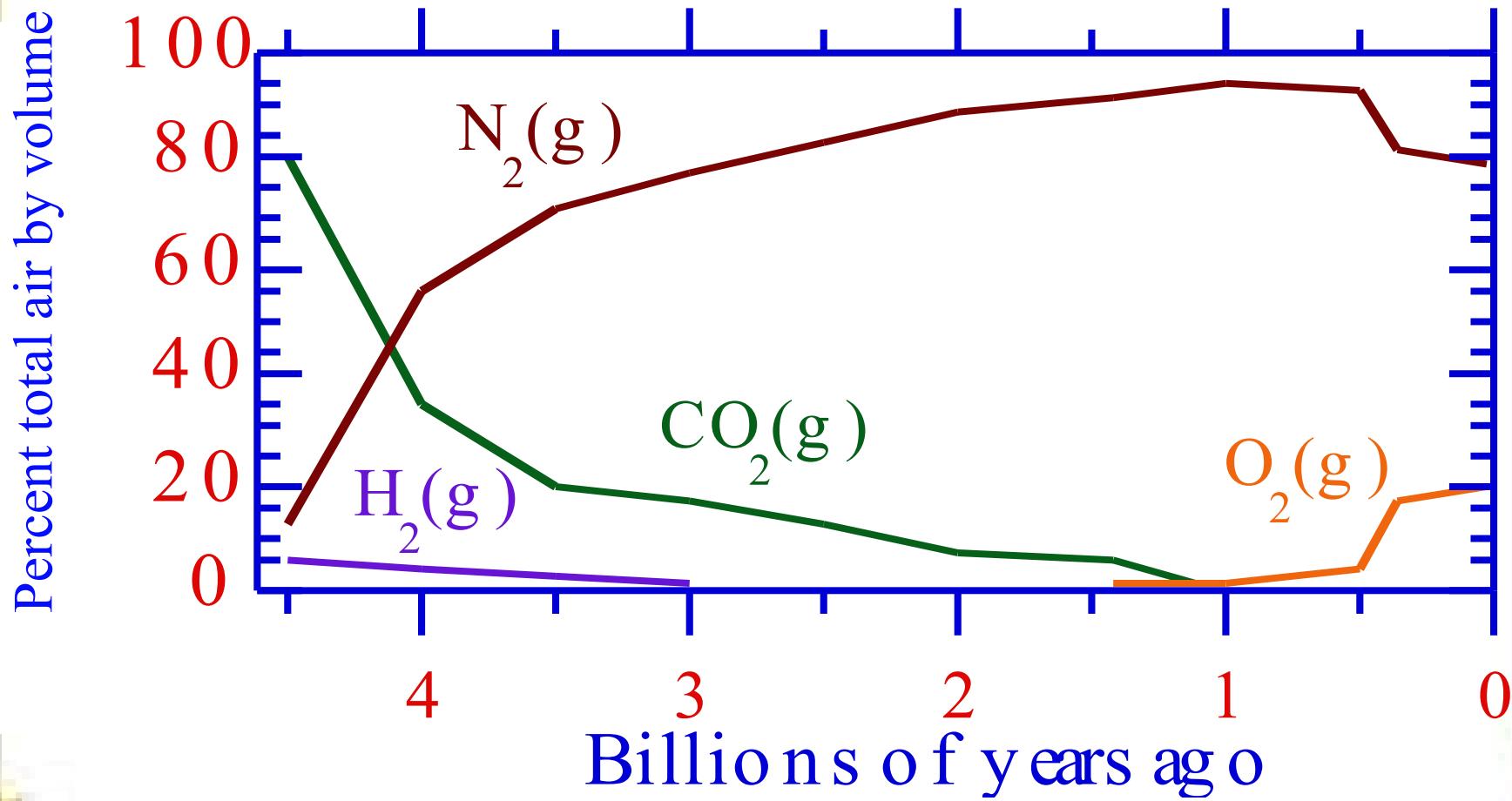


Figure 2.12