



Atmospheric Physics

Lecture 8

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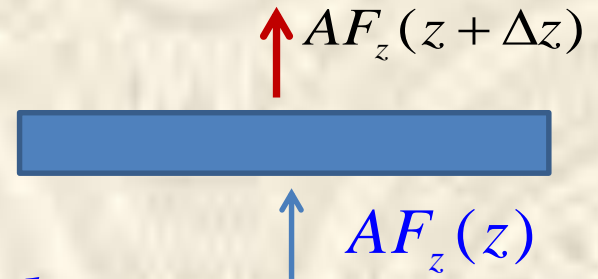
<http://www.razi.ac.ir/sahraei>



Heating rates

Basic ideas

$$A\Delta z$$



$$A[F_z(z) - F_z(z + \Delta z)] \approx -(A\Delta z)dF_z / dz$$

$$-dF_z / dz$$

$$Q = -\frac{1}{\rho(z)} \frac{dF_z}{dz}$$

$$Q / c_p$$

$$F_z (= F^\uparrow - F_\downarrow)$$

Short-wave heating

$$\rho Q_v^{sw} = \rho_a z \quad \chi_v(z) = \int_z^\infty k_v(z') \rho_a(z') dz'$$

$$L_v(s) = \int_0^{\chi_v} J_v(\chi') e^{-(\chi_v - \chi')} d\chi' + L_{v0} e^{-\chi_v}$$

$$F_v^\downarrow(z) = F_{v\infty}^\downarrow e^{-\chi_v(z)} \quad F_{v\infty}^\downarrow e^{-\chi_v(z)}$$

$$\tau_v(z, \infty) = F_v^\uparrow \quad F_{z\nu}^\downarrow(z) = -F_{v\infty}^\downarrow e^{-\chi_v(z)}$$

$$\begin{aligned} \rho Q_v^{sw} &= \frac{d}{dz} (F_{v\infty}^\downarrow e^{-\chi_v(z)}) = F_{v\infty}^\downarrow \left(-\frac{d\chi_v}{dz} \right) e^{-\chi_v(z)} \\ &= F_{v\infty}^\downarrow k_v(z) \rho_a(z) e^{-\chi_v(z)} \end{aligned}$$

$$\rho_a(z) = \rho_a(0)e^{-z/H_a}$$

$$\chi_v(z) = H_a k_v \rho_a(0)e^{-z/H_a} = \chi_v(0)e^{-z/H_a}$$

this shows how the optical depth increases as the solar radiation penetrates downwards, i.e. as z decreases.

Substitution into $F_{zV}(z) = -F_{\infty}^{\downarrow} e^{-\chi_v(z)}$

then gives the vertical irradiance $F_{zV} = -F_{\infty}^{\downarrow} e^{-\chi_v(0)} e^{-z/H_a}$

and differentiation gives the monochromatic volume heating rate,

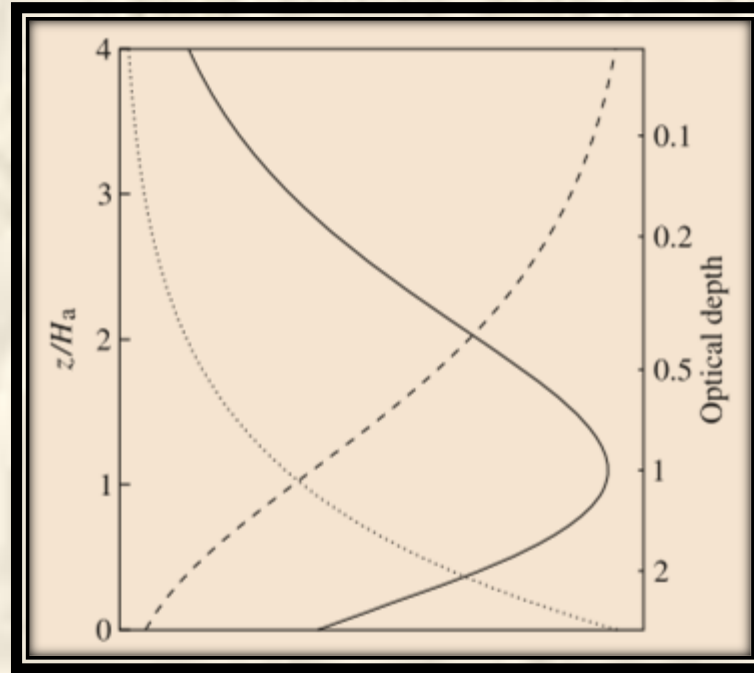
$$\rho Q_v^{sw}(z) = F_{\infty}^{\downarrow} k_v \rho_a(0) e^{-z/H_a - \chi_v(0)} e^{-z/H_a}$$

$$\rho Q_v^{sw}$$

$$F_{z\nu}(z)$$

$$\rho_a(z)$$

$$\rho_a \propto e^{-z/H_a}$$



Long-wave heating and cooling

The upward thermal irradiance at frequency ν and height z is

$$F_{\nu}^{\uparrow}(z) = \pi \int_0^z B_{\nu}(z') \frac{\partial \tau_{\nu}^*(z', z)}{\partial z'} dz' + \pi B_{\nu}(0) \tau_{\nu}^*(0, z)$$

$\tau_{\nu}^*(z', z)$ The spectral transmittance,

$B_{\nu}(0)$ The Planck function

$$J_{\nu} = B_{\nu}$$

Similarly, the downward irradiance is

$$F_{\nu}^{\downarrow}(z) = -\pi \int_z^{\infty} B_{\nu}(z') \frac{\partial \tau_{\nu}^*(z', z)}{\partial z'} dz'$$

The net upward long-wave spectral irradiance

$$F_{z\nu}(z) = F_{\nu}^{\uparrow}(z) - F_{\nu}^{\downarrow}(z)$$

Q_{ν}^{lw} can be calculated using $Q = -\frac{1}{\rho(z)} \frac{dF_z}{dz}$

These heating and cooling terms can also be obtained directly

The derivation of the radiative-transfer

$$\frac{dL_{\nu}}{ds} = -k_{\nu}\rho_a(L_{\nu} - J_{\nu}).$$

shows that the spectral power emitted in a vertical direction from this slab is

$$k_{\nu}\rho_a J_{\nu} A \Delta z$$

where J_{ν} is the source function, equal to B_{ν} under LTE

The fraction of this power that escapes to space is given by the transmittance

$$\tau_v(z, \infty) = \exp\left(-\int_z^{\infty} k_v \rho_a dz'\right)$$

Noting that
$$\frac{\partial \tau_v(z, \infty)}{\partial z} = k_v(z) \rho_a(z) \tau_v(z, \infty)$$

we find that the power escaping to space from the slab in a purely vertical beam is

$$B_v(z) \frac{\partial \tau_v(z, \infty)}{\partial z} A \Delta x$$

Now, integrating over all slanting paths as above and replacing τ_v by τ_v^* we obtain a contribution to the heating rate per unit mass

$$Q_v^{cts}(z) = \frac{\pi B_v(z)}{\rho(z)} \frac{\partial \tau_v^*(z, \infty)}{\partial z}$$

A useful simplification for some purposes is the cooling-to-space approximation, in which the loss of photon energy to space dominates the other contributions; therefore, under this approximation,

$$Q_{\nu}^{lw} \approx Q_{\nu}^{cts}$$

All gases that absorb and emit at frequency ν must in principle be included in τ_{ν}^*

Then $Q_{\nu}^{lw}(z)$ must be integrated over all relevant frequencies to obtain the total long-wave cooling $Q^{lw}(z)$

Net radiative heating rates

$$Q^{sw} / c_p$$

$$-Q^{lw} / c_p$$

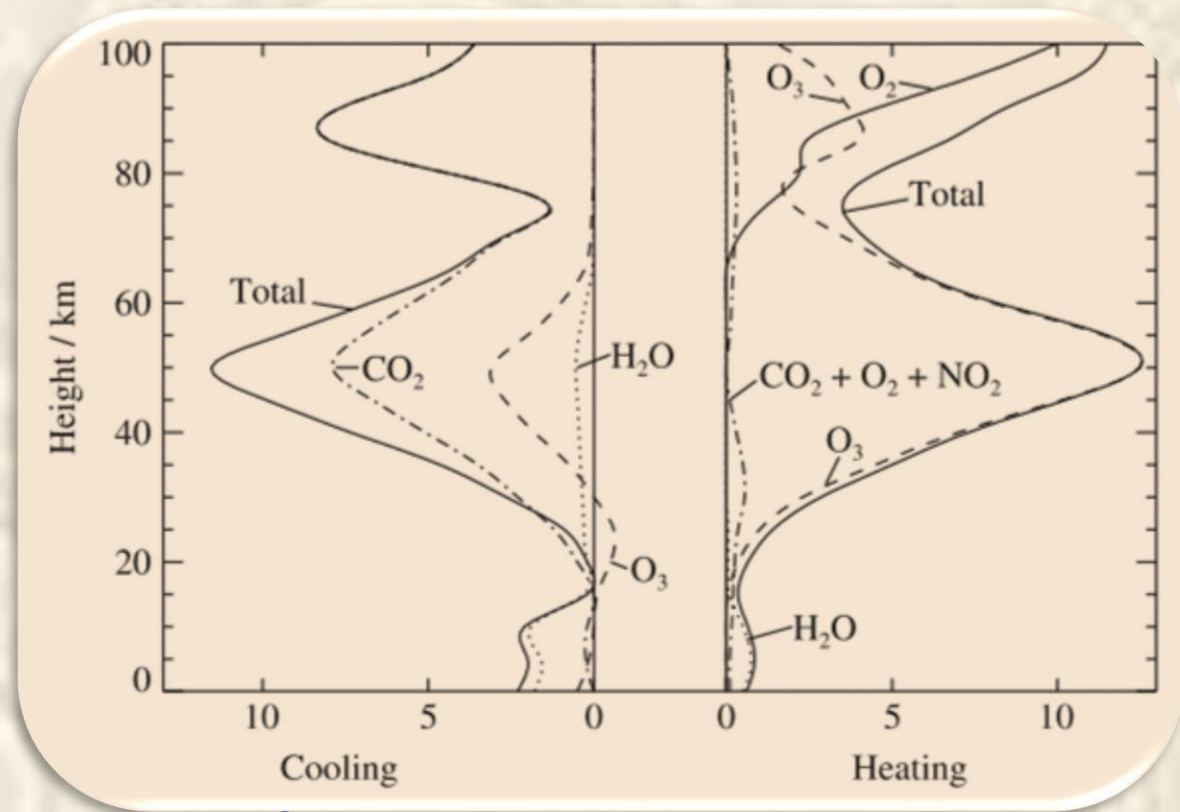
$$Q = Q^{sw} + Q^{lw}$$

$$Q = 0$$

$$T_r(r)$$

$$Q(T_r(r)) = 0$$

$$T = T_r + \delta T$$



$$Q(T_r + \delta T) \approx Q(T_r) + \delta T \left. \frac{\partial Q}{\partial T} \right|_{T=T_r} = \delta T \left. \frac{\partial Q}{\partial T} \right|_{T=T_r}$$

$$Q(T_r) = 0 \qquad = -c_p \frac{\delta T}{\tau_r}$$

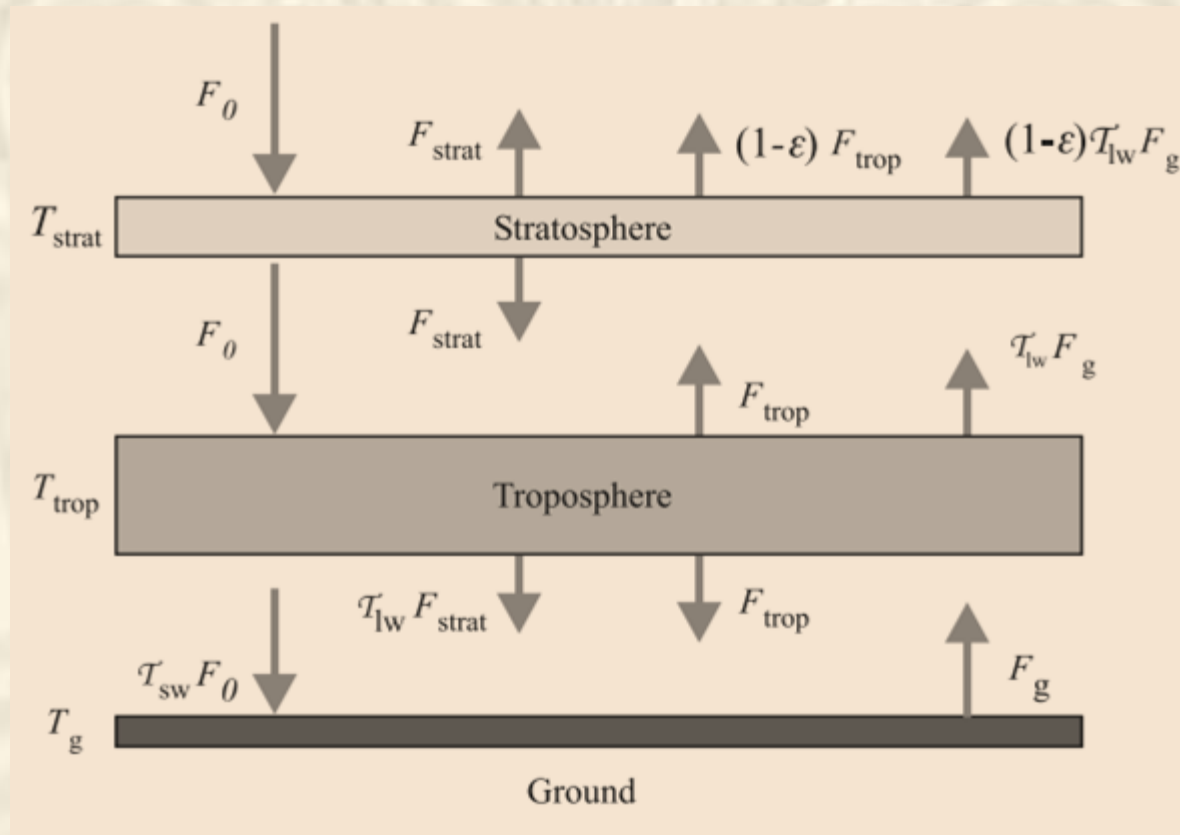
$$\tau_r = c_p \left(\left. \frac{\partial Q}{\partial T} \right|_{T=T_r} \right)^{-1}$$

The greenhouse effect revisited

Two-layer atmosphere in radiative equilibrium, including an optically thin stratosphere

$$T_{trop} \quad \tau_{sw} \quad \tau_{lw}$$

$$T_c \equiv \left(\frac{F_0}{\sigma}\right)^{1/4} \approx 255 K$$



$$F_0 = F_{strat} + (1 + \epsilon)(F_{trop} + \tau_{lw}F_g)$$

$$F_{strat} = \sigma \epsilon T_{strat}^4, \quad F_{trop} = \sigma (1 - \tau_{lw}) T_{trop}^4, \quad F_g = \sigma T_g^4$$

$$F_0 + F_{strat} = F_{trop} + \tau_{lw} F_g$$

$$2F_{strat} = \varepsilon(F_{trop} + \tau_{lw} F_g)$$

$$F_{trop} + \tau_{lw} F_g$$

$$F_0 + F_{strat} = (1 - \varepsilon)(F_0 + F_{strat})$$

$$\sigma \varepsilon T_{strat}^4 = F_{strat} = \frac{\varepsilon F_0}{2 - \varepsilon}$$

$$\varepsilon \ll 1 \quad \sigma T_{strat}^4 \approx \frac{F_0}{2} \quad T_{strat} \approx \frac{T_c}{2^{1/4}} = 214K$$

$$F_{trop} + \tau_{lw} F_g$$

$$F_{trop} = \frac{2F_0}{2 - \varepsilon} - \tau_{lw} F_g$$

$$\tau_{sw} F_0 + F_{lw} F_{strat} + F_{trop} = F_g$$

Continuously stratified atmosphere in radiative equilibrium

$$-\frac{dF^\uparrow}{d\chi^*} + F^\uparrow = \pi B(T)$$

$$\pi B(T) = \sigma T^4$$

$$\frac{dF^\downarrow}{d\chi^*} + F^\downarrow = \pi B(T)$$

$$Q^{sw} = 0$$

$$Q^{lw} = 0$$

$$F_z = F^\uparrow - F^\downarrow = \text{constant} \quad F^\downarrow(0) = 0 \quad F_z = F^\uparrow(0)$$

$$F_z = F^\uparrow - F^\downarrow = F_0$$

$$-\frac{d}{d\chi^*} (F^\uparrow - F^\downarrow) + F^\uparrow - F^\downarrow = 2\pi B(T)$$

$$\pi B(T) = \frac{1}{2} (F^\uparrow + F^\downarrow)$$

$$\frac{d}{d\chi^*} (F^\uparrow + F^\downarrow) = F^\uparrow - F^\downarrow = F_0$$

$$F^{\uparrow} + F^{\downarrow} = F_0 \chi^* + \text{constant}$$

$$F^{\uparrow} + F^{\downarrow} = F_0(1 + \chi^*)$$

$$F^{\uparrow} = \frac{1}{2} F_0(2 + \chi^*)$$

$$F^{\downarrow} = \frac{1}{2} F_0 \chi^*$$

$$\pi B(T) = \sigma T^4 = \frac{1}{2} F_0(1 + \chi^*)$$

$$F_0(1 + \chi_g^*)/2$$

$$\pi B(T_g) = \sigma T_g^4$$

$$\sigma T_g^4 = F_0 \left(1 + \frac{1}{2} \chi_g^*\right) = \sigma T_c^4 \left(1 + \frac{1}{2} \chi_g^*\right)$$

$$T_c \approx 255K \quad \chi_g^* > 0$$

$$T_g > T_c$$

$$\rho_c(z) = \rho_a(0) e^{-z/H_a}$$

$$\chi^*(z) = \chi_g^* e^{-z/H_a}$$

$$F^\uparrow(z) = \frac{1}{2} F_0 (2 + \chi_g^* e^{-z/H_a}) \quad F^\downarrow(z) = \frac{1}{2} F_0 \chi_g^* e^{-z/H_a}$$

$$T(z) = \left[\frac{F}{2\sigma_0} (1 + \chi_g^* e^{-z/H_a}) \right]^{1/4}$$

$$z / H_a$$

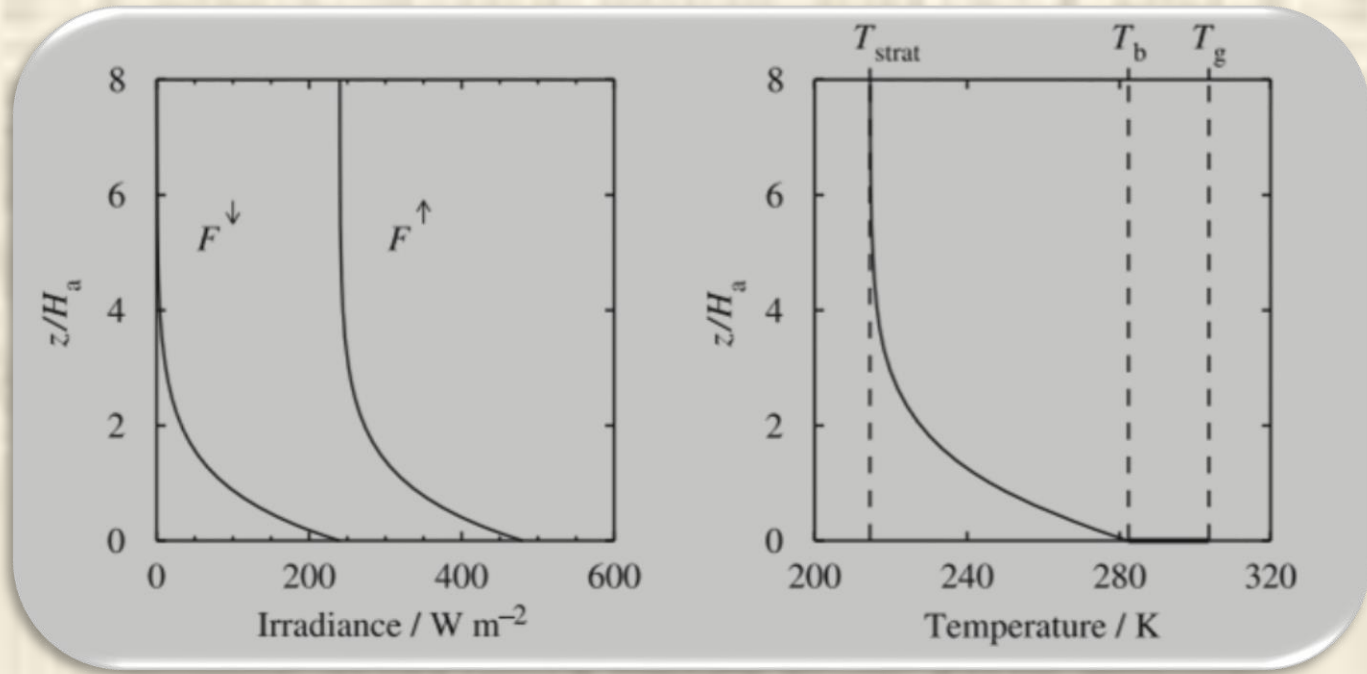
$$\chi_g^* = 2$$

$$F_0 = 240 \text{ W/m}^2$$

$$T \rightarrow \left(\frac{F}{2\sigma}\right)^{1/4} \text{ as } z \rightarrow \infty$$

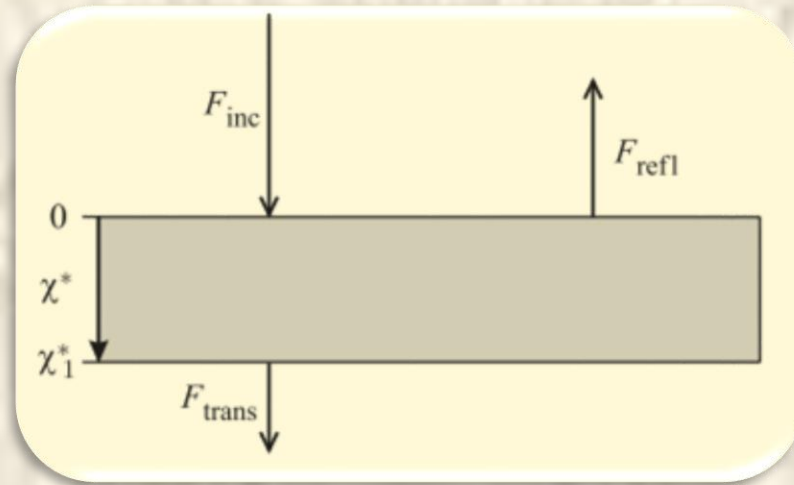
$$T_{strat} = 2^{-1/4} T_c$$

$$T(z) \rightarrow T_b \equiv T_c \left(\frac{1 + \chi_g^*}{2}\right)^{1/4} \text{ as } z \downarrow 0$$



$$T_g \equiv T_c \left(\frac{2 + \chi_g^*}{2} \right)^{1/4}$$

A simple model of scattering



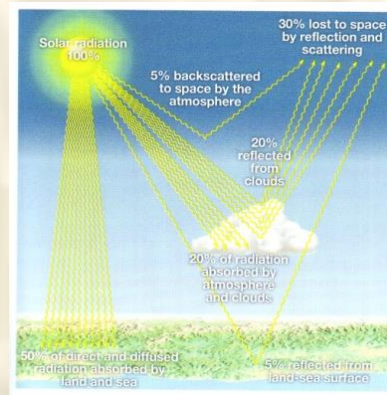
Shortwave Radiation

$S_0 = 1368 \text{ w m}^{-2}$ is the **solar constant** for Earth

Insolation

$$R_0 = S_0 \left(\frac{d_m}{d} \right)^2 \cos \gamma$$

$$I_0 = \int_{t_1}^{t_2} R_0(t) dt$$



Stefan-Boltzmann Law

This law expresses the rate of radiation emission per unit area

$$R = \sigma T^4 \quad \sigma = 5.67 \times 10^{-8} \text{ W / m}^2 \text{ K}^4$$

Compare the difference between the radiation emission from the sun and the Earth.

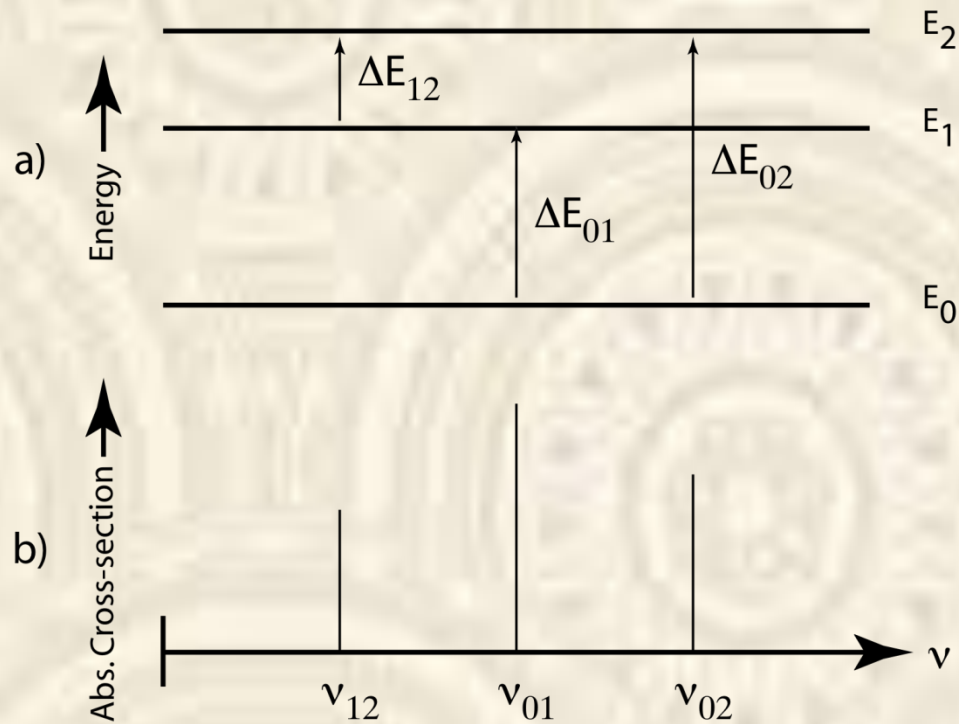
The sun with an average temperature of 6000 K emits 73,483,200 W/m²

By contrast, Earth with an average temperature of 300 K emits 459 W/m²

The sun has a temperature 20 times higher than Earth and thus emits about 160,000 times more radiation

This makes sense, $20^4 = 160,000$

Absorption spectra of molecules



Hypothetical molecule
with three allowed
energy levels

Note relationship to
emission!

$$\nu_{ij} = \Delta E_{ij}/h$$

a) allowed transitions

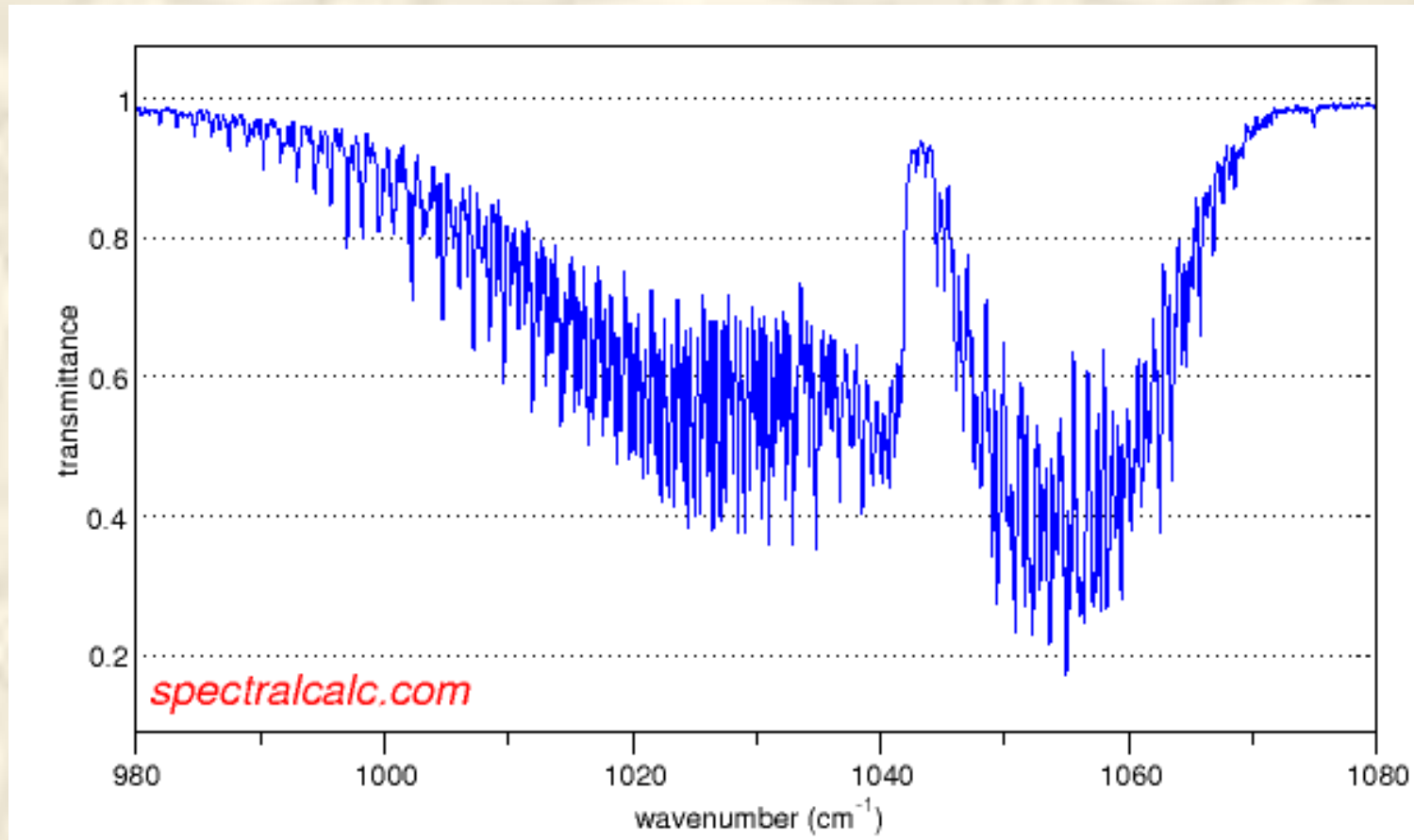
b) positions of the absorption lines in the spectrum of the molecule

Line positions are determined by the **energy changes** of allowed transitions

Line strengths are determined by the **fraction of molecules** that are in a particular initial state required for a transition

Multiple **degenerate** transitions with the same energy may combine

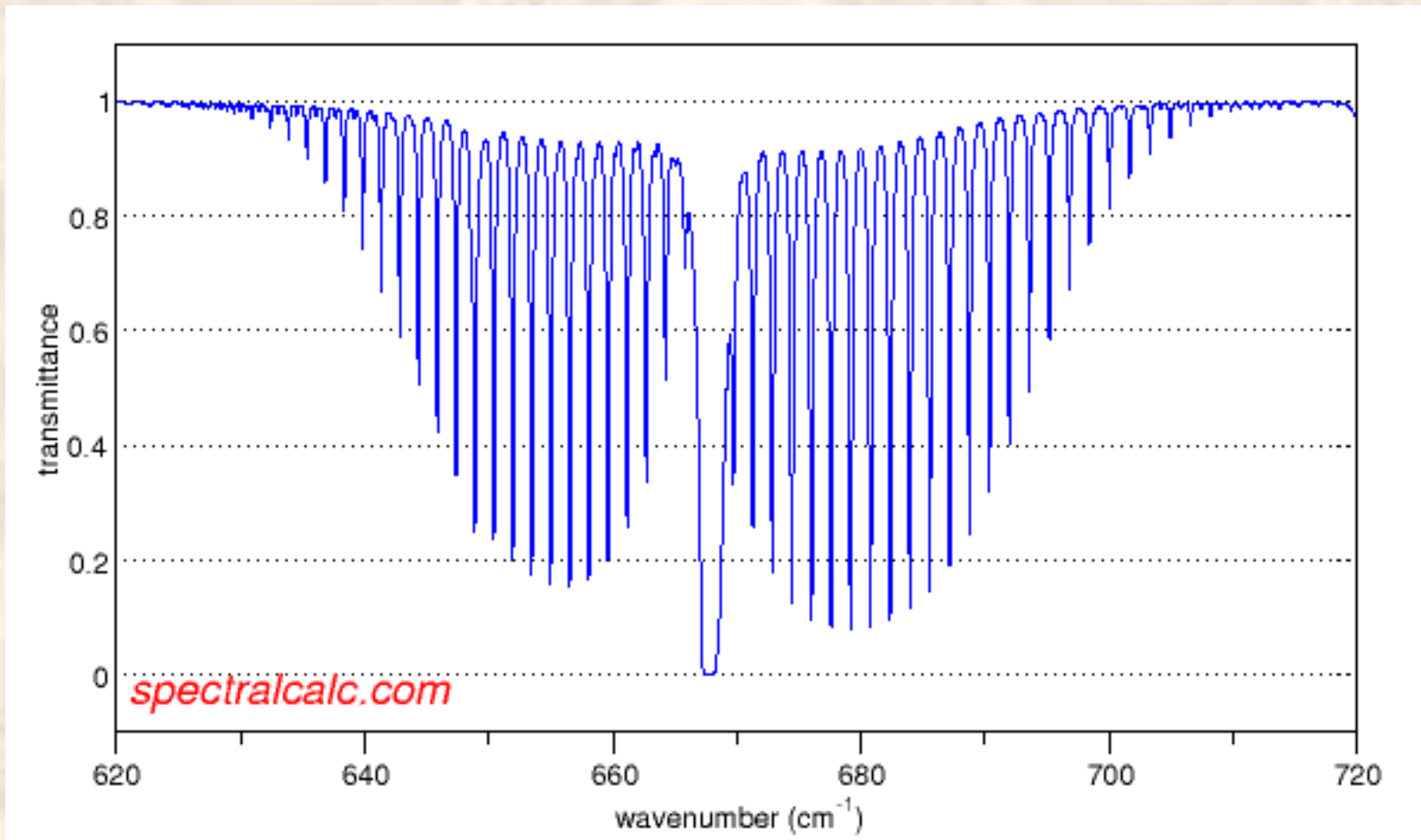
Transmittance spectrum for ozone (O_3)



<http://www.spectralcalc.com/calc/spectralcalc.php>

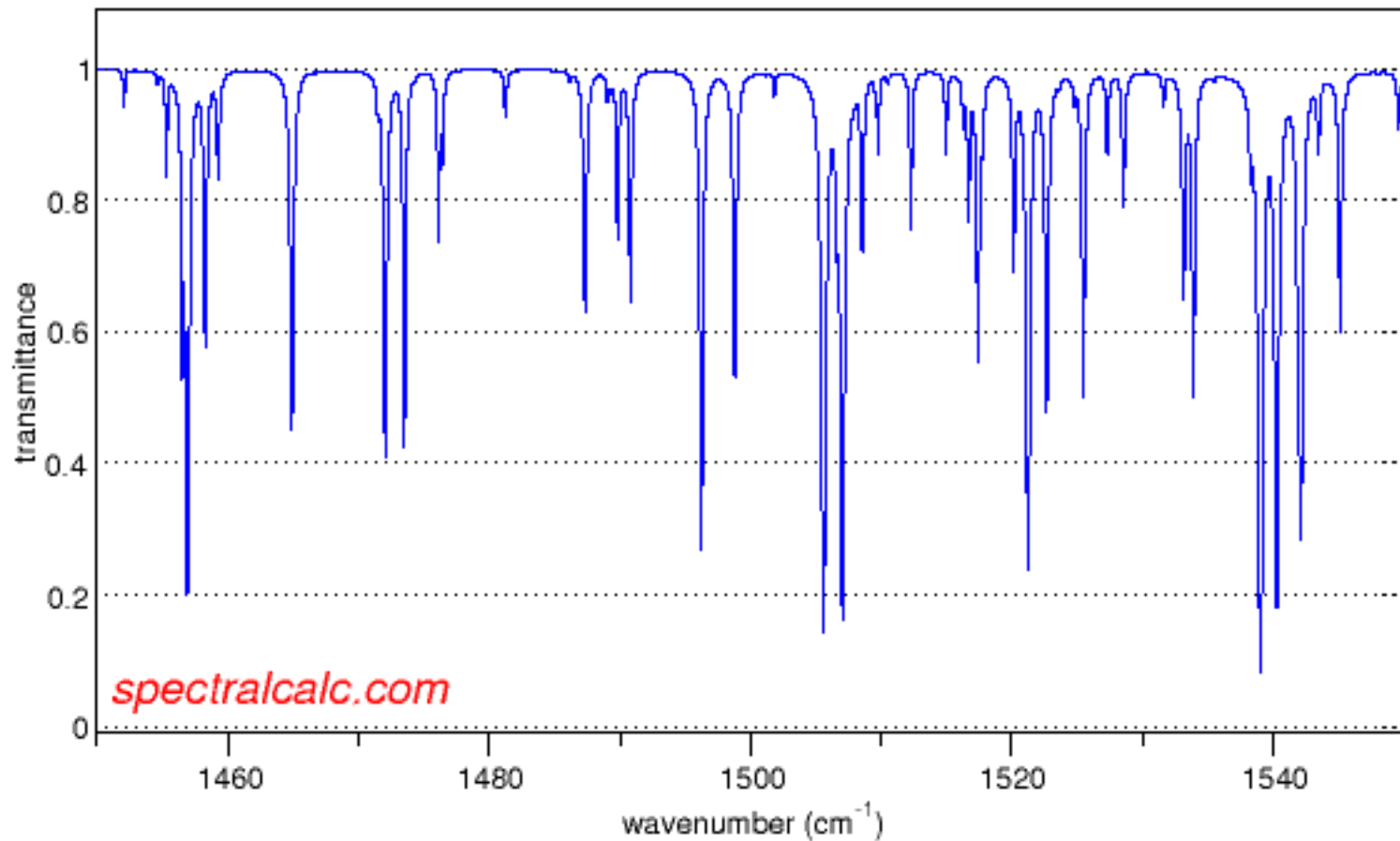
Transmittance spectrum for CO₂

<http://www.spectralcalc.com/calc/spectralcalc.php>



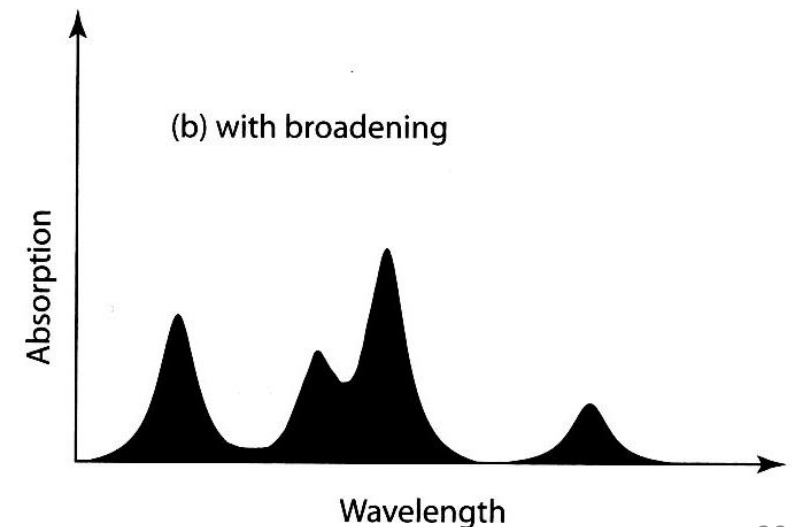
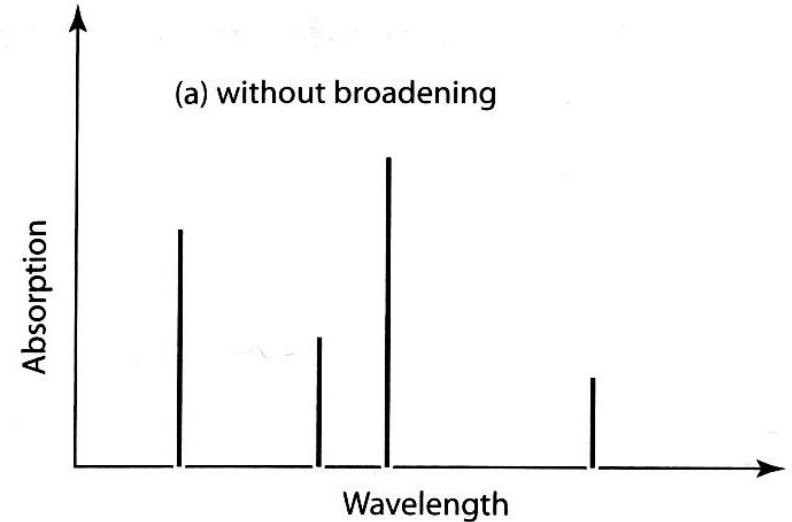
Transmittance spectrum for H₂O

<http://www.spectralcalc.com/calc/spectralcalc.php>



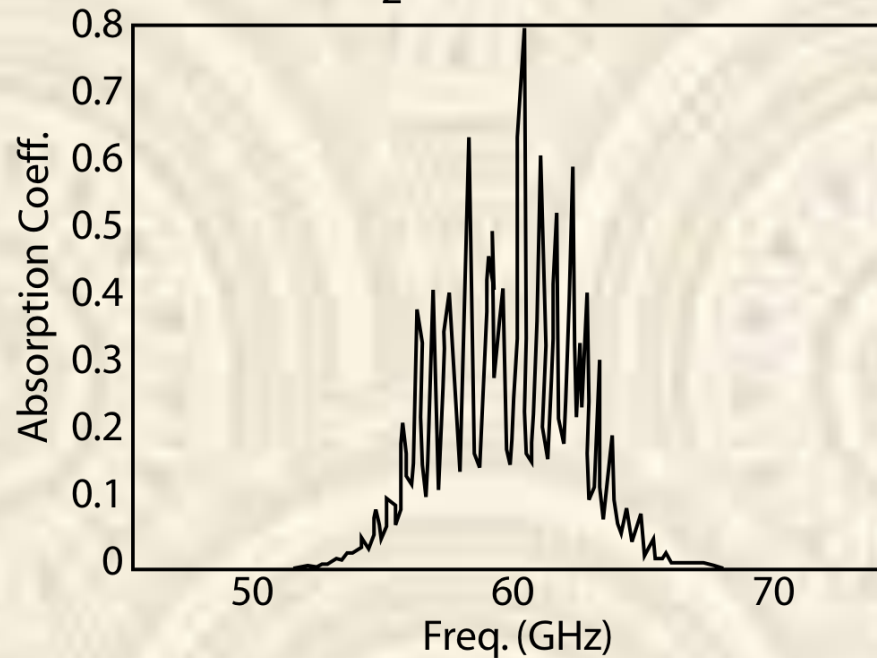
Absorption line shapes

- **Doppler broadening:** random translational motions of individual molecules in any gas leads to Doppler shift of absorption and emission wavelengths (important in upper atmosphere)
- **Pressure broadening:** collisions between molecules randomly disrupt natural transitions between energy states, so that absorption and emission occur at wavelengths that deviate from the natural line position (important in troposphere and lower stratosphere)
- Line broadening closes gaps between closely spaced absorption lines, so that the atmosphere becomes opaque over a continuous wavelength range.

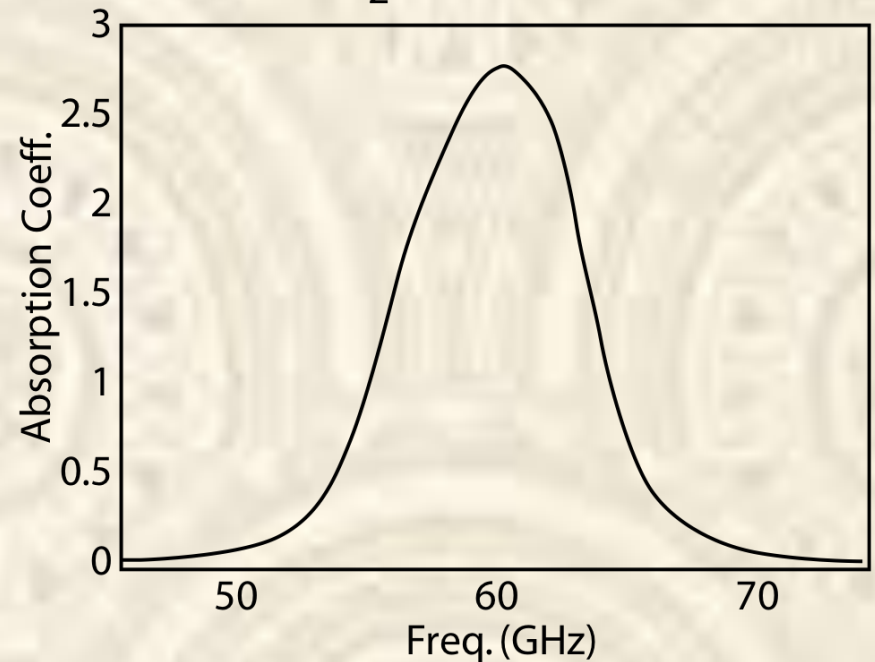


Pressure broadening

a) O₂ at 100 mb pressure



b) O₂ at 1000 mb pressure



- Absorption coefficient of O₂ in the microwave band near 60 GHz at two different pressures. Pressure broadening at 1000 mb obliterates the absorption line structure.