



Atmospheric Physics

Lecture 7

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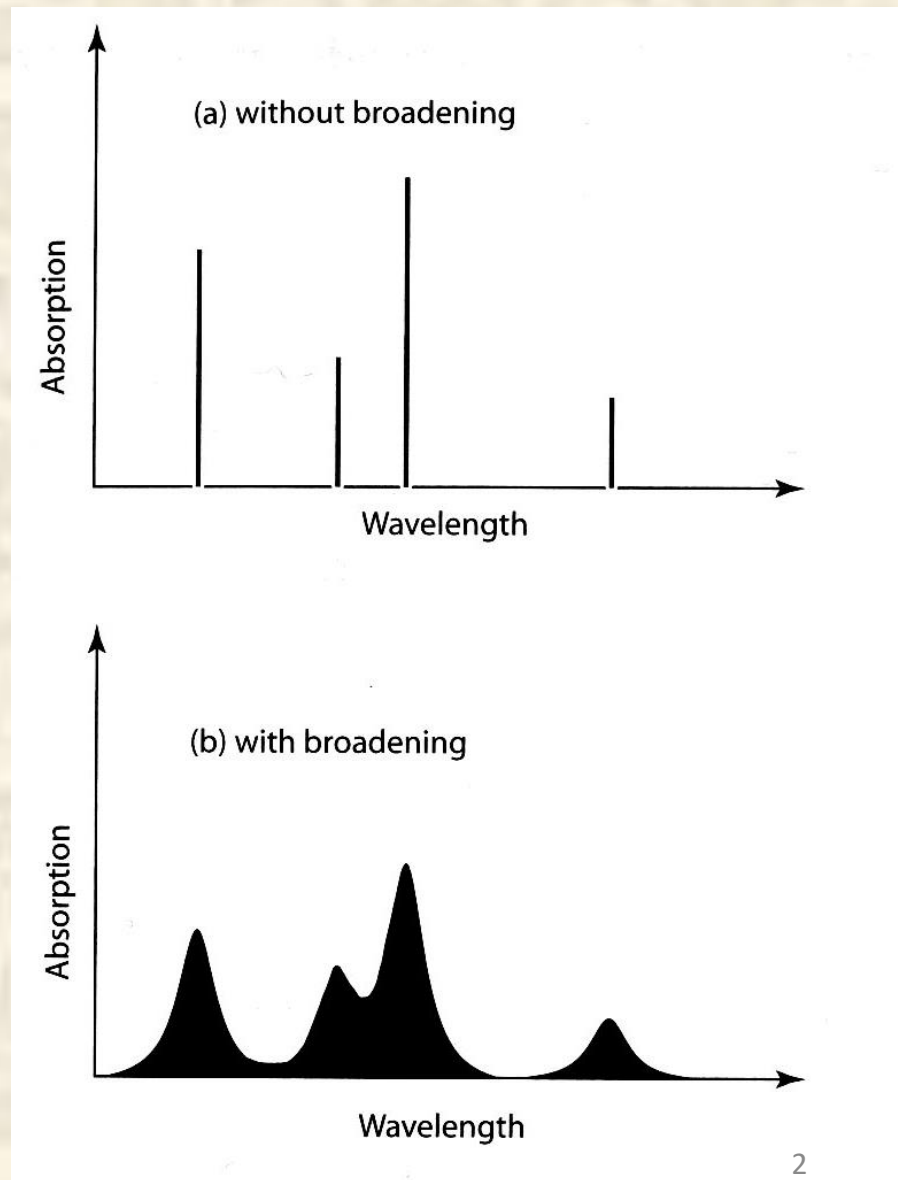
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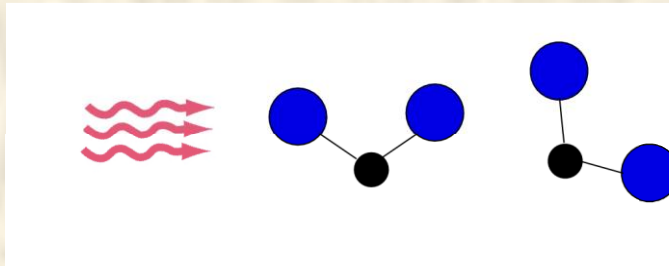
<http://www.razi.ac.ir/sahraei>

Absorption line shapes

Doppler broadening: random translational motions of individual molecules in any gas leads to Doppler shift of absorption and emission wavelengths (important in upper atmosphere)

Pressure broadening: collisions between molecules randomly disrupt natural transitions between energy states, so that absorption and emission occur at wavelengths that deviate from the natural line position (important in troposphere and lower stratosphere)





Molecules absorb radiation at particular wavelengths, depending on amount of energy required to cause vibration or rotation of atomic bond.

Two essential things for the greenhouse effect:

The Earth's atmosphere is mostly transparent to visible radiation (why not totally)

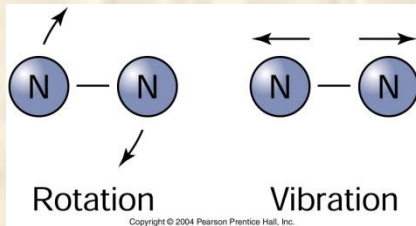
The Earth's atmosphere is mostly opaque to infrared radiation.

The composition of the Earth's atmosphere

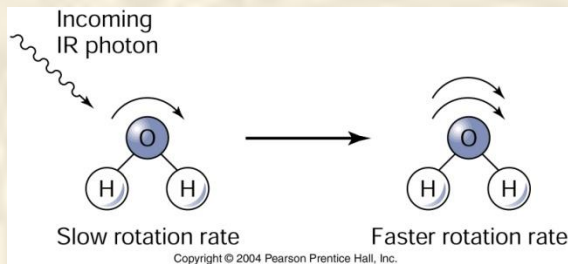
Major Constituents of Earth's Atmosphere Today

<i>Name and Chemical Symbol</i>	<i>Concentration (% by volume)</i>
Nitrogen, N ₂	78
Oxygen, O ₂	21
Argon, Ar	0.9
Water vapor, H ₂ O	0.00001 (South Pole)–4 (tropics)
Carbon dioxide, CO ₂	0.037*

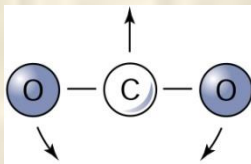
(Plus other trace components, e.g. methane, CFCs, ozone)



Bi-atomic molecules (O₂, N₂) can only absorb high energy photons, meaning ultraviolet wavelengths and shorter.



Tri-atomic molecules (H₂O, CO₂) can absorb lower energy photons, with wavelengths in the infrared



Radiation in the Atmosphere

Deviations from blackbody due to absorption by the solar atmosphere, absorption and scattering by the earth's atmosphere (below).

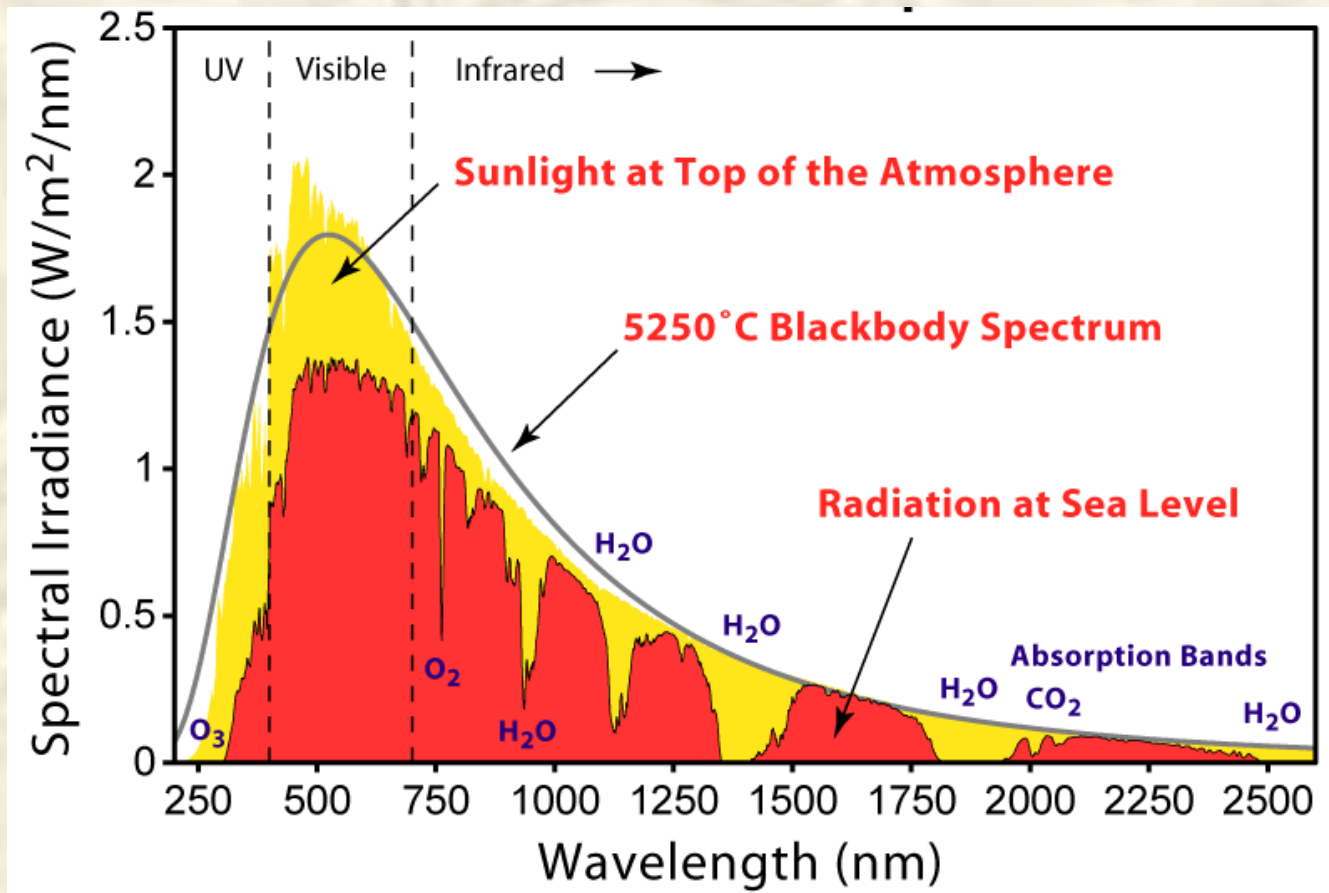
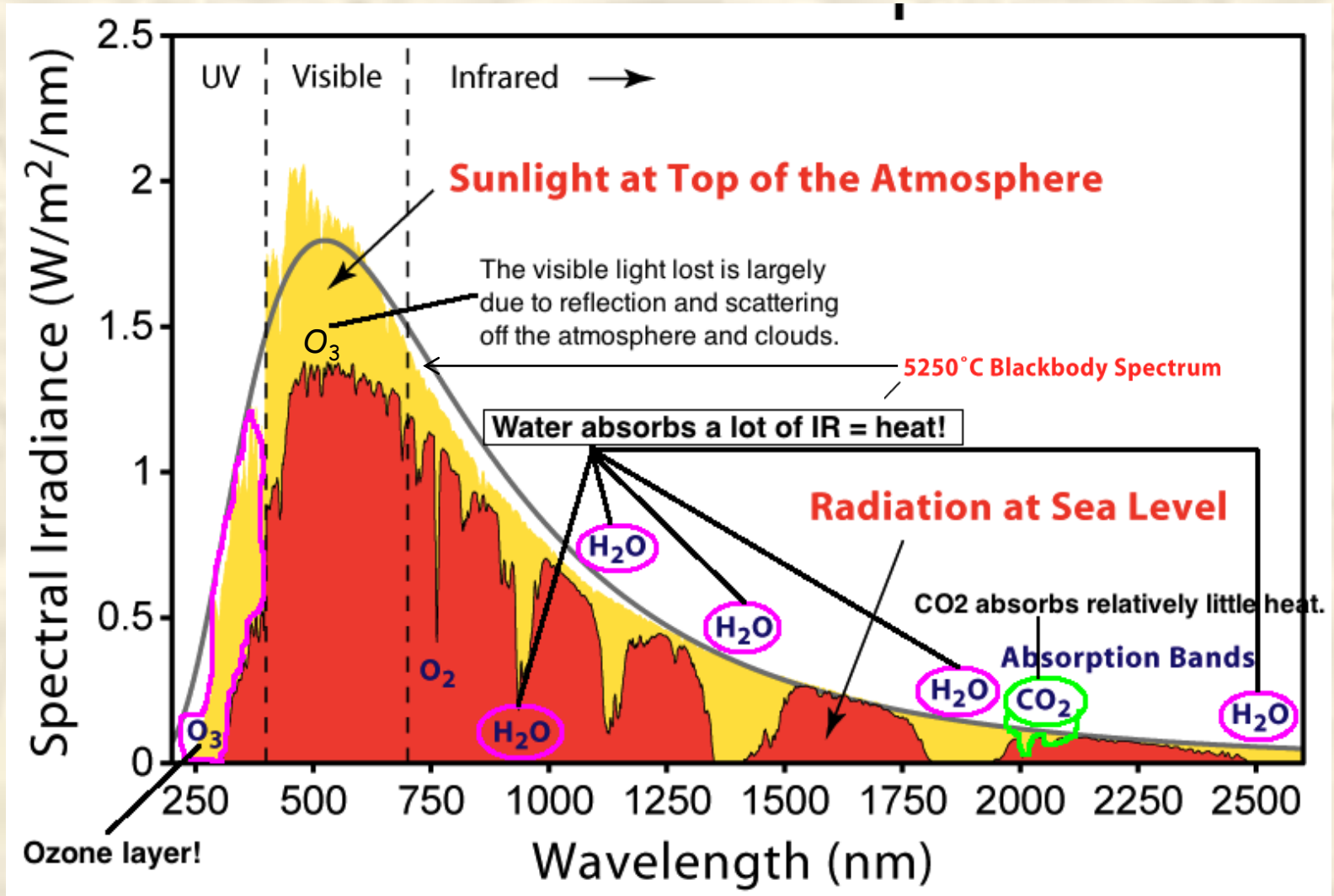


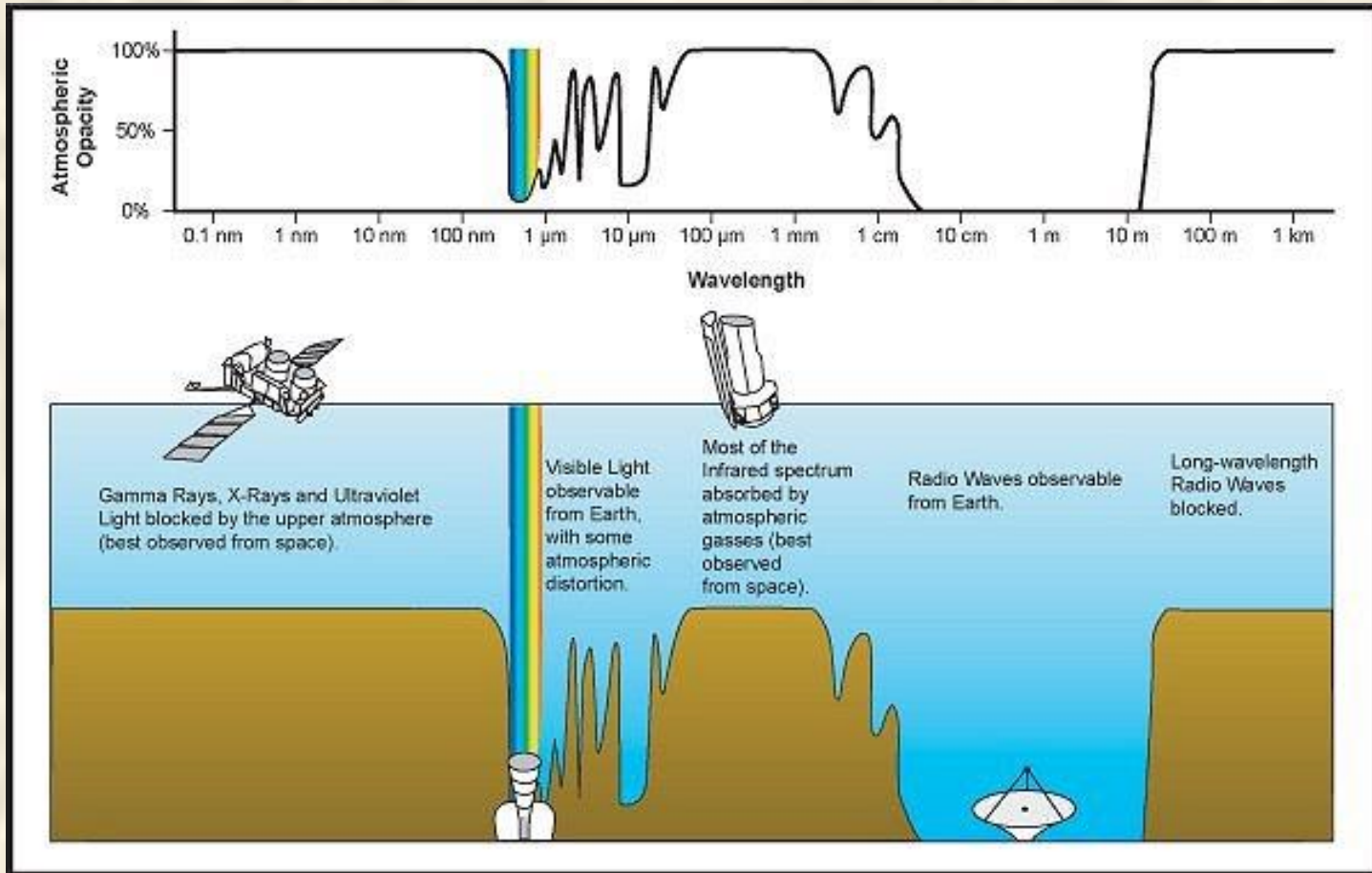
Fig. 3.13

Absorption by atmospheric gases

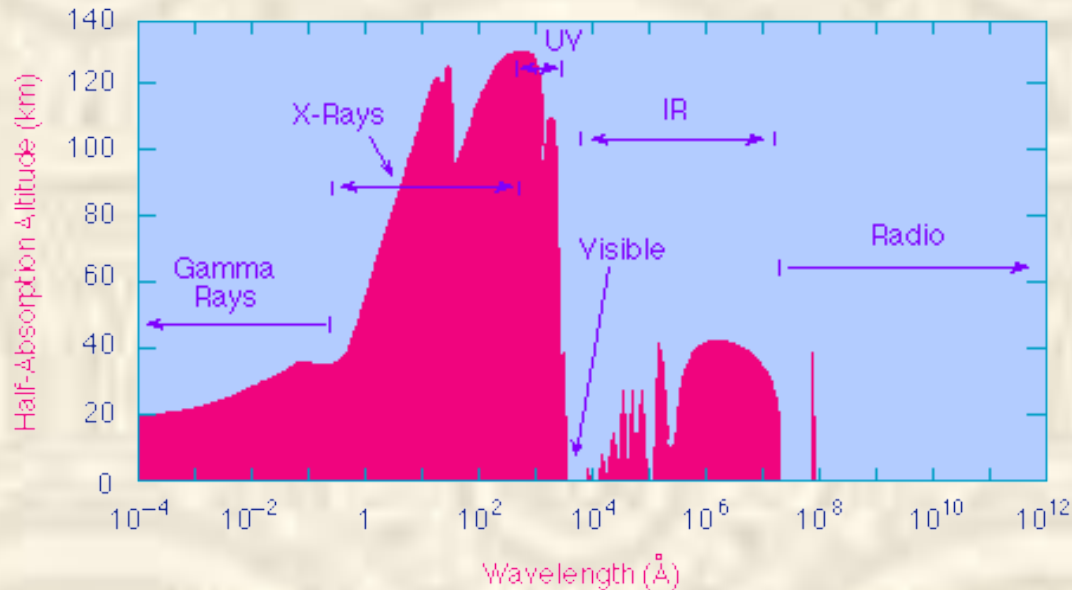
The solar spectrum



Atmospheric Windows



Atmospheric Windows



The dominant windows in the atmosphere are seen to be in the **visible and radio frequency regions**, while X-Rays and UV are seen to be very strongly absorbed and Gamma Rays and IR are somewhat less strongly absorbed.

Infrared Windows in the Atmosphere

Wavelength Range	Sky Transparency	Sky Brightness
1.1 - 1.4 microns	high	low at night
1.5 - 1.8 microns	high	very low
2.0 - 2.4 microns	high	very low
3.0 - 4.0 microns	3.0 - 3.5 microns: fair 3.5 - 4.0 microns: high	low
4.6 - 5.0 microns	low	high
7.5 - 14.5 microns	8 - 9 microns and 10 - 12 microns: fair others: low	very high
17 - 40 microns	very low	very high
330 - 370 microns	very low	low

The Primary Greenhouse Gases

Infra-red absorption

CH₄

- Fundamental modes

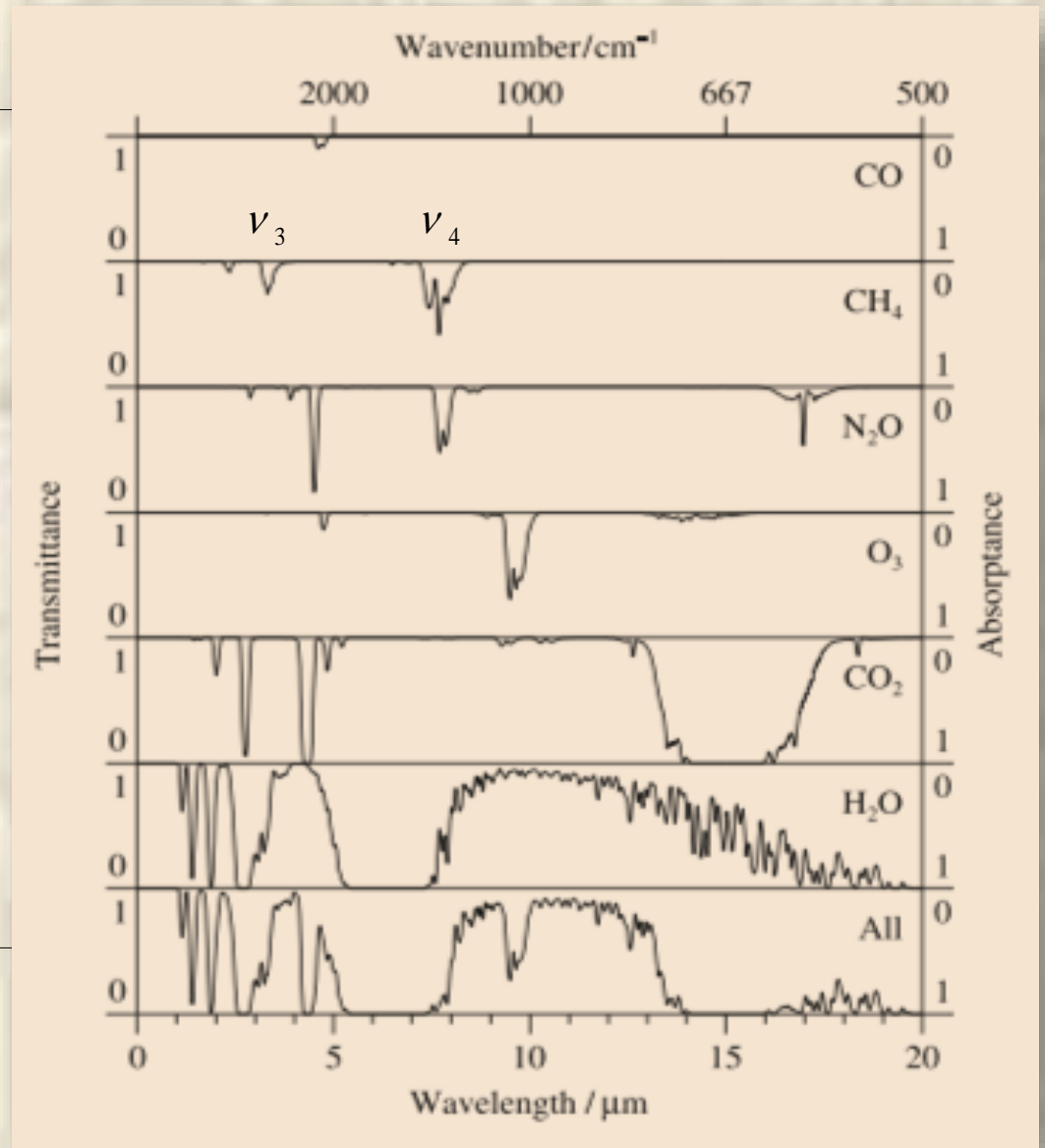
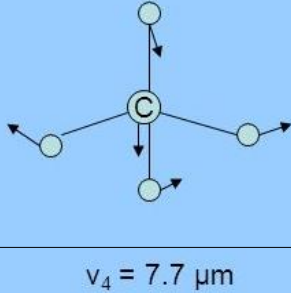
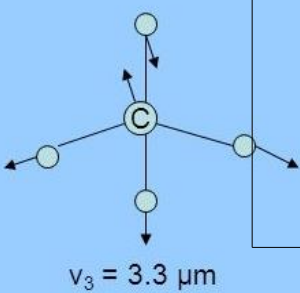
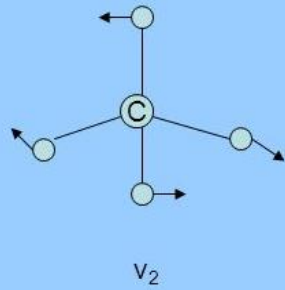
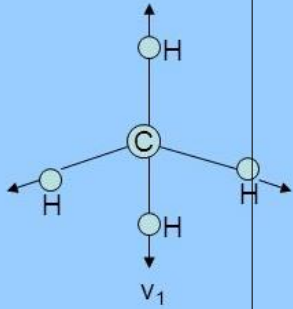


Fig. 3.14

Infra-red absorption

Nitrous oxide (N_2O)

Fundamental modes

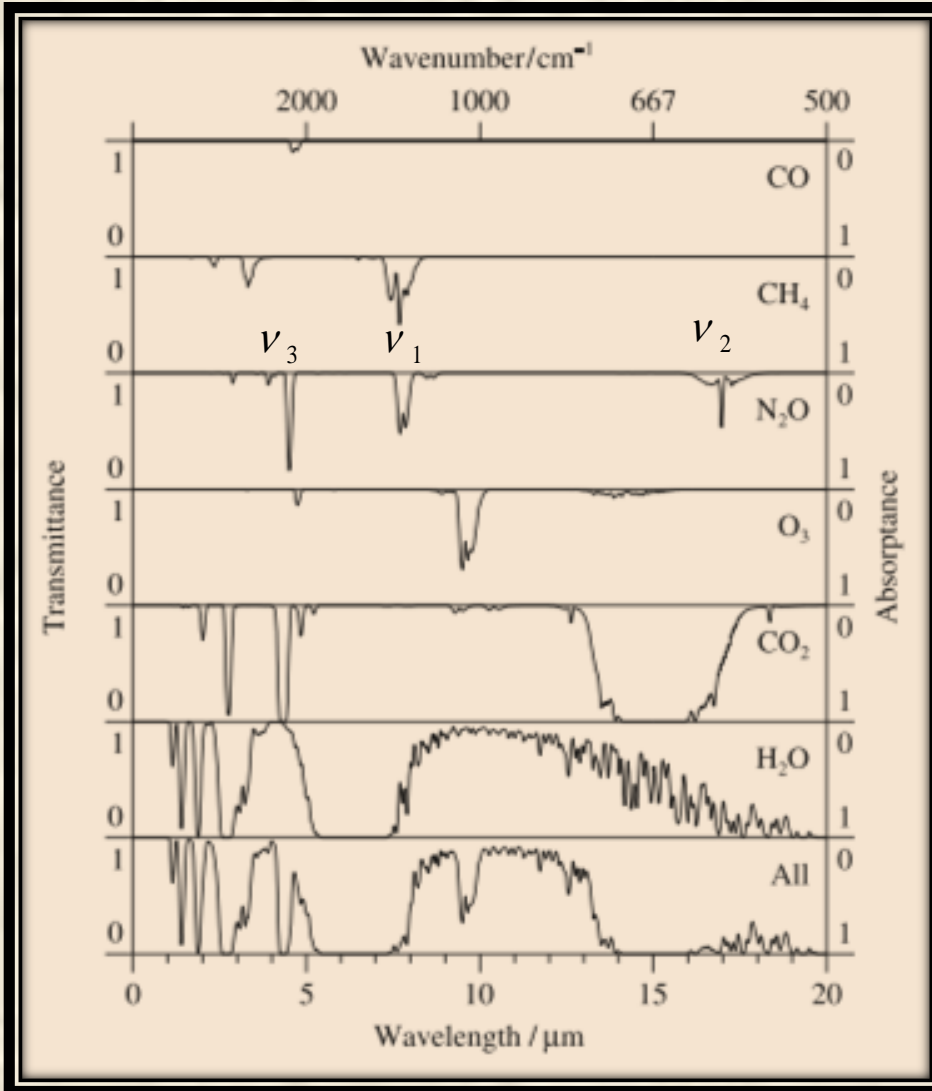
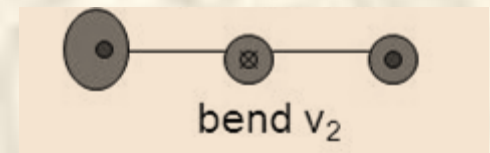
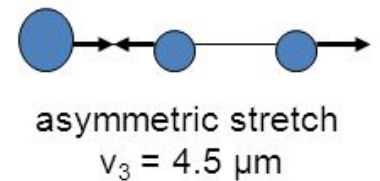
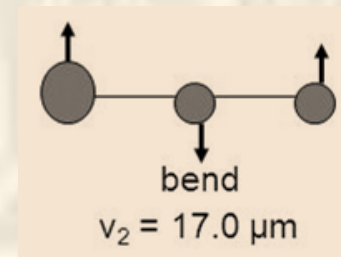
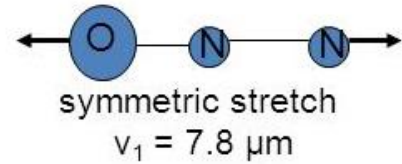


Fig. 3.14

Infra-red absorption

Ozone (O_3)

Fundamental vibrational modes

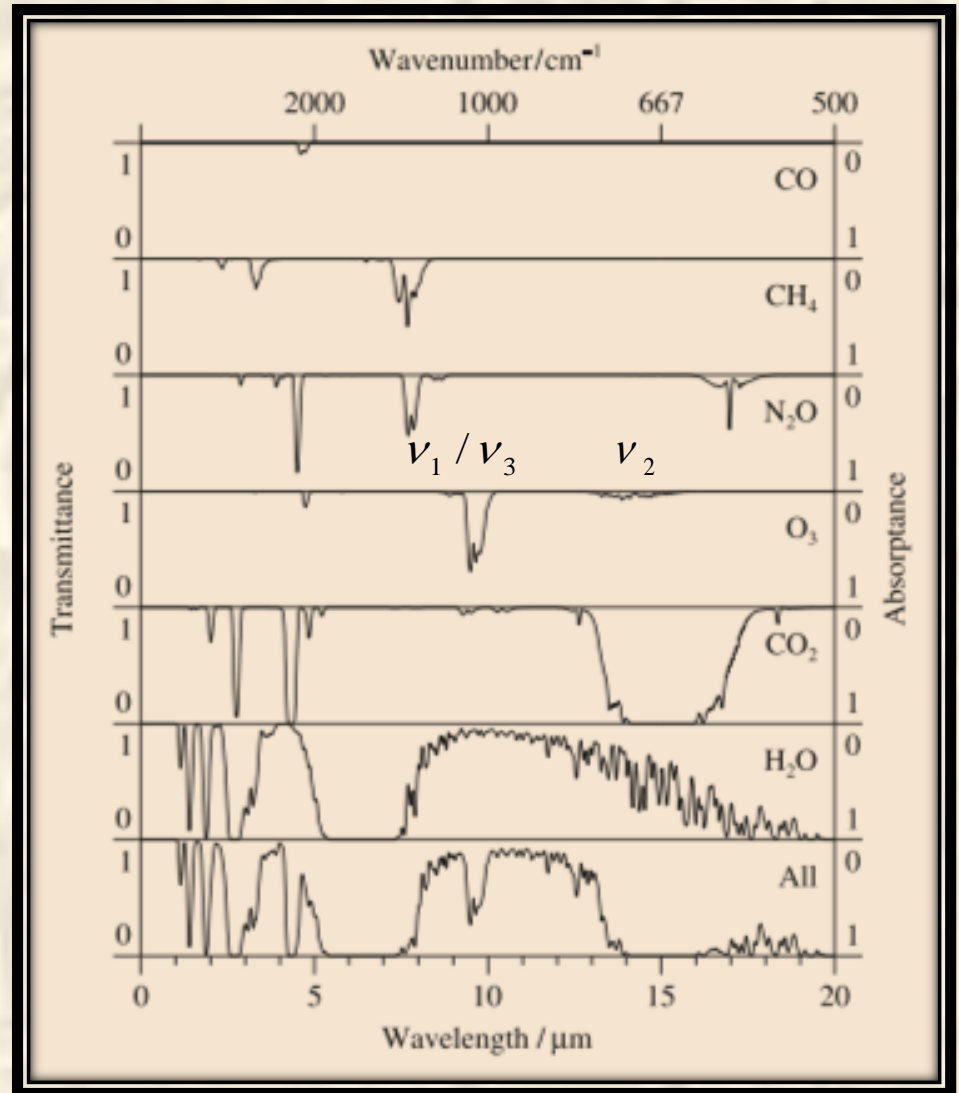
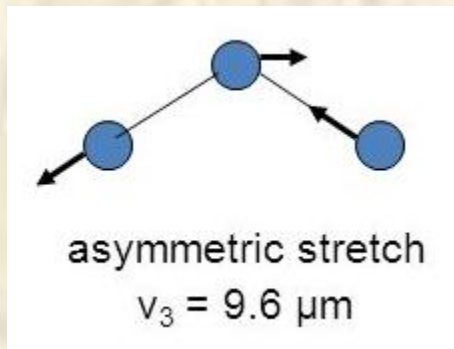
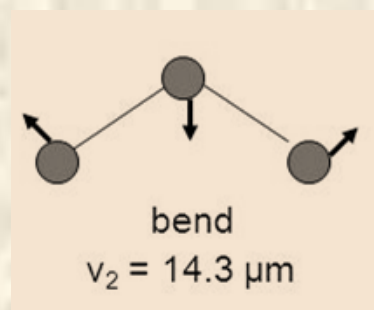
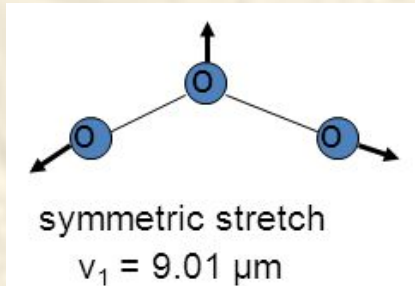
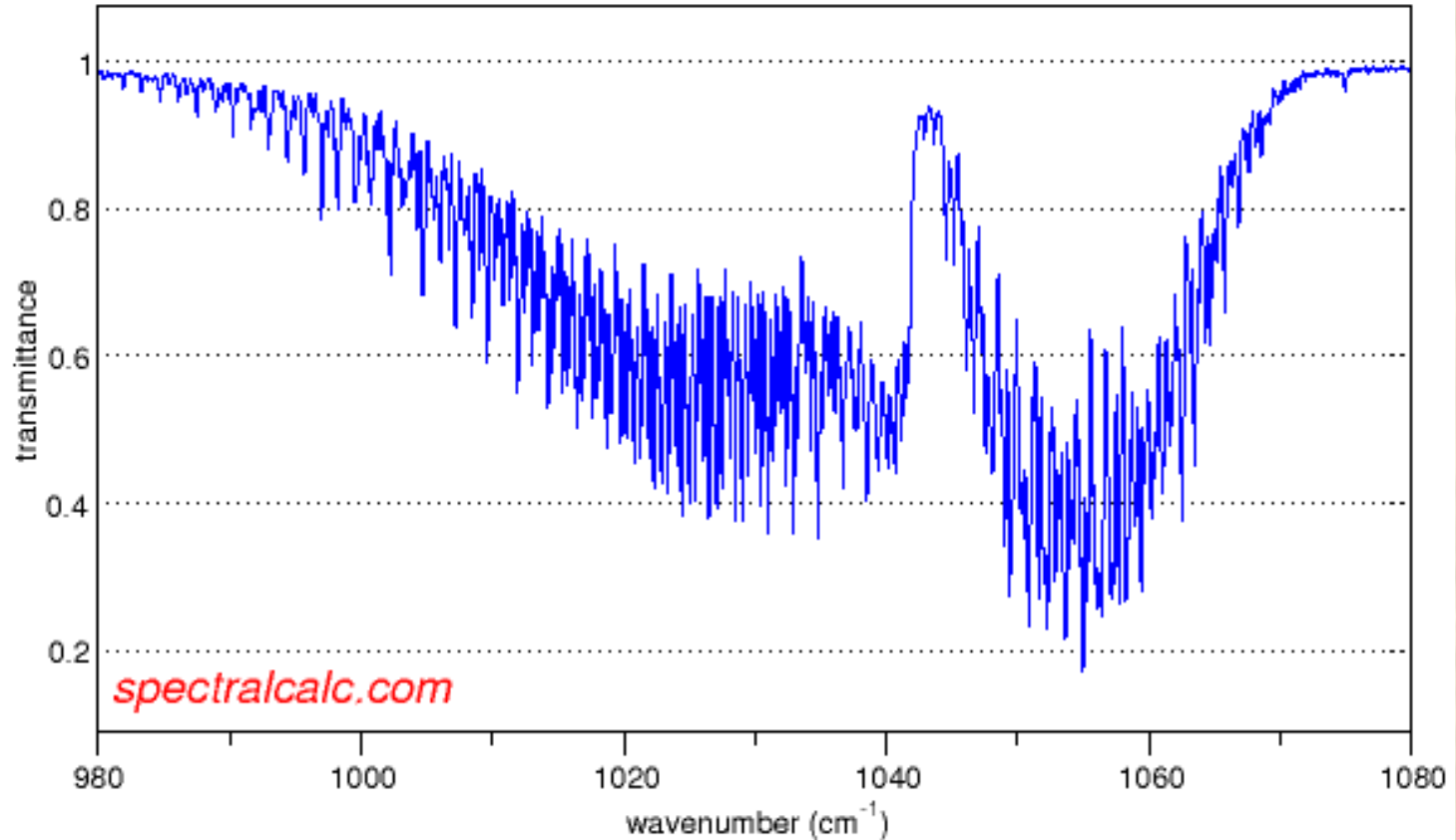


Fig. 3.14

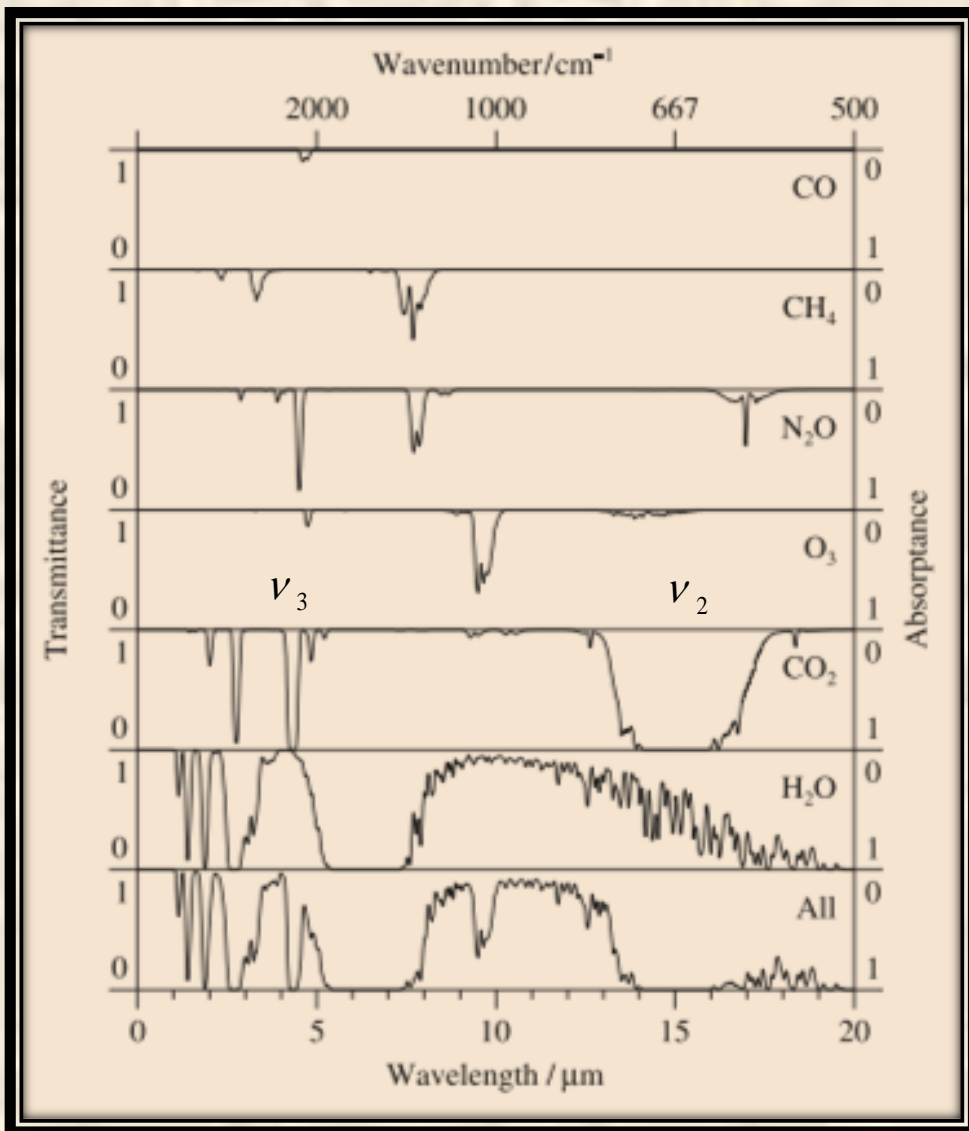
Transmittance spectrum for ozone (O_3)



<http://www.spectralcalc.com/calc/spectralcalc.php>

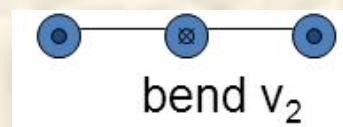
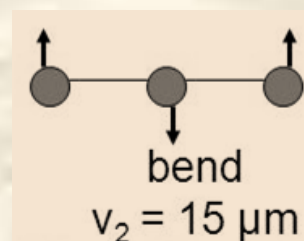
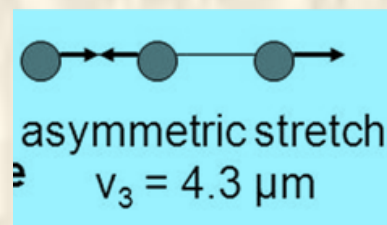
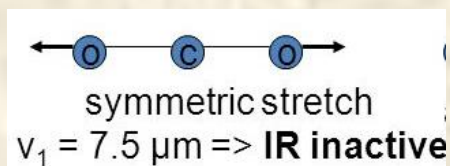
Infra-red absorption

Fig. 3.14

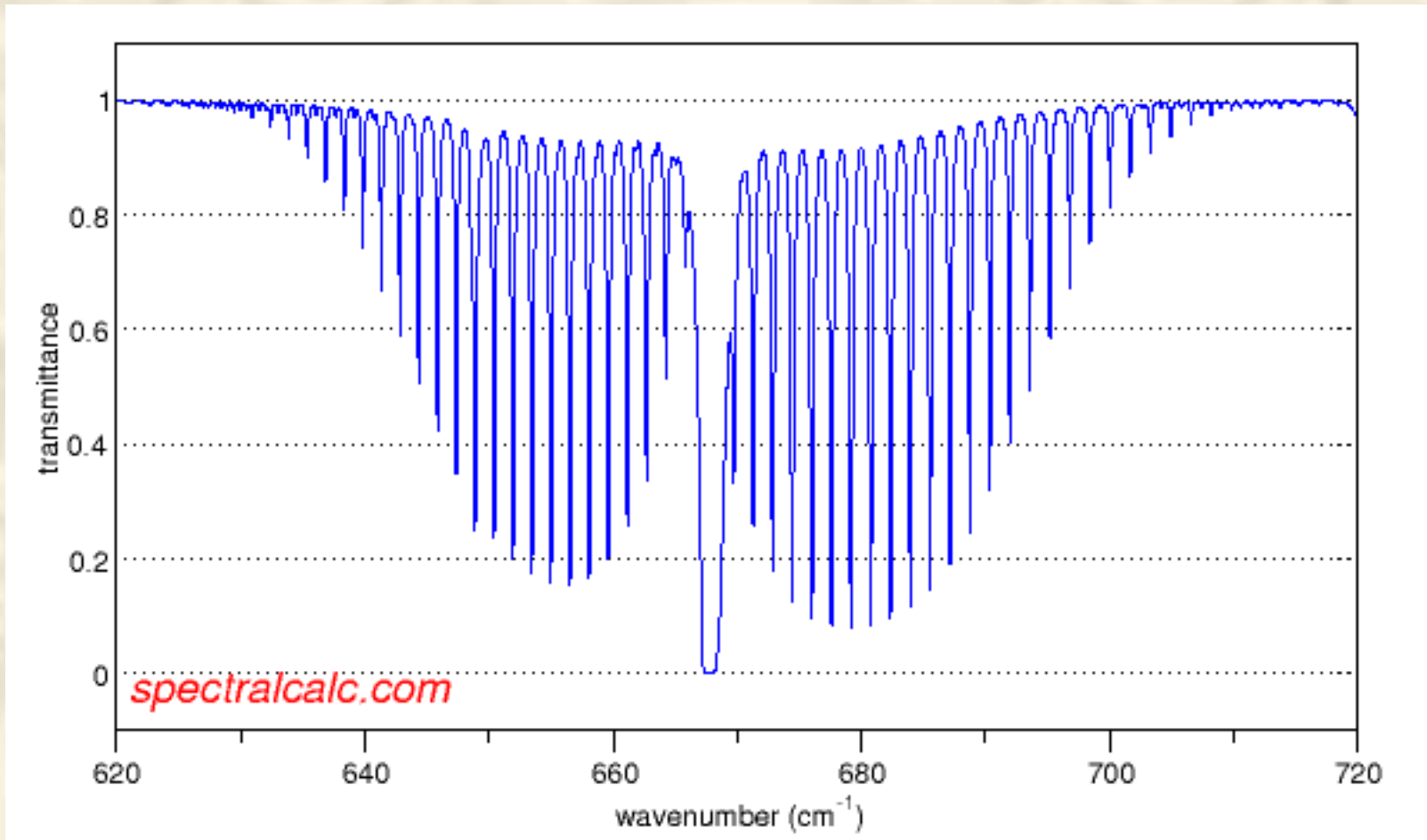


Carbon dioxide (CO₂)

Fundamental modes:



Transmittance spectrum for CO₂



<http://www.spectralcalc.com/calc/spectralcalc.php>

Infra-red absorption

Water vapor (H_2O)

3 vibrational fundamental modes

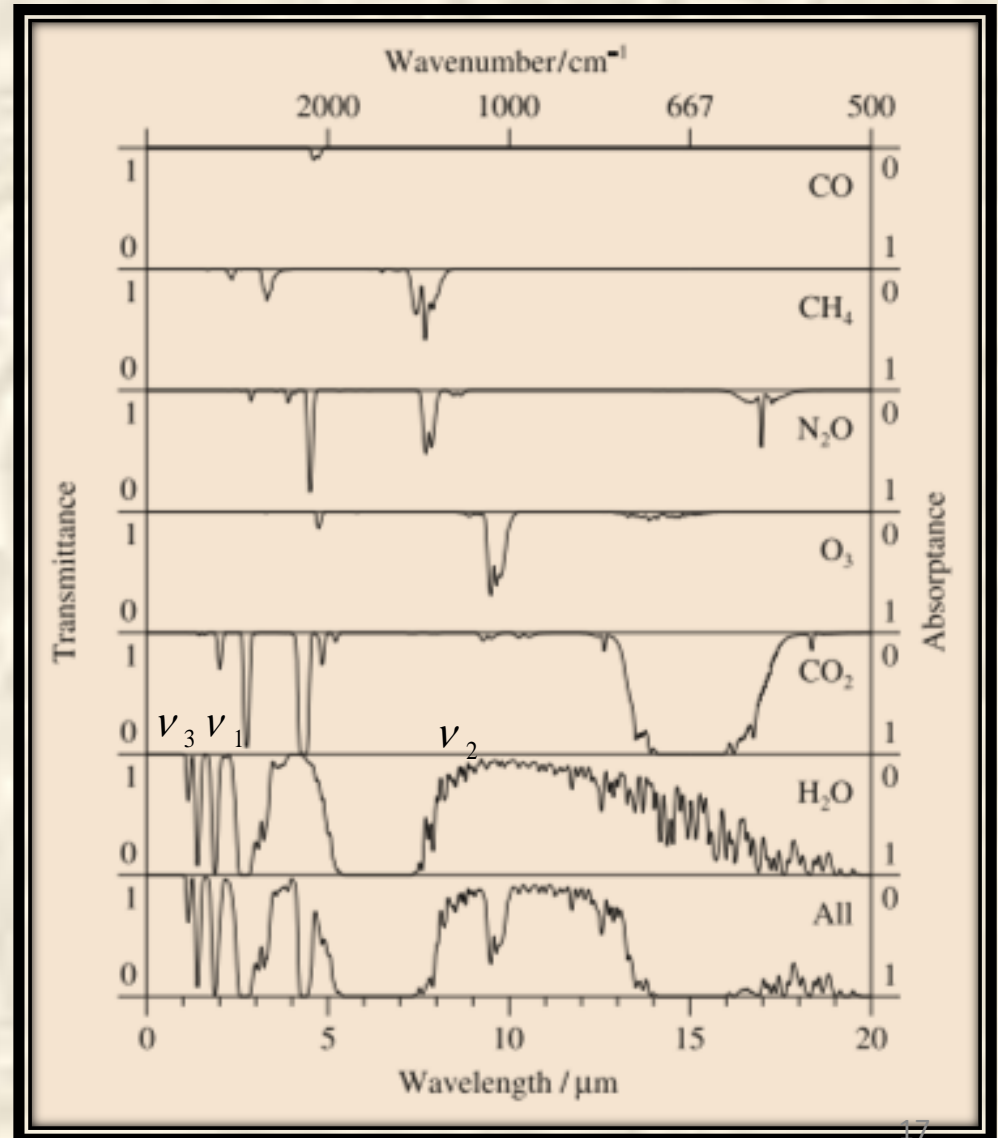
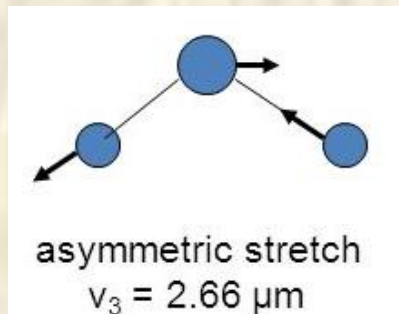
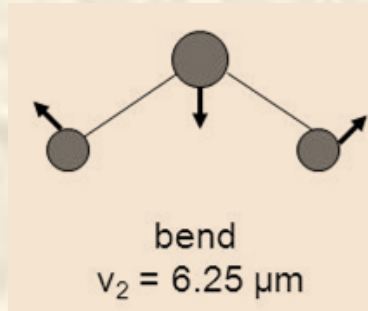
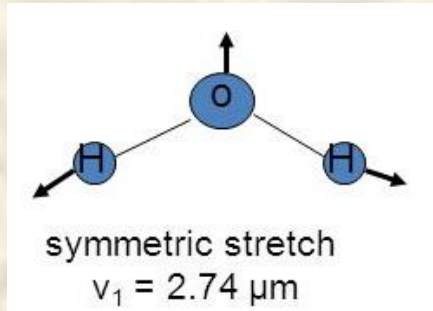
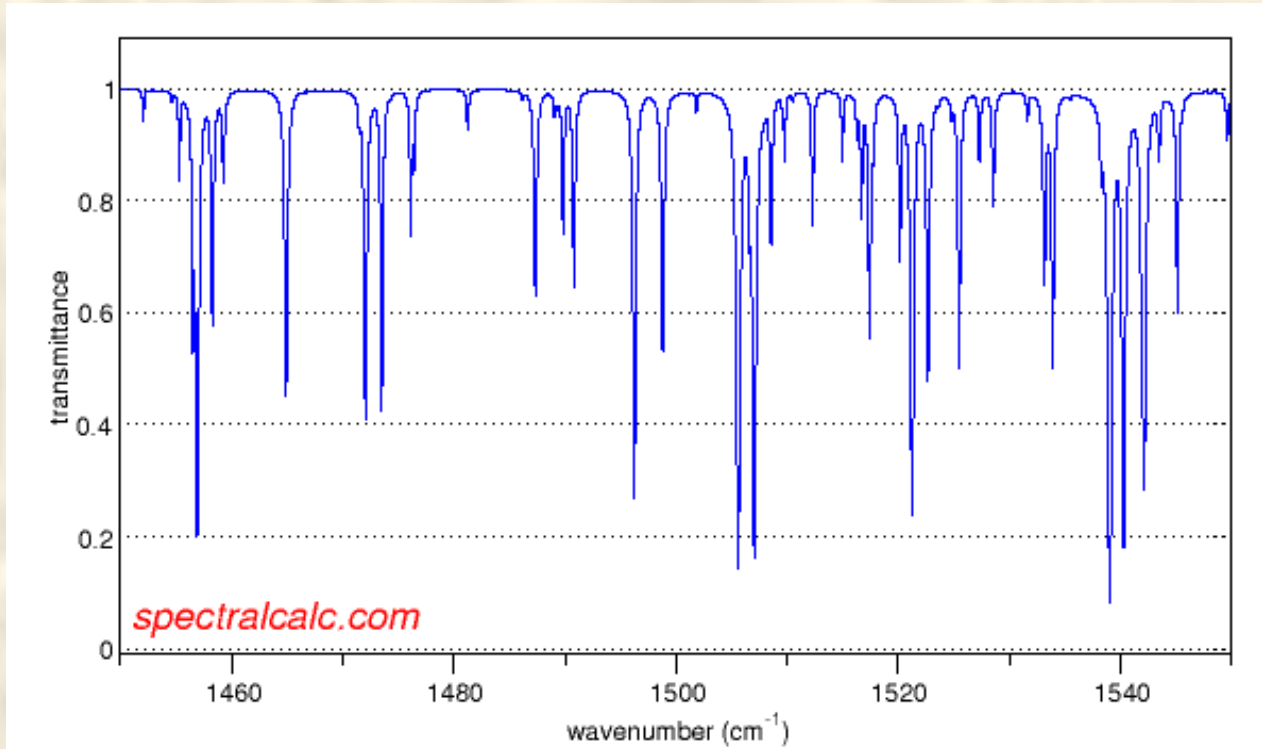


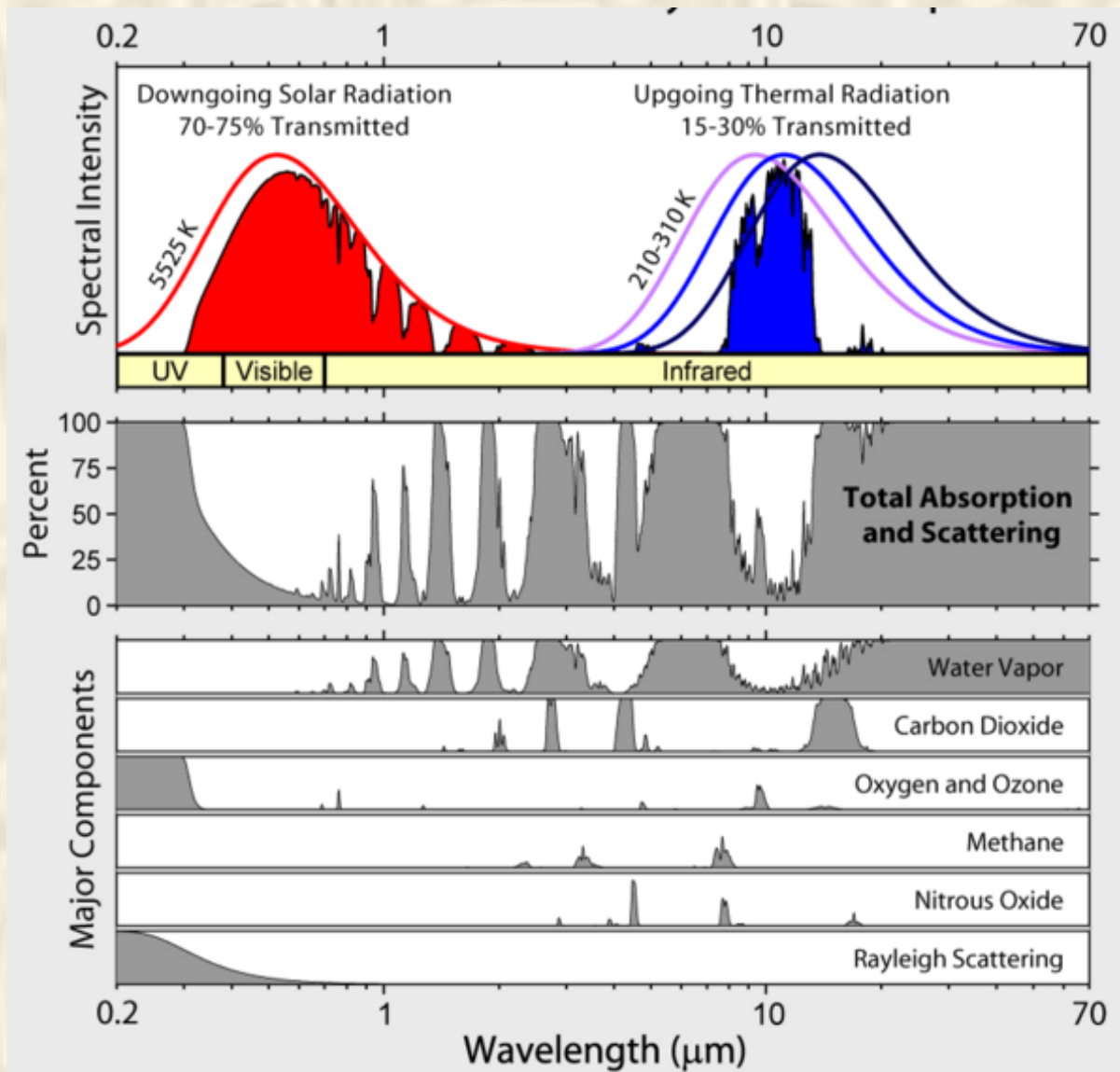
Fig. 3.14

Transmittance spectrum for H₂O



<http://www.spectralcalc.com/calc/spectralcalc.php>

Radiation Transmitted by the Atmosphere



Ultra-violet absorption

The absorption cross-section for O_2 has large values due to ionisation at Wavelengths below 100 nm in the range 100-300 nm there are irregular Bands of unknown origin.

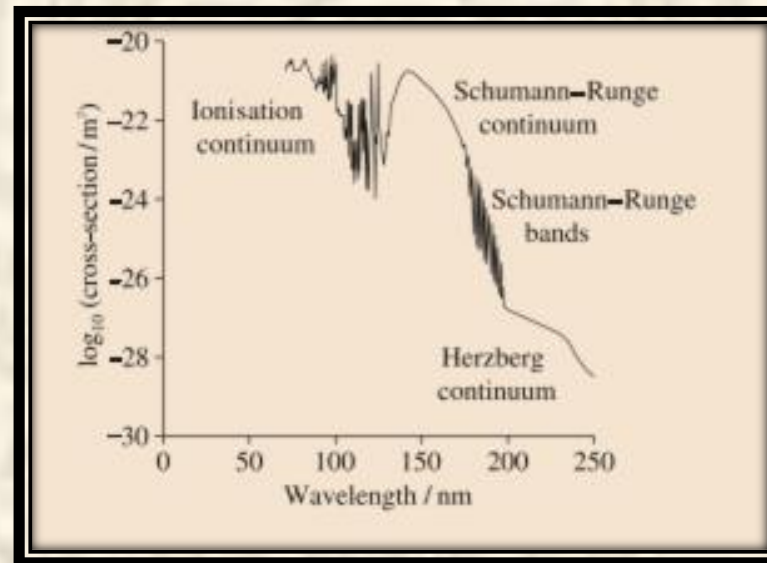
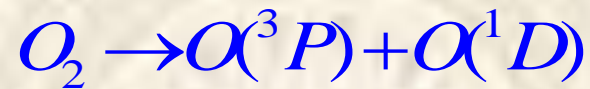
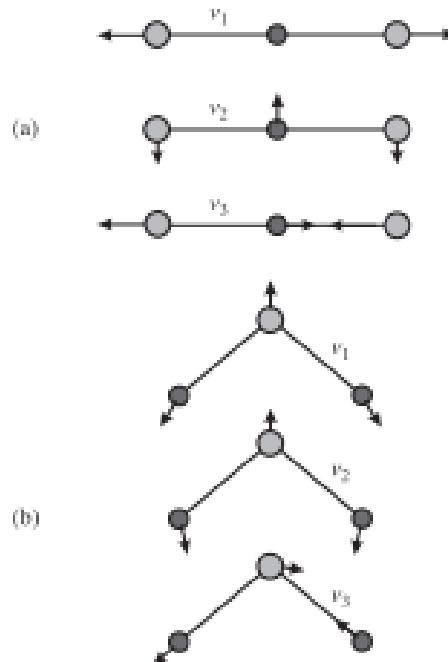
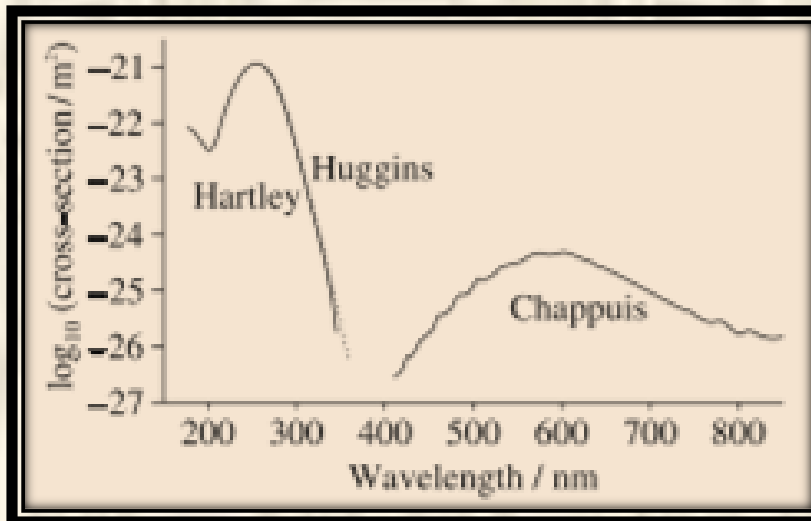


Fig. 3.15

General shape of the absorption cross-section as a function of wavelength for O_2

Fig. 3.16

The absorption cross-section as a function of wavelength for O_3 . Details of the line structure of the Huggins band have been suppressed. In the Huggins band the solid line corresponds to a temperature of 203 K and the dashed line to a temperature of 273 K.



The altitude of unit optical depth for vertical solar radiation.
The principal absorption bands are shown

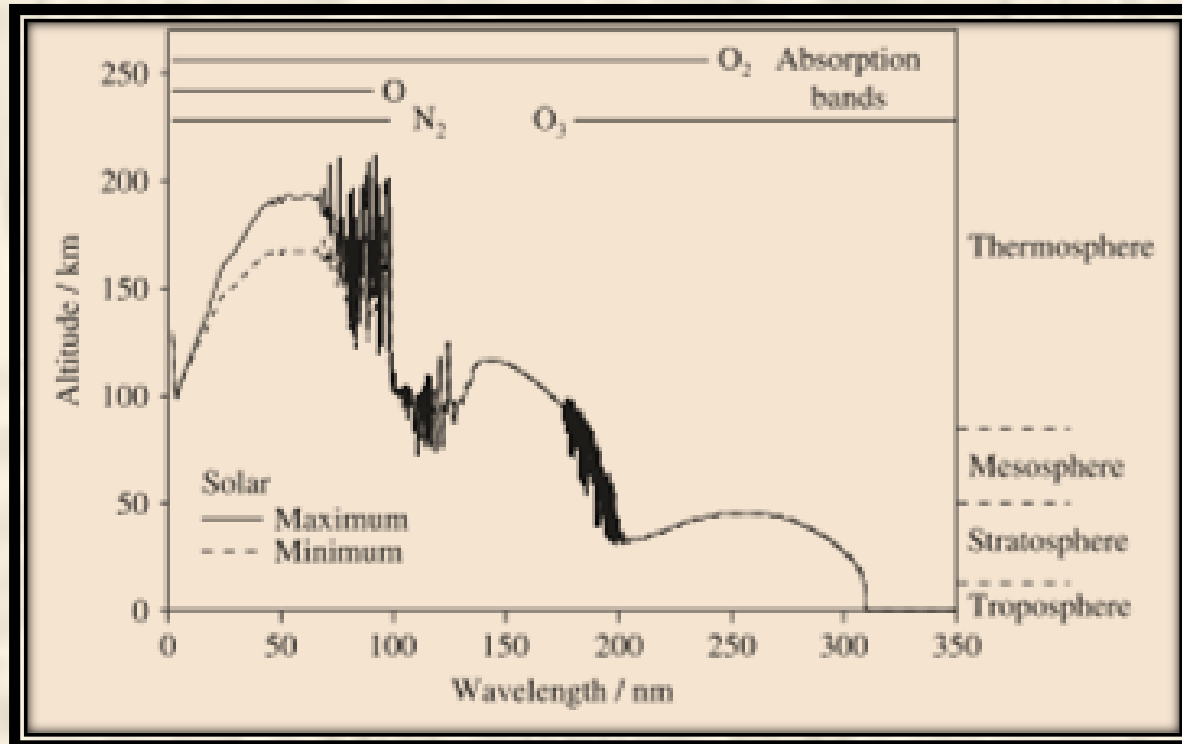


Fig. 3.17

Heating rates

Basic ideas

$$AF_z(z)$$

$$A\Delta z$$

$$AF_z(z + \Delta z)$$

$$A[F_z(z) - F_z(z + \Delta z)] \approx -(A\Delta z)dF_z / dz$$

$$-dF_z / dz$$

$$Q = -\frac{1}{\rho(z)} \frac{dF_z}{dz}$$

$$Q / c_p$$

$$F_z (= F^\uparrow - F^\downarrow)$$

Short-wave heating

$$\rho Q_v^{sw} \quad \rho_a z$$

$$\chi_v(z) = \int_z^\infty k_v(z') \rho_a(z') dz'$$

$$F_v^\downarrow(z) = F_{\infty}^\downarrow e^{-\chi_v(z)}$$

$$F_{\infty}^\downarrow \quad e^{-\chi_v(z)} \quad \tau_v(z, \infty)$$

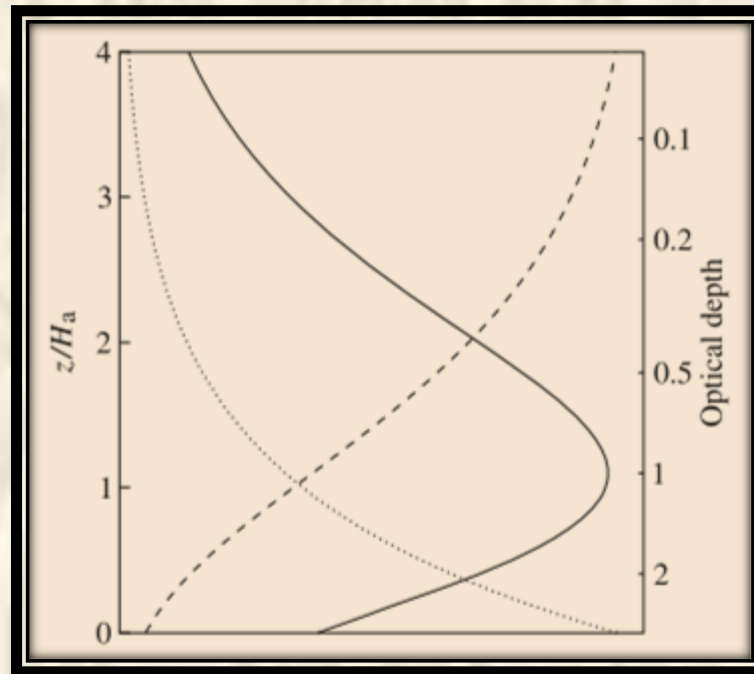
$$F_{z_v}(z) = -F_{\infty}^\downarrow e^{-\chi_v(z)}$$

$$\begin{aligned} \rho Q_v^{sw} &= \frac{d}{dz} (F_{\infty}^{\downarrow} e^{-\chi_v(z)}) = F_{\infty}^{\downarrow} \left(-\frac{d\chi_v}{dz} \right) e^{-\chi_v(z)} \\ &= F_{\infty}^{\downarrow} k_v(z) \rho_a(z) e^{-\chi_v(z)} \end{aligned}$$

$$\rho_a(z) = \rho_a(0) e^{-z/H_a}$$

$$\chi_v(z) = H_a k_v \rho_a(0) e^{-z/H_a} = \chi_v(0) e^{-z/H_a}$$

$$F_{zV} = -F_{\infty}^{\downarrow} e^{-\chi_v(0)} e^{-z/H_a}$$



$$\rho Q_v^{sw}(z) = F_{\infty}^{\downarrow} k_v \rho_a(0) e^{-z/H_a - \chi_v(0)} e^{-z/H_a}$$

Long-wave heating and cooling

$$F_{\nu}^{\uparrow}(z) = \pi \int_0^z B_{\nu}(z') \frac{\partial \tau_{\nu}^*(z', z)}{\partial z'} dz' + \pi B_{\nu}(0) \tau_{\nu}^*(0, z)$$

$$\tau_{\nu}^*(z', z) \quad B_{\nu}(0) \quad J_{\nu} = B_{\nu}$$

$$F_{\nu}^{\downarrow}(z) = -\pi \int_z^{\infty} B_{\nu}(z') \frac{\partial \tau_{\nu}^*(z', z)}{\partial z'} dz'$$

$$F_{z\nu}(z) = F_{\nu}^{\uparrow}(z) - F_{\nu}^{\downarrow}(z)$$

$$Q_{\nu}^{lw} \quad k_{\nu} \rho_a J_{\nu} A \Delta z$$

$$\tau_v(z, \infty) = \exp\left(-\int_z^{\infty} k_v \rho_a dz'\right)$$

$$\frac{\partial \tau_v(z, \infty)}{\partial z} = k_v(z) \rho_a(z) \tau_v(z, \infty)$$

$$B_v(z) \frac{\partial \tau_v(z, \infty)}{\partial z} A \Delta x$$

$$\tau_v \quad \tau_v^*$$

$$Q_v^{cts}(z) = \frac{\pi B_v(z)}{\rho(z)} \frac{\partial \tau_v^*(z, \infty)}{\partial z}$$

$$Q_v^{lw} \approx Q_v^{cts}$$

$$Q_v^{lw}(z)$$

$$Q_v^{lw}(z)$$

$$\tau_v^*$$

Net radiative heating rates

$$Q^{sw} / c_p$$

$$-Q^{lw} / c_p$$

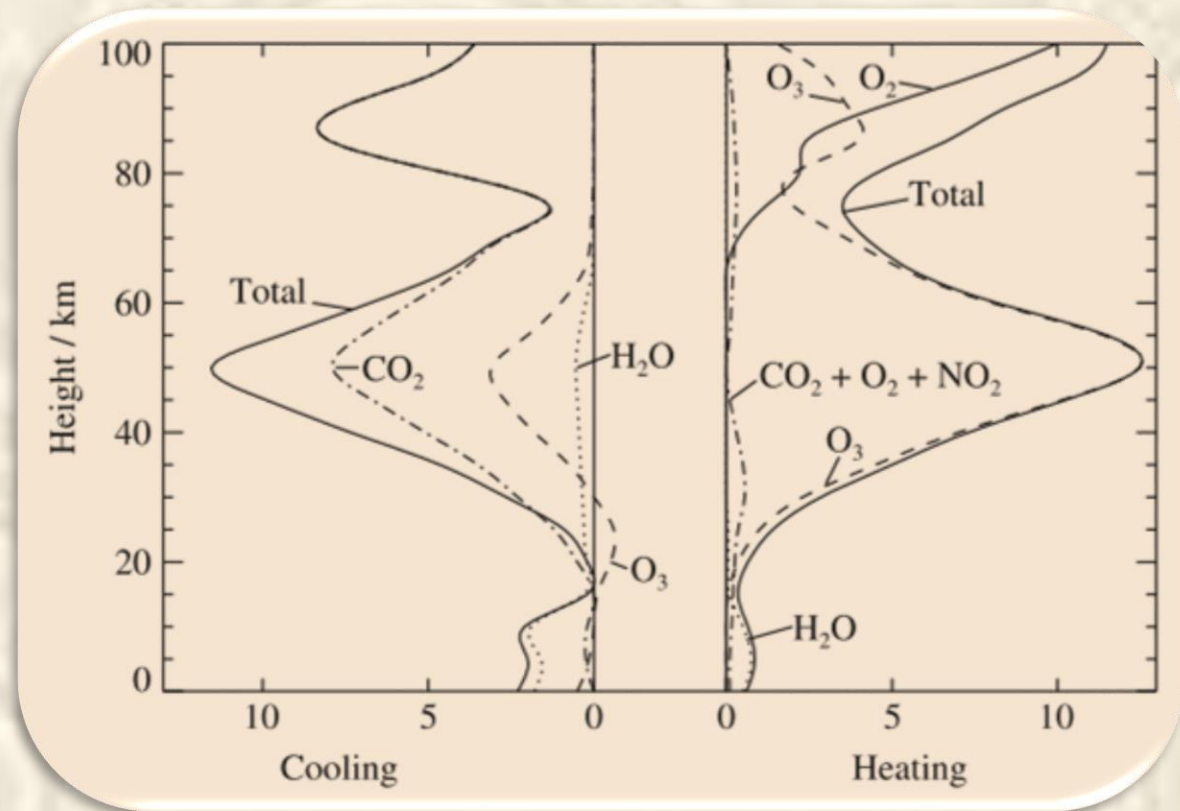
$$Q = Q^{sw} + Q^{lw}$$

$$Q = 0$$

$$T_r(r)$$

$$Q(T_r(r)) = 0$$

$$T = T_r + \delta T$$



$$Q(T_r + \delta T) \approx Q(T_r) + \delta T \left. \frac{\partial Q}{\partial T} \right|_{T=T_r} = \delta T \left. \frac{\partial Q}{\partial T} \right|_{T=T_r}$$

$$Q(T_r) = 0 \qquad = -c_p \frac{\delta T}{\tau_r}$$

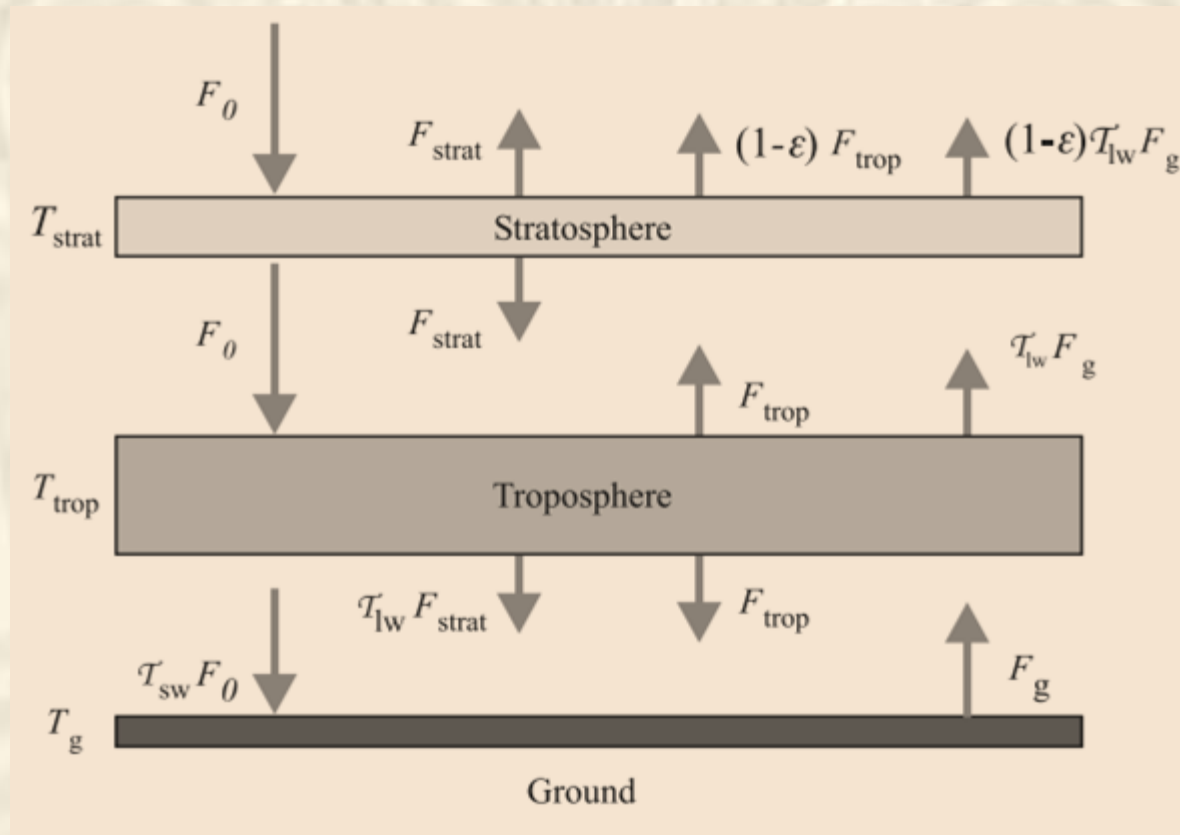
$$\tau_r = c_p \left(\left. \frac{\partial Q}{\partial T} \right|_{T=T_r} \right)^{-1}$$

The greenhouse effect revisited

Two-layer atmosphere in radiative equilibrium, including an optically thin stratosphere

$$T_{trop} \quad \tau_{sw} \quad \tau_{lw}$$

$$T_c \equiv \left(\frac{F_0}{\sigma}\right)^{1/4} \approx 255 K$$



$$F_0 = F_{\text{strat}} + (1 + \varepsilon)(F_{\text{trop}} + \tau_{\text{lw}}F_g)$$

$$F_{\text{strat}} = \sigma \varepsilon T_{\text{strat}}^4, \quad F_{\text{trop}} = \sigma (1 - \tau_{\text{lw}}) T_{\text{trop}}^4, \quad F_g = \sigma T_g^4$$

$$F_0 + F_{strat} = F_{trop} + \tau_{lw} F_g$$

$$2F_{strat} = \varepsilon(F_{trop} + \tau_{lw} F_g)$$

$$F_{trop} + \tau_{lw} F_g$$

$$F_0 + F_{strat} = (1 - \varepsilon)(F_0 + F_{strat})$$

$$\sigma \varepsilon T_{strat}^4 = F_{strat} = \frac{\varepsilon F_0}{2 - \varepsilon}$$

$$\varepsilon \ll 1 \quad \sigma T_{strat}^4 \approx \frac{F_0}{2} \quad T_{strat} \approx \frac{T_c}{2^{1/4}} = 214K$$

$$F_{trop} + \tau_{lw} F_g$$

$$F_{trop} = \frac{2F_0}{2 - \varepsilon} - \tau_{lw} F_g$$

$$\tau_{sw} F_0 + F_{lw} F_{strat} + F_{trop} = F_g$$

Continuously stratified atmosphere in radiative equilibrium

$$-\frac{dF^\uparrow}{d\chi^*} + F^\uparrow = \pi B(T)$$

$$\pi B(T) = \sigma T^4$$

$$\frac{dF^\downarrow}{d\chi^*} + F^\downarrow = \pi B(T)$$

$$Q^{sw} = 0 \quad Q^{lw} = 0$$

$$F_z = F^\uparrow - F^\downarrow = \text{constant} \quad F^\downarrow(0) = 0 \quad F_z = F^\uparrow(0)$$

$$F_z = F^\uparrow - F^\downarrow = F_0$$

$$-\frac{d}{d\chi^*} (F^\uparrow - F^\downarrow) + F^\uparrow - F^\downarrow = 2\pi B(T)$$

$$\pi B(T) = \frac{1}{2} (F^\uparrow + F^\downarrow)$$

$$\frac{d}{d\chi^*} (F^\uparrow + F^\downarrow) = F^\uparrow - F^\downarrow = F_0$$

$$F^{\uparrow} + F^{\downarrow} = F_0 \chi^* + \text{constant}$$

$$F^{\uparrow} + F^{\downarrow} = F_0(1 + \chi^*)$$

$$F^{\uparrow} = \frac{1}{2} F_0(2 + \chi^*)$$

$$F^{\downarrow} = \frac{1}{2} F_0 \chi^*$$

$$\pi B(T) = \sigma T^4 = \frac{1}{2} F_0(1 + \chi^*)$$

$$F_0(1 + \chi_g^*)/2$$

$$\pi B(T_g) = \sigma T_g^4$$

$$\sigma T_g^4 = F_0 \left(1 + \frac{1}{2} \chi_g^*\right) = \sigma T_c^4 \left(1 + \frac{1}{2} \chi_g^*\right)$$

$$T_c \approx 255K \quad \chi_g^* > 0$$

$$T_g > T_c$$

$$\rho_c(z) = \rho_a(0) e^{-z/H_a}$$

$$\chi^*(z) = \chi_g^* e^{-z/H_a}$$

$$F^\uparrow(z) = \frac{1}{2} F_0 (2 + \chi_g^* e^{-z/H_a}) \quad F^\downarrow(z) = \frac{1}{2} F_0 \chi_g^* e^{-z/H_a}$$

$$T(z) = \left[\frac{F}{2\sigma_0} (1 + \chi_g^* e^{-z/H_a}) \right]^{1/4}$$

$$z / H_a$$

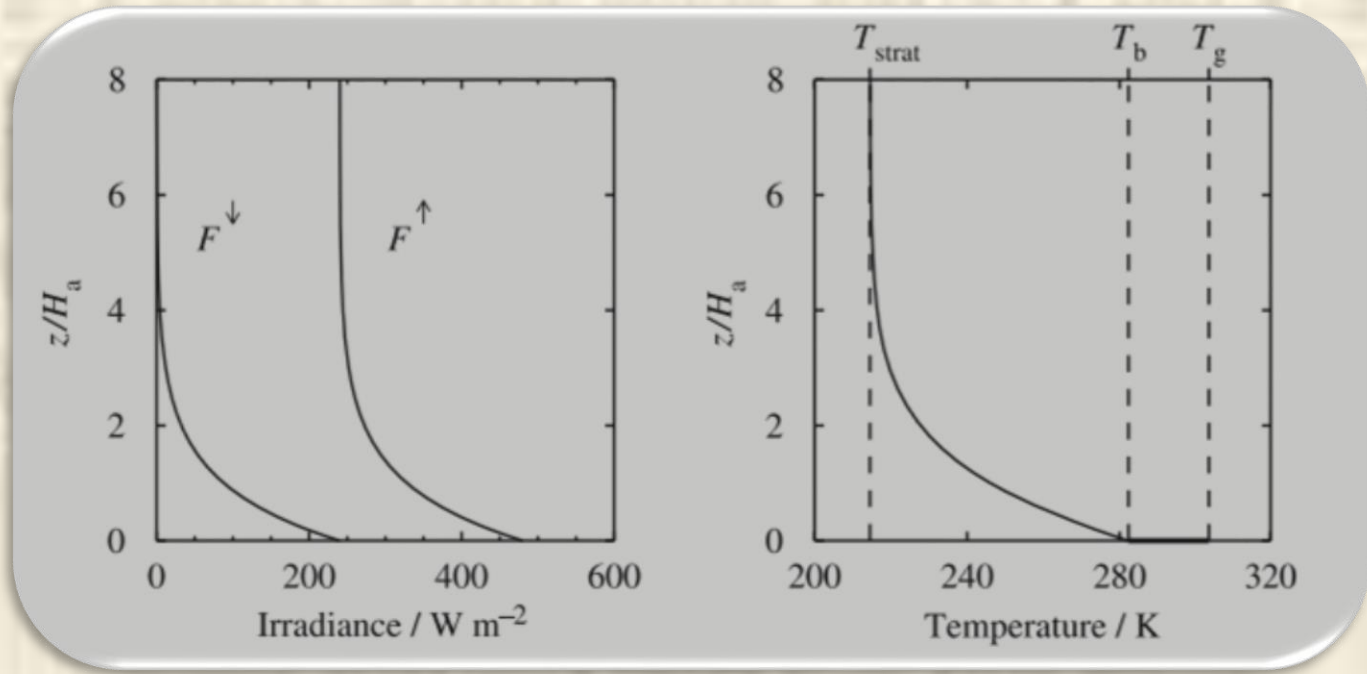
$$\chi_g^* = 2$$

$$F_0 = 240 \text{ W/m}^2$$

$$T \rightarrow \left(\frac{F}{2\sigma}\right)^{1/4} \text{ as } z \rightarrow \infty$$

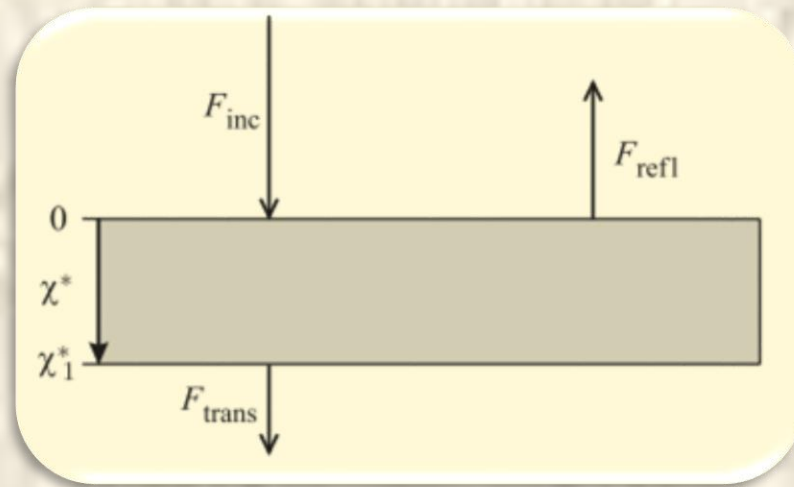
$$T_{strat} = 2^{-1/4} T_c$$

$$T(z) \rightarrow T_b \equiv T_c \left(\frac{1 + \chi_g^*}{2}\right)^{1/4} \text{ as } z \downarrow 0$$



$$T_g \equiv T_c \left(\frac{2 + \chi_g^*}{2} \right)^{1/4}$$

A simple model of scattering



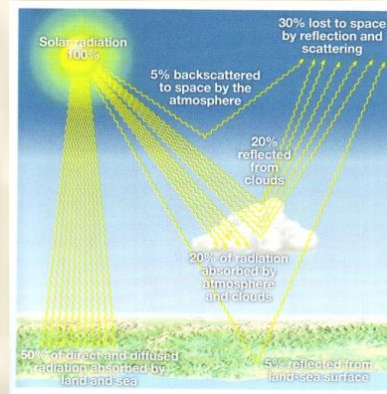
Shortwave Radiation

$S_0 = 1368 \text{ w m}^{-2}$ is the **solar constant** for Earth

Insolation

$$R_0 = S_0 \left(\frac{d_m}{d} \right)^2 \cos \gamma$$

$$I_0 = \int_{t_1}^{t_2} R_0(t) dt$$



Stefan-Boltzmann Law

This law expresses the rate of radiation emission per unit area

$$R = \sigma T^4 \quad \sigma = 5.67 \times 10^{-8} \text{ W / m}^2 \text{ K}^4$$

Compare the difference between the radiation emission from the sun and the Earth.

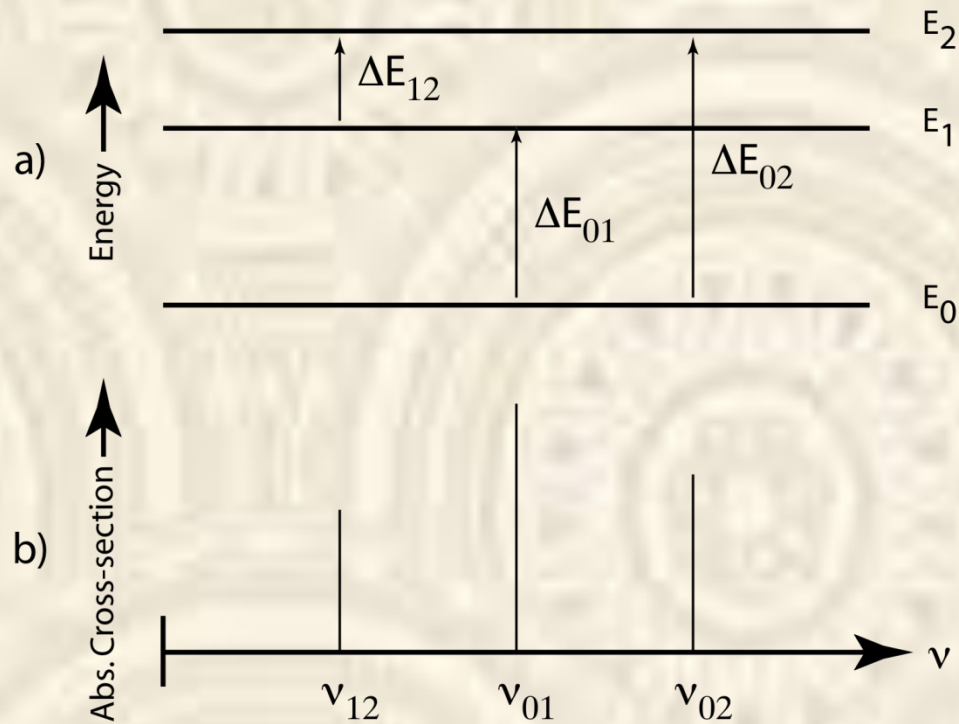
The sun with an average temperature of 6000 K emits 73,483,200 W/m²

By contrast, Earth with an average temperature of 300 K emits 459 W/m²

The sun has a temperature 20 times higher than Earth and thus emits about 160,000 times more radiation

This makes sense, $20^4 = 160,000$

Absorption spectra of molecules



Hypothetical molecule
with three allowed
energy levels

Note relationship to
emission!

$$\nu_{ij} = \Delta E_{ij}/h$$

a) allowed transitions

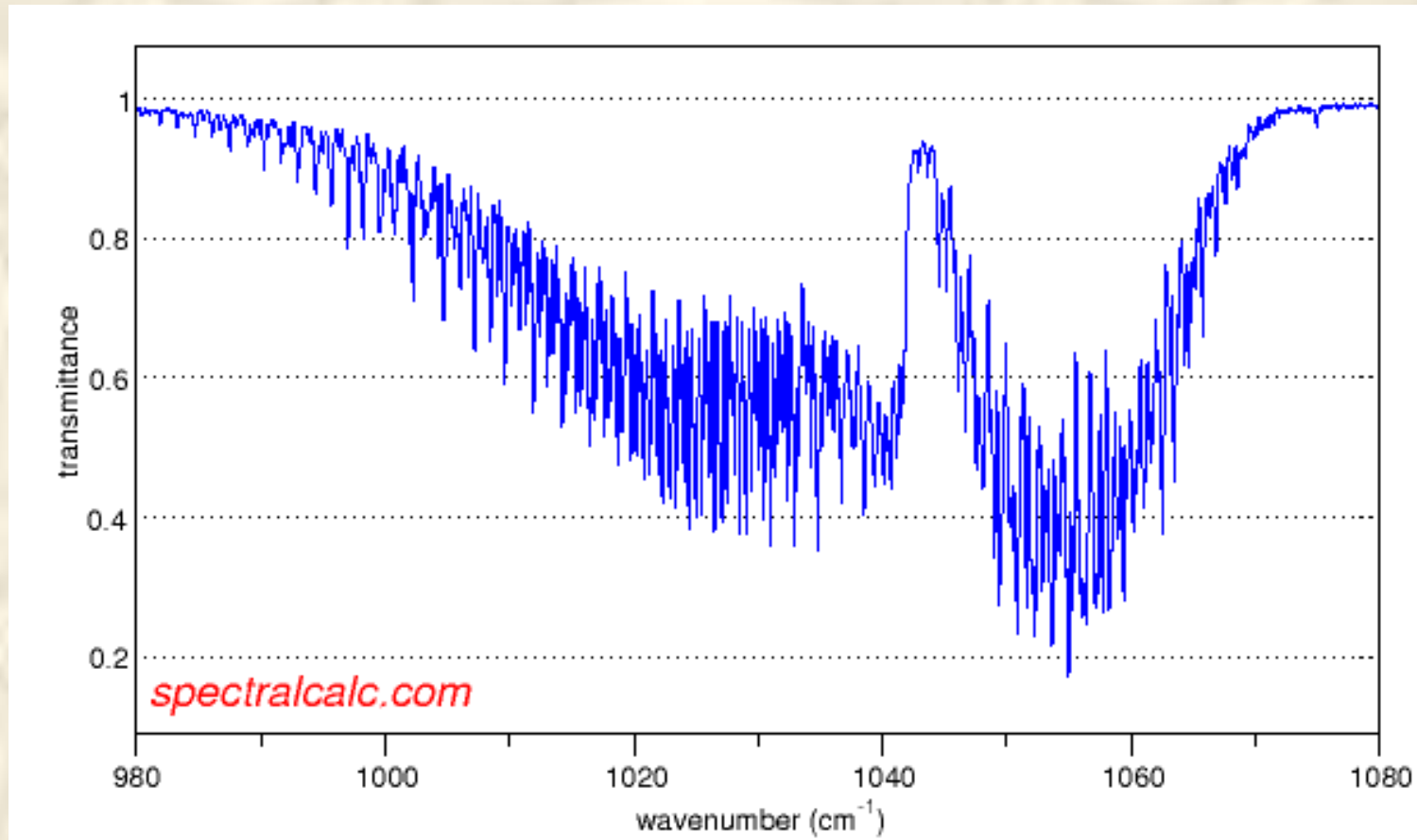
b) positions of the absorption lines in the spectrum of the molecule

Line positions are determined by the **energy changes** of allowed transitions

Line strengths are determined by the **fraction of molecules** that are in a particular initial state required for a transition

Multiple **degenerate** transitions with the same energy may combine

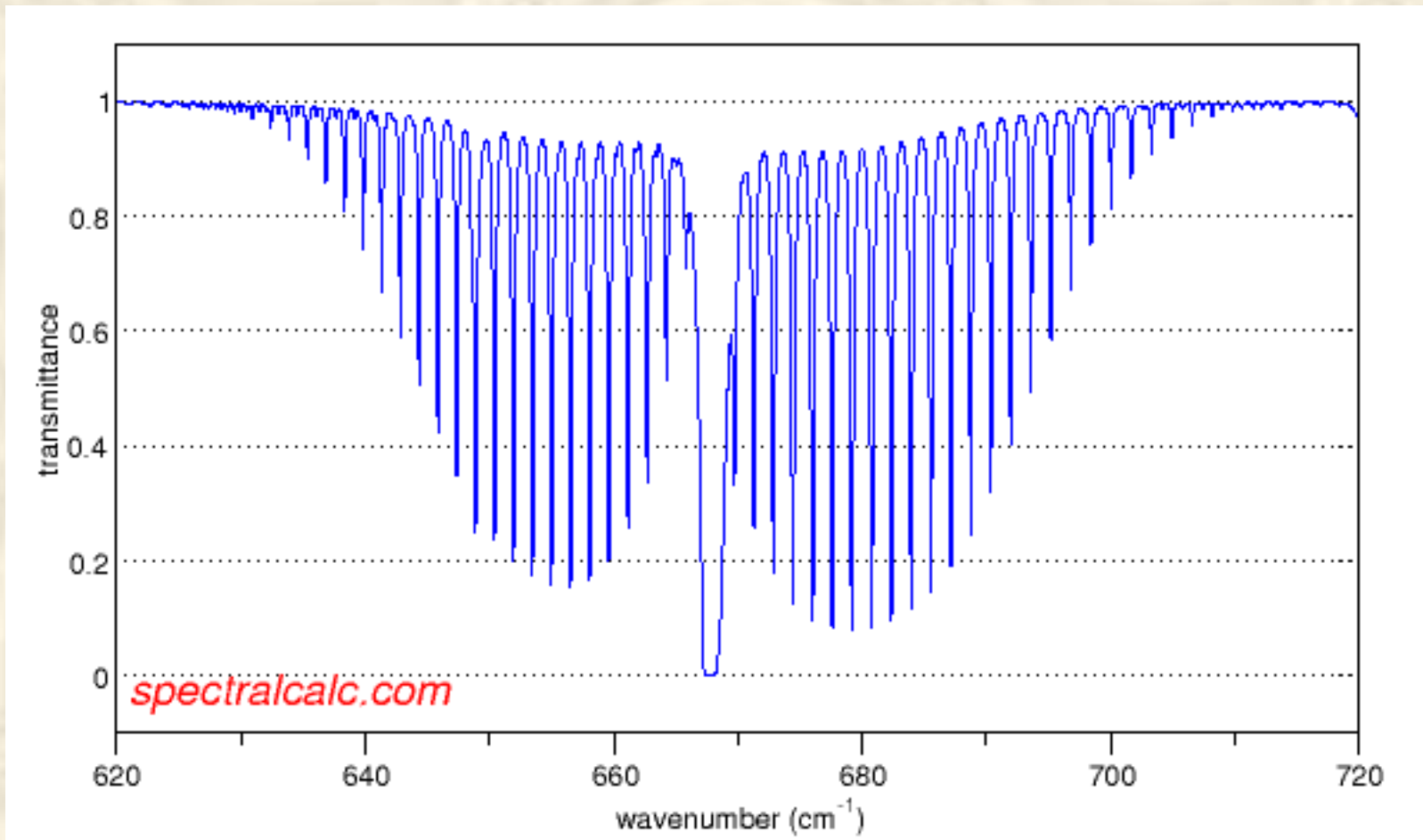
Transmittance spectrum for ozone (O_3)



<http://www.spectralcalc.com/calc/spectralcalc.php>

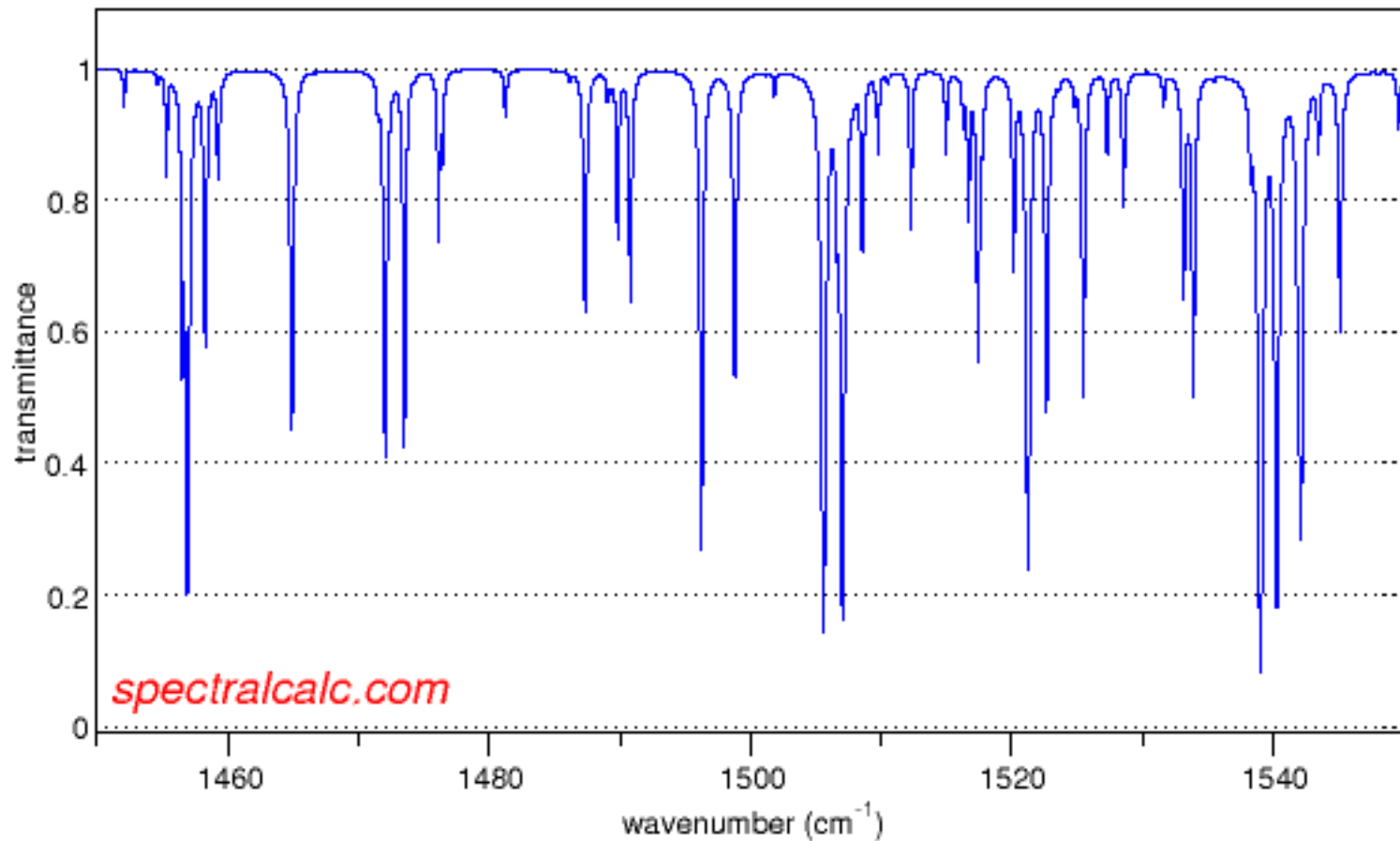
Transmittance spectrum for CO₂

<http://www.spectralcalc.com/calc/spectralcalc.php>



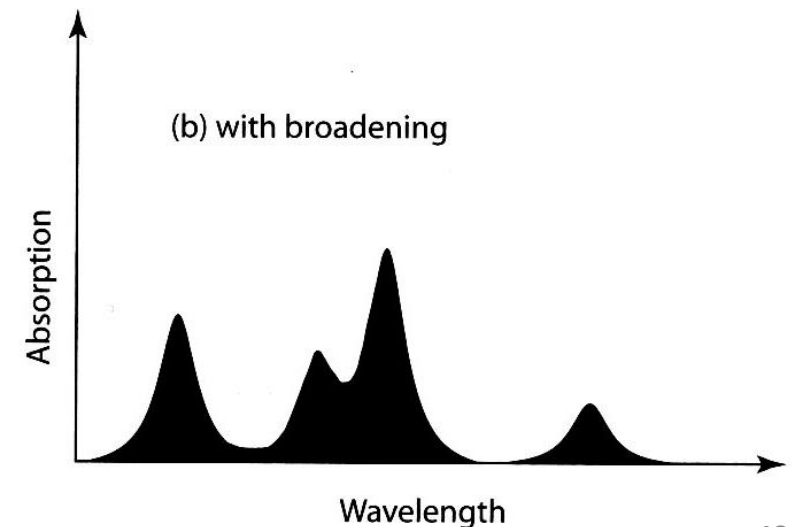
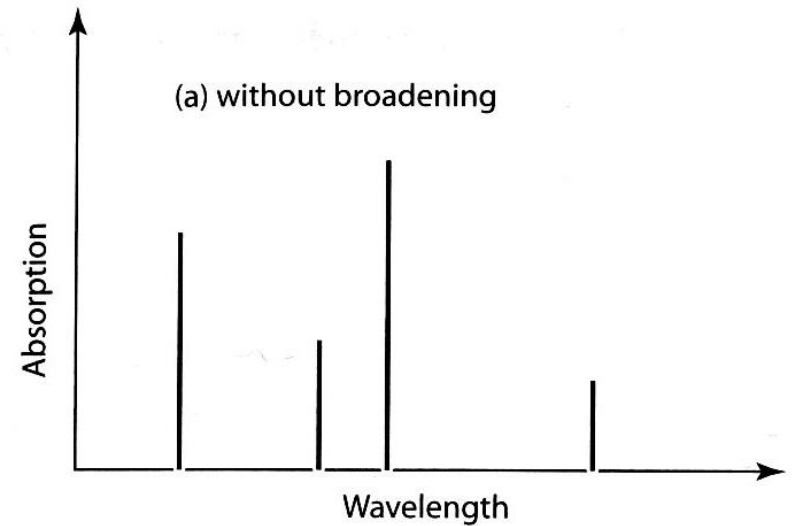
Transmittance spectrum for H₂O

<http://www.spectralcalc.com/calc/spectralcalc.php>



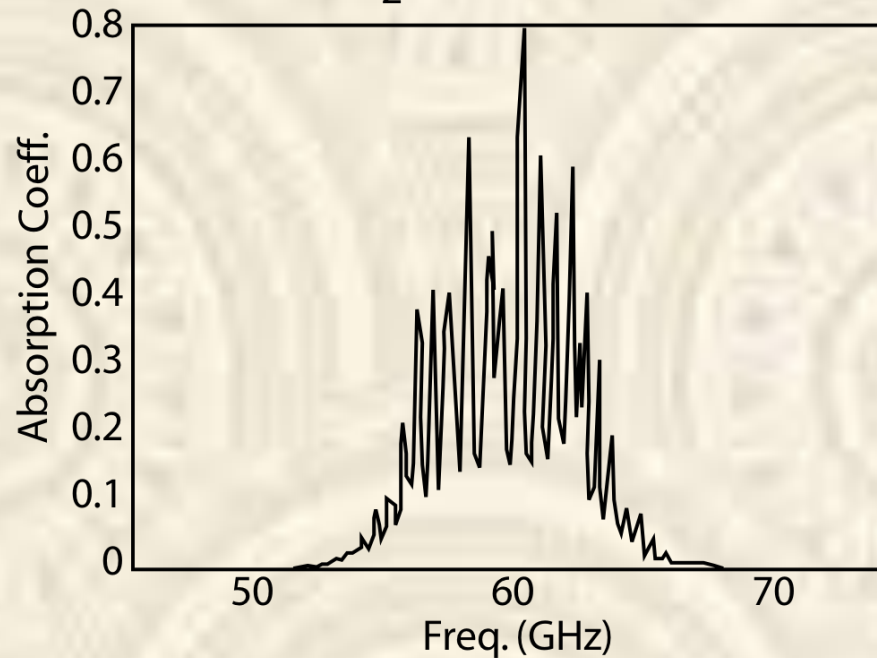
Absorption line shapes

- **Doppler broadening:** random translational motions of individual molecules in any gas leads to Doppler shift of absorption and emission wavelengths (important in upper atmosphere)
- **Pressure broadening:** collisions between molecules randomly disrupt natural transitions between energy states, so that absorption and emission occur at wavelengths that deviate from the natural line position (important in troposphere and lower stratosphere)
- Line broadening closes gaps between closely spaced absorption lines, so that the atmosphere becomes opaque over a continuous wavelength range.

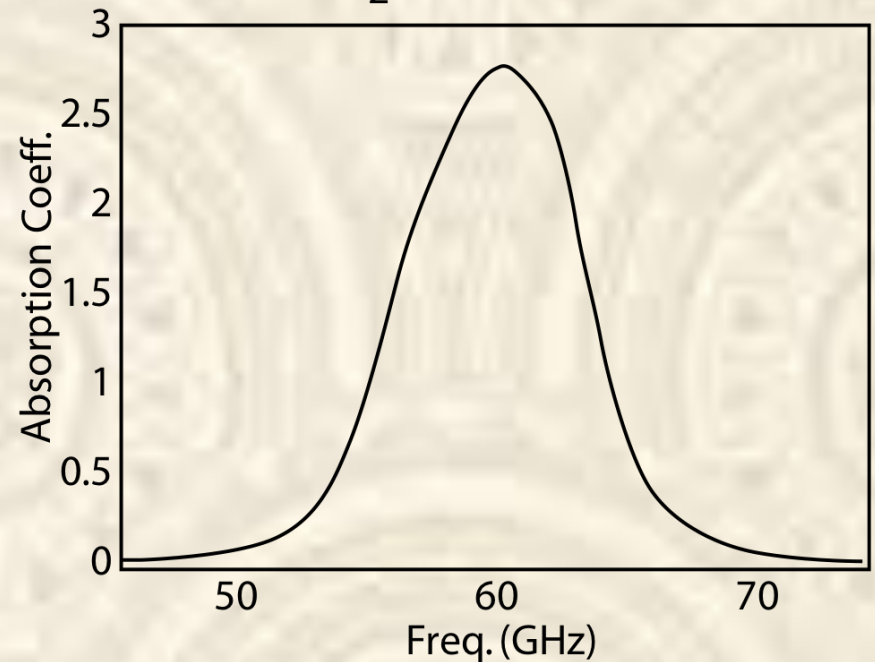


Pressure broadening

a) O₂ at 100 mb pressure

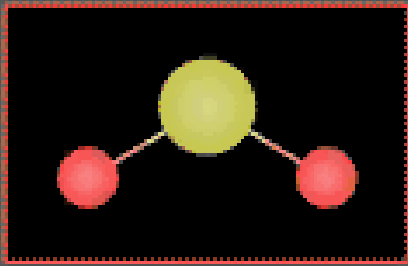


b) O₂ at 1000 mb pressure

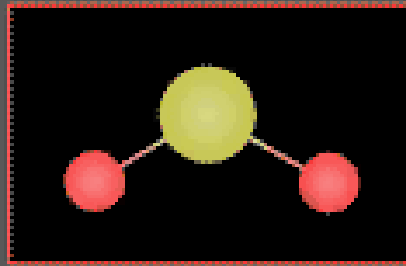


- Absorption coefficient of O₂ in the microwave band near 60 GHz at two different pressures. Pressure broadening at 1000 mb obliterates the absorption line structure.

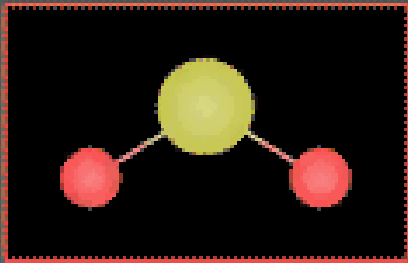
Sulfur dioxide (SO₂)



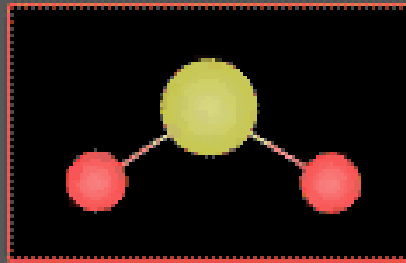
symmetric stretching



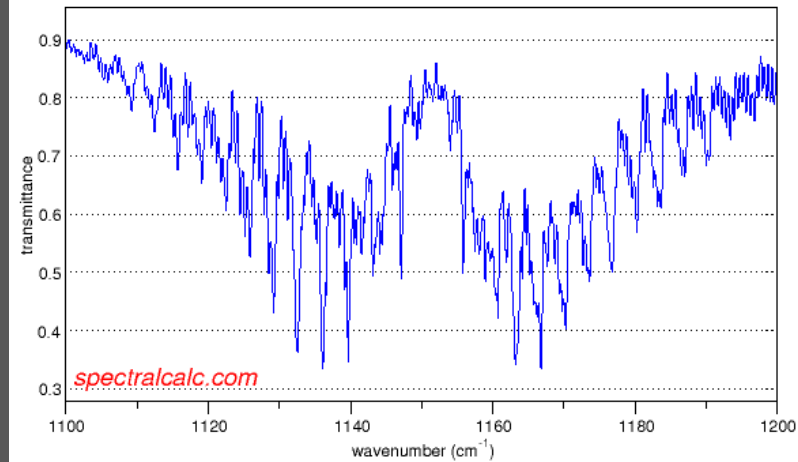
asymmetric stretching



bending

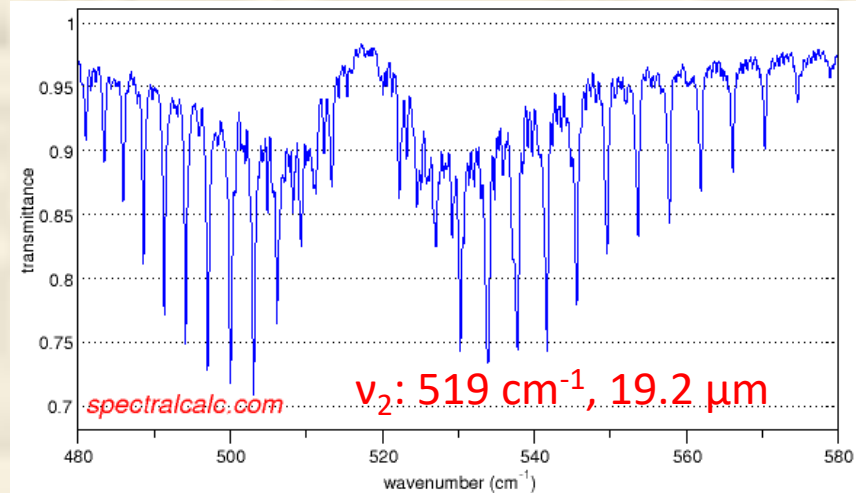


superposition

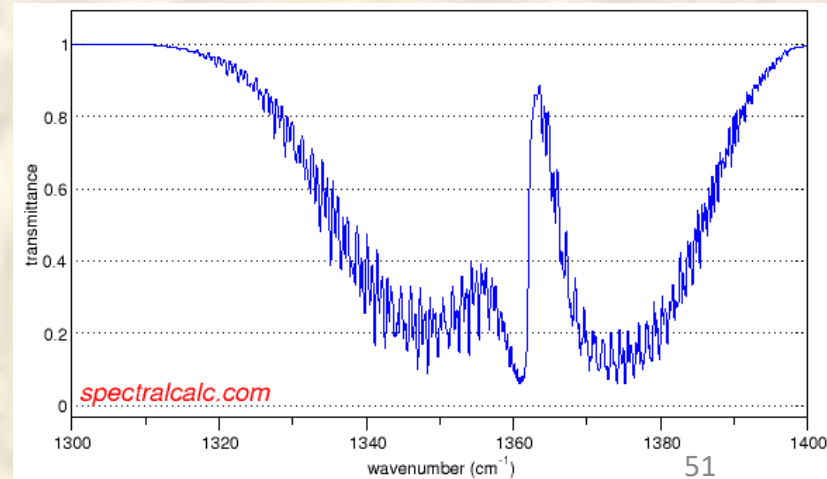


ν_1 : 1151 cm⁻¹, 8.6 μm

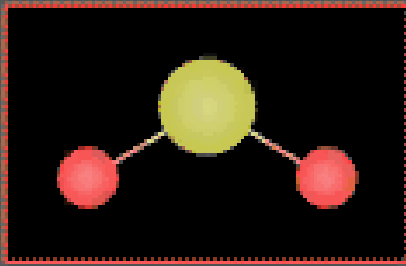
ν_3 : 1361 cm⁻¹, 7.3 μm



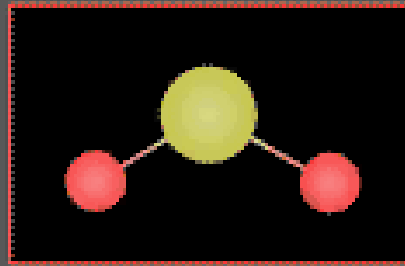
ν_2 : 519 cm⁻¹, 19.2 μm



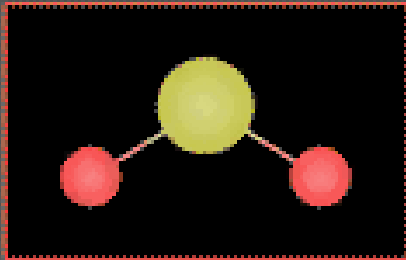
Sulfur dioxide (SO₂)



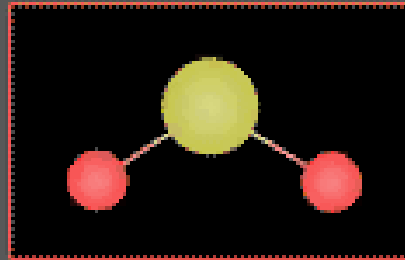
symmetric stretching



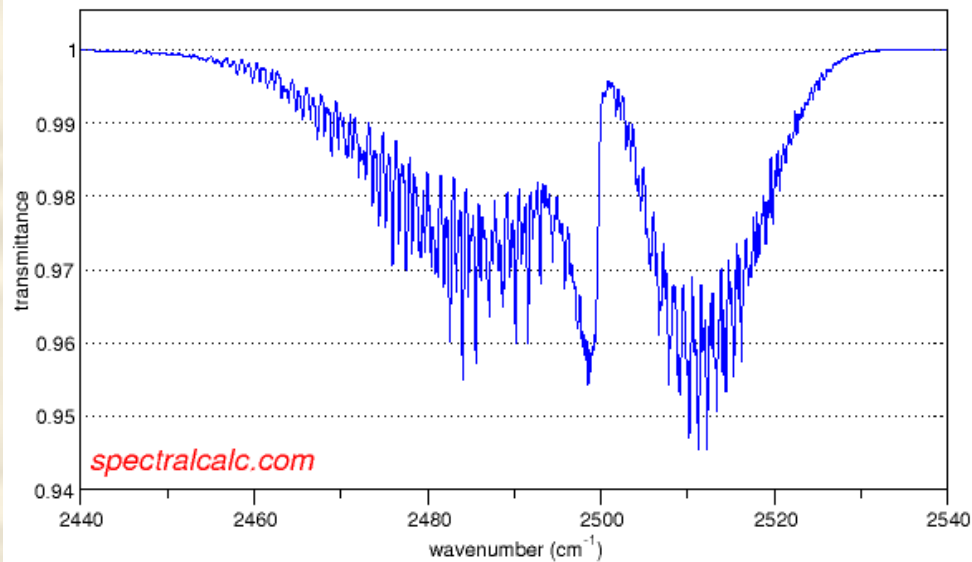
asymmetric stretching



bending

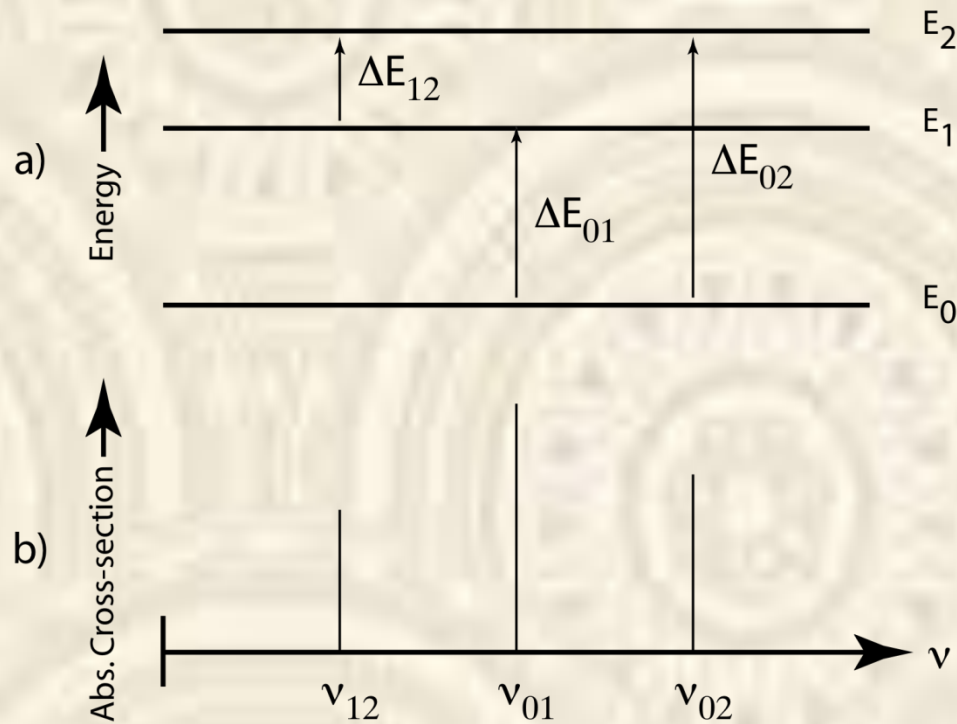


superposition



$\nu_1 + \nu_3$: 2500 cm⁻¹, 4 μm

Absorption spectra of molecules



Hypothetical molecule
with three allowed
energy levels

Note relationship to
emission!

$$\nu_{ij} = \Delta E_{ij}/h$$

a) allowed transitions

b) positions of the absorption lines in the spectrum of the molecule

Line positions are determined by the **energy changes** of allowed transitions

Line strengths are determined by the **fraction of molecules** that are in a particular initial state required for a transition

Multiple **degenerate** transitions with the same energy may combine